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**WIND ENERGY RESOURCE
EVALUATION IN A SITE OF CENTRAL
ITALY BY CFD SIMULATIONS**

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*Se un uomo sogna da solo,
il sogno rimane solo un sogno,
ma se molti uomini sognano la stessa cosa,
il sogno diventa realtà.*

Dom Helder Camara (1909-1999)

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Introduction

The present work aims to improve part of the technologies nowadays used in the production of electrical energy by renewable sources, particularly energy coming from the wind. This tendency is supported by the deeper aim, especially in the European Union, to be more and more independent from extra-communitarian Countries in fossil energy sources import and to be even more awake in protecting the environment from pollutant.

The actual world situation demonstrates how much, from an economic point of view, to dispose of primary energy sources is important. In fact, the first consequence of this subordination condition is the high cost of the energy, which is going to grow more and more due to the limited fossil energy sources, causing a risk of degeneration in the war to own the energy resources.

It is evident, nowadays, that a change in the direction of the diversification of the energy resources is needed, giving importance to the renewable energies.

Wind Energy is an ancient Energy source, used in its first applications already during Persian Emperor, more than 2000 years ago. Its use has been widely diffused in various Countries of the world and, in the present days, the wind energy technologies arrived to a mature level.

The first step in projecting a wind farm is the prediction of the wind resources and, for this scope, various techniques are used. Linear models, like Was'P (Wind Atlas Analysis and Application Program), are widely used where the terrain is characterized by simple orography. Linear models use linear equations to describe the behavior of the flow over a territory, but can encounter some problems in evaluating wind characteristics when the region becomes complex, where hills and steep mountains are present.

In these cases the use of other methods, like Computational Fluid Dynamics (CFD) is preferred [15]. CFD solves, by using numerical methods, the Reynolds Averaged Navier-Stokes (RANS) equations to predicts the flow over the site. In this thesis this calculation is done by the software WindSim.

WindSim is a software developed by the company WindSim AS of Tonsberg, Norway, which bears on the CFD software PHOENICS, to solve the RANS equations using the finite volume method and has a series of packages allowing wind energy analysis.

The present Master Thesis is the final work of a training period of 6 months spent by the author at the main office of WindSim company, during the last year of his studying in Mechanical Engineering.

The work encompasses a starting validation phase of the CFD numerical code, focused to obtain the correct parameters and boundary condition able to describe with a fair accuracy the behavior of the Atmospheric Boundary Layer (ABL).

Various problematic about wind farms projects have been afforded. A collaboration with the company Enel Produzione of Catania, Italy, has been set up and carried on, in order to perform the study of the wind energy resources in a site of central Italy, about which the company was interested in. Particularly, two short-period sets of available experimental wind data have been analysed by a procedure addressed to validate the CFD calculation over the complex orography of the site. Wind resources maps of the site were obtained and the project of a wind farm siting was proposed.

This Master Thesis is subdivided in 7 chapters: the first one describes the general characteristics of the ABL and its mathematical formulation. Moreover, a brief description of the WindSim numerical code and its graphical interface is presented. In chapter 2 and 3 the correct boundary conditions, grid and solver settings to be applied to the model in order to describe the ABL are analyzed, by means of comparison of the CFD predictions with results found in literature. Next chapter describes the wind characteristics of a site located in central Italy, together with the analysis of the experimental wind data available. Particular attention is done in finding the dominant sectors of the wind rose. Chapters 5 and 6 are related to the simulation of the flow field over the site. The “nesting technique” have been used in chapter 5 to determine the free stream velocity boundary conditions needed to describe the geostrophic wind over the site. Chapter 6 describes the geometrical and numerical model used to simulate the wind field over the whole site area. The procedure to match numerical results and of the available experimental data is

explained and performed. Finally, chapter 7 presents the wind resources map of the site and a proposal for a wind farm layout.

CHAPTER 1. Atmospheric Boundary Layer.

1.1 The Atmospheric Boundary Layer

The Atmospheric Boundary Layer (ABL) is the lowest part of the atmosphere and its behavior is directly influenced by the interaction with a planetary surface, according to Stull, R. B. (1988) [12].

It's known that the ABL responds to surface forcing in a timescale of one hour or less, while the typical space scale is few kilometers. In this layer physical quantities such as flow velocity, temperature, moisture, etc. display relatively rapid fluctuations.

Above the ABL there is the "free atmosphere", where the wind is considered geostrophic (parallel to isobars), while inside the ABL the wind is affected by surface effects and turns across the isobars.

1.2 Description of the ABL

The ABL is a part of the troposphere which is the lowest portion of Earth's atmosphere. It contains approximately 75% of the atmosphere's mass and almost all of its water vapor and aerosols.

The average depth of the troposphere is about 11 km in the middle latitudes. It is deeper in the tropical regions (up to 20 km) and shallower near the poles (about 7 km).

As already said, the lowest part of the troposphere is the ABL and this layer is about 2 km deep, depending on the landform and time of day.

In a neutral ABL, where heat transfer is not taken into account, several sub-layers can be identified.

The *canopy layer* is the first sub-layer near the terrain, where obstacles are present; above it, the *surface layer* can be found, in which Coriolis effects are negligible; finally there is the so called Ekman layer, where Coriolis effects are dominant.

1.3 Mathematical formulation of the ABL

The Atmospheric Boundary Layer can be described through very simple mathematical equations using the so called surface-layer similarity, also known as Monin-Obukhov similarity.

According to this, the speed vertical profile of the wind can be described as a function of the height by the logarithmic profile,

$$1.1) \quad u(z) = \frac{u_\tau}{VKC} \cdot \ln\left(\frac{z}{z_0}\right)$$

where,

u_τ	is the friction velocity	[m/s]
VKC	is the Von Karman constant	[]
z_0	is the roughness height	[m]

The friction velocity is calculated by the equation

$$1.2) \quad u_\tau = \sqrt{\frac{\tau_w}{\rho}}$$

where,

τ_w	is the shear stress at the wall boundary	[N/m ²]
ρ	is the air density	[kg/m ³]

An example of a velocity vertical profile, is shown in figure 1.1

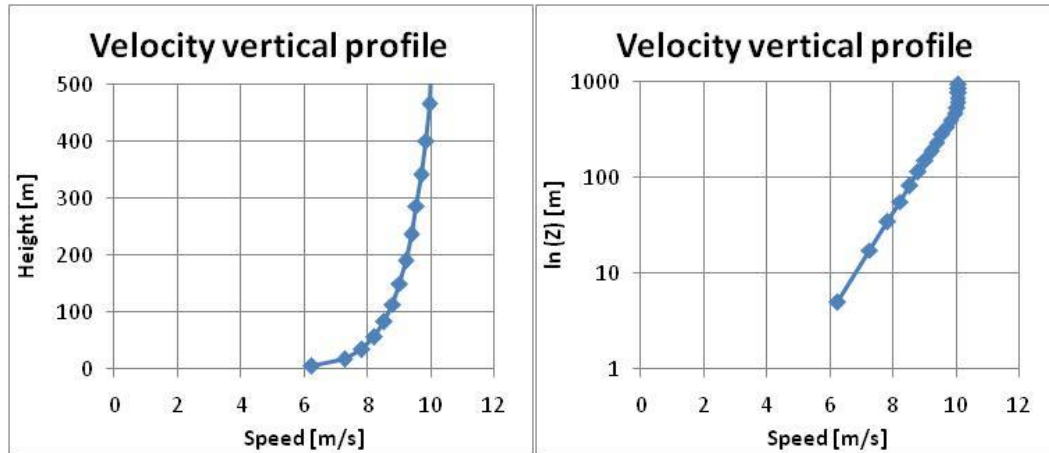


Figure 1.1: Example of an atmospheric speed vertical profile – $U=10$ m/s – $\delta=500$ m.

1.4 Navier-Stokes and RANS Equations

The most general equations that govern the fluid flow in the ABL are the Navier-Stokes equations [10]. These equations express the physical principles of conservation of mass, momentum and energy. Since a neutral boundary layer is to be analysed, no thermal effect is taken into account and the energy equation will not be solved

Conservation of mass ("continuity equation")

$$1.3) \quad \frac{\partial(u_i)}{\partial x_i} = 0$$

Conservation of momentum ("momentum equation")

$$1.4) \quad \rho \left(\frac{\partial(u_i)}{\partial t} + u_k \frac{\partial(u_i)}{\partial x_k} \right) = - \frac{\partial p}{\partial x_i} + \rho f_i + \frac{\partial}{\partial x_k} \left(\mu \frac{\partial(u_i)}{\partial x_k} \right)$$

The numerical solution of the Navier-Stokes equation is possible only if a Direct Numerical Simulation is sought. Nowadays this kind of CFD technique is applied to simple geometries and very low Reynolds numbers. Even with the today's computer power, It is only used for academic research.

Turbulence is typical of industrial and atmospheric flows. Two different ways can be pursued with the goal of solving the turbulence problem for real flows using CFD. The first method numerically solves the Large Eddy Simulation (LES) equation, while the second uses the Reynolds Averaged Navier-Stokes equation.

LES equation are deduced from Navier-Stokes Equation with a spatial averaging process. In this way only the large scale vortices are numerically solved, while the small scale turbulence is “modelled”. LES simulation were introduced [7] to solve the mesoscale atmospheric flow and today are also used for industrial flow, even though a very high computer power in terms of CPU and RAM is required.

The most used way to solve real flows with a relatively low computer cost is solving the RANS equations. These equation, displayed in equations 1.5 and 1.6, are again deduced from the N-S equation, but using a time averaging procedure instead. In this way all the turbulence is modeled and only the time averaged variables are found with the simulations.

Continuity equation

$$1.5) \quad \frac{\partial U_i}{\partial x_j} = 0$$

Momentum equation

$$1.6) \quad U_i \cdot \frac{\partial U_i}{\partial x_j} = \frac{1}{\rho} \cdot \frac{\partial}{\partial x_j} (\tau_{ij} - \rho \overline{u'_i u'_j}) - \frac{1}{\rho} \cdot \frac{\partial p}{\partial x_j} + F_i$$

In the RANS equations, terms that take account for turbulence appear. A relationship between these terms and the time averaged variables is needed to close and solve the system of RANS equations. Different turbulence models are used to close the system of RANS equations, mainly based on the Boussinesque hypothesis [6]. An example is shown by equation

$$1.7) \quad -\rho \overline{u'_i u'_j} = \mu_t \cdot \left(\frac{\partial U_i}{\partial x_j} + \frac{\partial U_j}{\partial x_i} \right) - \frac{2}{3} \left(\rho k + \mu_t \cdot \frac{\partial U_i}{\partial x_j} \right) \cdot \delta_{ij}$$

1.5 The WindSim numerical code

WindSim is a CFD code used for evaluate the wind resources in a site by solving the RANS equations. It is produced by the company WindSim AS, Norway.

WindSim uses a core constituted by the solver Phoenix (developed by Cham, UK), which solves the RANS equations with the finite volume method [11].

WindSim contains a group of accessory software which help in the solution of atmospheric flows.

WindSim is constituted by 6 modules, which have to be run in sequence: the user is not able to run the further module if the previous one hasn't completed its work.

The modules are listed below.

- 1 Terrain
- 2 Windfield
- 3 Objects
- 4 Results
- 5 Wind Resources
- 6 Energy

A brief description of each module is now reported, in order to analyze the main input parameters and output results.

1.5.1 TERRAIN

In this module all informations regarding orography and roughness of the site of interest are inserted, through the file `***.gws`. The `***.gws` file is a characteristic WindSim format that can be created by the automatic conversion module, also present in WindSim. In this way, *WASP* files and `***.xyz` files can be used to set terrain informations.

The Terrain module also allows the user to define the parameters of the geometrical model, setting the grid discretization and the number of cells. Moreover, a lot of other important settings needed for CFD simulations can be defined.

1.5.2 WINDFIELD

The Windfield module calculate the wind characteristics of the site of interest. To do this the Reynolds Averaged Navier-Stokes equations are solved, using the finite volume numerical method with the CFD solver “PHOENICS”

In this module the boundary conditions, the turbulence model and the solver settings can be defined by the user.

Boundary conditions

The wind velocity vertical profile must be set by defining the height of the ABL, the free stream velocity of the geostrophic wind and the wind direction. The ABL is described through the log law (refer to chapter 1), while the wind direction can be set by subdividing the wind rose in a specific number of sector, usually 12 in order to obtain a main wind direction every 30 degrees.

In this way, 12 CFD simulations of the flow field are needed, in which the settings for the ABL are kept constant.

The geometrical model of the flow field can be considered as a rectangular wind tunnel, with one inlet section and one outlet section. At the ground of the wind tunnel the informations on the digital terrain map are given (chapter 6.2.1), while at the top of the model the “*wall with no friction*” condition is assigned.

Concerning the choice of the wind vertical profile, the value 500 m is given to the height of the ABL. This value is the most used by scientists according to literature in order to describe the height of the geostrophic wind; concerning to the choice of the free stream velocity, the value used in this CFD simulation is the result of the test reported in the chapter 5, where 20 m/s was the best value to avoid convergence problem in a such complex site.

Turbulence model

The turbulence model chosen to run the CFD simulation was the default standard κ - ϵ model. In WindSim, it is possible to define other turbulence models, like the

($k-\epsilon$ RNG or the Low Reynolds), by manipulating the Phoenics input file. An accurate explanation on the use of Phoenics can be found in the specific technical manual [5].

Solver

In WindSim two types of solver are available, the coupled and the segregated ones. The first is a multigrid solver.

Moreover, the number of iterations can be set and a field variable to monitor the convergence level can be chosen.

1.5.3 OBJECTS

The Objects module allows the user to define all the elements needed to study a wind farm, like turbines, climatologies and transferred climatologies.

Turbines

With the turbine object the position of the machine and its characteristics (power curve and thrust curve) can be defined.

In the place of each turbine the velocity vertical profile referred to the CFD results can be extracted.

Climatologies

Using this object the wind experimental data, velocity and direction, related to an anemometric station inside the site, can be inserted. Data can be assigned by a frequency table (****.wws* file) or by a time-history (****.twc* file).

By inserting one climatology in the project, the CFD results can be weighted with the experimental data; this procedure allows to forecast the wind characteristics of the site, through the module Wind Resources which will be described in the further paragraphs.

Transferred climatologies

The transferred climatology is a virtual climatology calculated from experimental data available in one climatology already present into the site.

The transfer process is based on frequency tables (****.wws*). In other words, if a climatology is known in the position A, containing velocity and direction of the wind,

it is possible to forecast a climatology in a new position B inside the map. It means that is possible to “transfer” experimental data from the original position (A) to another one (B).

The procedure that allows to build a “transferred climatology”, making use of the CFD simulations, is explained in the following lines.

- 1) CFD simulations are carried out using boundary conditions somehow different from the actual ones, in order to obtain a better convergence of the simulations as will be explained in chapter 5.
- 2) As a result of the CFD calculation the speed-up map is found, in a zone where the experimental climatology can be placed. The speed-up is defined as the ratio between the wind velocity in one point of the map and the wind velocity in a particular reference point.
- 3) This map can be modified (with a linear scaling operation) in the way to have the speed-up value equal to 1 in the location where the climatology is known. A new speed-up map is generated.
- 4) Finally, the scalar value of the experimental velocity in the climatology position must be multiplied per every value of the CFD speed-up map, in order to generate a new velocity map, that simulates the “real” velocity in the numerical domain.
- 5) This procedure is repeated for each sector of the wind rose (12 in the specific case) and for each velocity bin defined in the frequency table of the climatology.

The use of the “transferred climatologies” is needed when doing a “cross checking” of experimental data, for example when inside the map more than one climatology is present. In this way, it is possible to use experimental data of one station and forecast CFD wind data in an other position. Then, it is possible a comparison between the data forecasted by simulations and the experimental data, in a second station.

1.5.4 RESULTS

The Results module allows the user to extract the informations related to the thermo-fluid dynamic variables calculated by the Windfield module.

All extracted variables are not weighted with climatology, but refer directly to the Phoenix database, containing the results of the pure CFD calculation.

1.5.5 WIND RESOURCES

Trough this module is possible to calculate the horizontal wind speed maps all over the site, by weighting the CFD results with the experimental climatology. The procedure to arrive to the wind resources maps is basically the same of the transferred climatology calculation.

By this module, the calculation of the wind turbine losses is possible, using 3 different models.

1.5.6 ENERGY

This module is the last one and calculates the Annual Energy Production (AEP) of the windfarm, with the possibility to include the losses calculation.

1.6 κ - ϵ turbulence model

To close the system of equations introduced by the Navier-Stokes equations a turbulence model is needed. In all the simulations the standard k-e model is used, which needs two new equations to add to the system of numerical equations and introduces two new unknown variables. One is the Turbulent Kinetic Energy (TKE), while the second is the Dissipation Rate (ϵ).

Equation for TKE

$$1.8) \quad \frac{\partial(\rho k)}{\partial t} + \frac{\partial(\rho k u_i)}{\partial x_i} = \frac{\partial}{\partial x_i} \left[\left(\mu + \frac{\mu_t}{\sigma_k} \right) \frac{\partial k}{\partial x_i} \right] + G_k - \rho \varepsilon$$

Equation for ε

$$1.9) \quad \frac{\partial(\rho \varepsilon)}{\partial t} + \frac{\partial(\rho \varepsilon u_i)}{\partial x_i} = \frac{\partial}{\partial x_i} \left[\left(\mu + \frac{\mu_t}{\sigma_\varepsilon} \right) \frac{\partial \varepsilon}{\partial x_i} \right] + C_{1\varepsilon} \cdot \frac{\varepsilon}{k} \cdot G_k - C_{1\varepsilon} \cdot \rho \cdot \frac{\varepsilon^2}{k}$$

The solution of the numerical system need to start with boundary conditions, both for k and ε .

Concerning to k , the values at the inlet section of the model can be described as a vertical profile of TKE, by the equation

$$1.10) \quad K(z) = \frac{u(z)^2}{\sqrt{C_\mu}} \cdot \left(1 - \frac{u(z)}{\delta} \right)^2$$

where,

C_μ is a constant

δ is the height of the boundary layer [m]

An example of the shape of a TKE profile connected to the velocity vertical profile described in figure 1.1 is shown in figure 1.2.

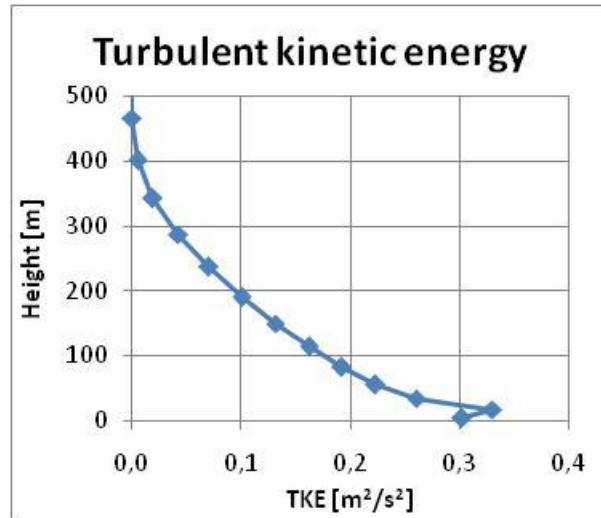


Figure 1.2: Example of TKE vertical profile at inlet section.

The boundary condition for ε used in WindSim is described by the equation

$$1.11) \quad \varepsilon(z) = \frac{u(z)^3}{VKC} \cdot \left(\frac{1}{z} + \frac{1}{L_{obu}} \right)$$

where

L_{obu} is the Obukhov length

A typical shape of the ε vertical profile for an ABL is displayed in figure 1.3.

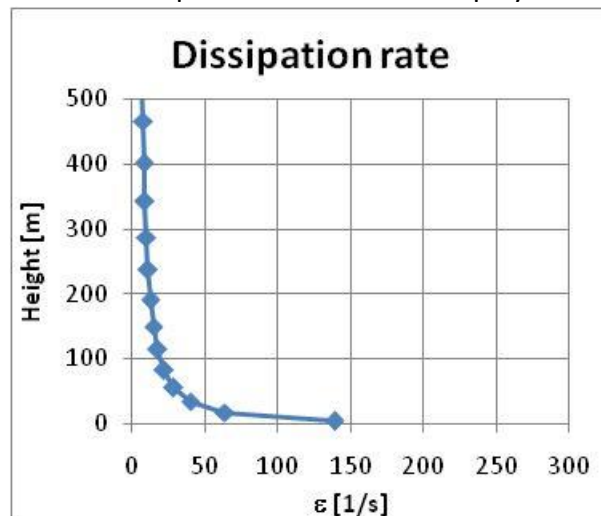


Figure 1.3: Example of ε vertical profile at inlet section

Finally, a turbulence intensity vertical profile can be constructed from the TKE vertical profile

$$1.12) \quad T.I. (z) = \frac{\sqrt{\frac{4}{3}K(z)}}{\sqrt{u^2+v^2}} * 100$$

The shape of the TI vertical profile for the usual example is reported in the figure 1.4

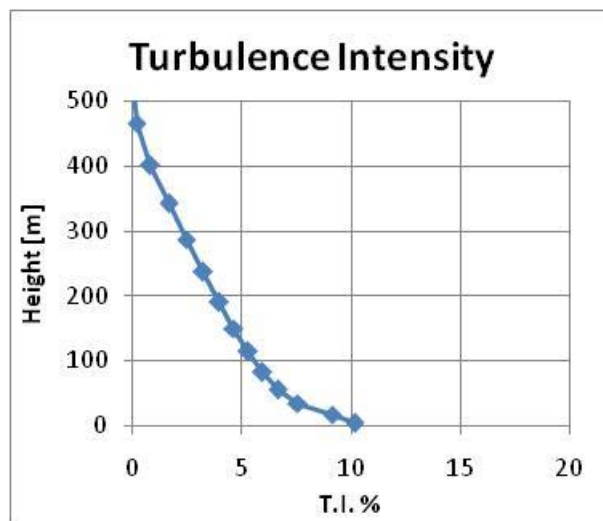


Figure 1.4: Example of TI % vertical profile at inlet section.

CHAPTER 2. Simulations over flat terrains.

2.1 Introduction

The purpose of this chapter is to analyze the influence of some parameters which can be set by the user in predicting wind velocity fields, by running the Terrain and the Windfield modules of WindSim. A series of 2D simulations has been done, by varying step by step one of the parameters considered, i.e. roughness value of the terrain, number of cells in vertical direction, height distribution factor and point of measurement [14]. Moreover, particular attention has been devoted in finding the correct boundary condition at the top of the model, able to describe the physical phenomena of the geostrophic wind and neutral ABL.

First of all, it is useful to recall the principle by which an incompressible turbulent flow develops over a flat plate, under a constant mass flow rate. The reader should remind the meaning of the figure 2.1 below.

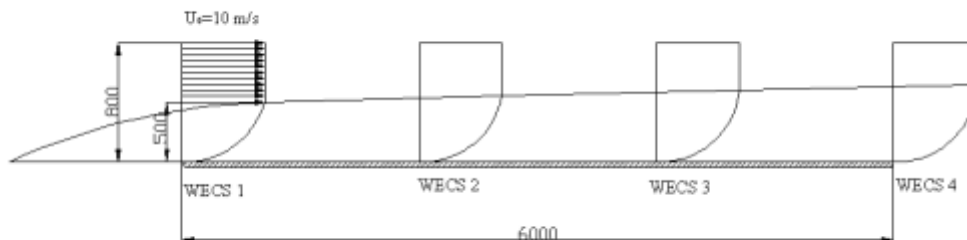


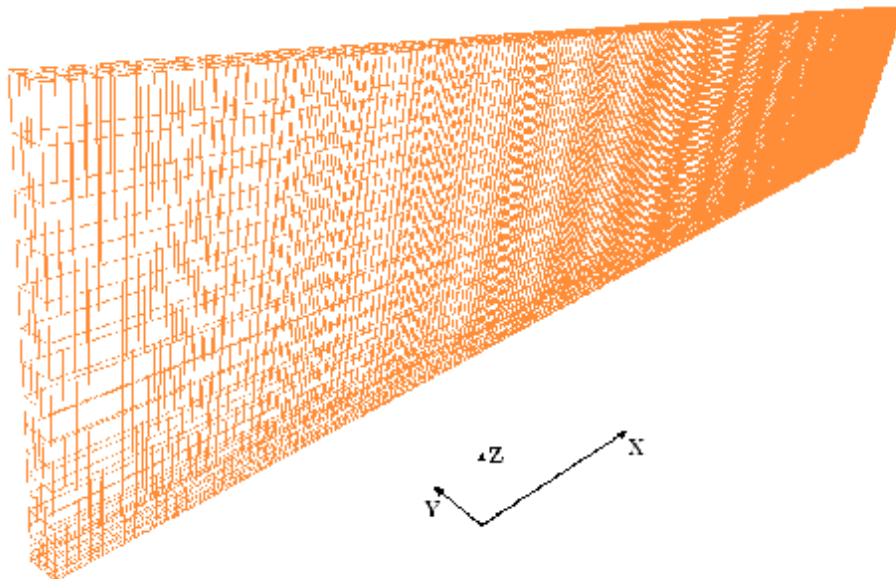
Figure 2.1: Development of the boundary layer thickness of an air flow upon a flat plate. All dimensions in meters.

Referring to the Atmospheric Boundary Layer (ABL), the situation can be supposed to be like a flow passing over a rough plate, with a known value of roughness. A neutral ABL can be considered as the case of the flow over a flat plate very far from the beginning of the plate. In this case the upper boundary of the model can be

considered as a virtual wall with no friction, describing a wind tunnel containing all the air flowing away, while the flow reaches the maximum velocity value (free stream velocity) at the top of the boundary, as shown in the previous figure 2.1.

2.2 Geometrical model and inlet BC

The model analyzed forward is a simple flat terrain 6000 meters long and 100 meters large, a height above the terrain set at 800 meters. The height of the ABL was fixed at 500 m and a geostrophic wind of 10 m/s was set above 500 m up to 800 m.



Different cases of terrain roughness have been studied and for each case the correct inlet velocity profile was chosen. The model used for the inlet profile was the log law, explained in chapter 1, while the turbulence profile was set by means of the particular equations also reported in the introductory chapter.

For the purpose of data reduction the vertical profiles of the wind speed have been extracted in 4 different stations along the flat terrain, respectively station 1 at the

inlet position (0 meters), station 2 at 2000 meters, station 3 at 4000 meters and station 4 at the outlet position (6000 meters).

2.3 Vertical discretization sensitivity

To determine the influence on the results of the vertical discretization a series of simulations has been run. These simulations have been done keeping constant the roughness value and varying the vertical number of cells, without any change of the height distribution factor (HDF). According to this, the first group of simulations have been done for roughness $R=0.01$ with 20, 40 and 60 vertical number of cells. The height distribution factor has been set at the value 0.1. Concerning the second group of simulations, the roughness value has been set to $R=0.6$ and the number of vertical cells was the same of the previous group.

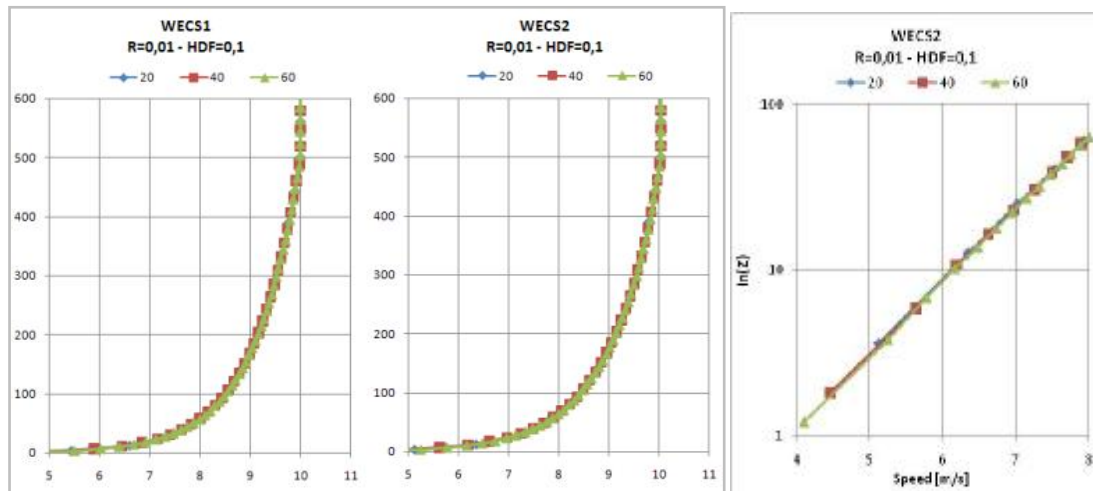


Figure 2.2: Vertical position in different stations for $R=0.01$ and $HDF=0.1$.

Figure 2.2 shows the overlapping of the vertical profiles relative to station 1 (inlet of the geometrical model) and those of station 2. As it can be seen there is no shift between the three curves and then the simulation with the lowest number of vertical cells is enough accurate to describe the vertical profile. The same result can be deduced from the right side of the figure where the wind profiles at the different stations were plotted in log scale.

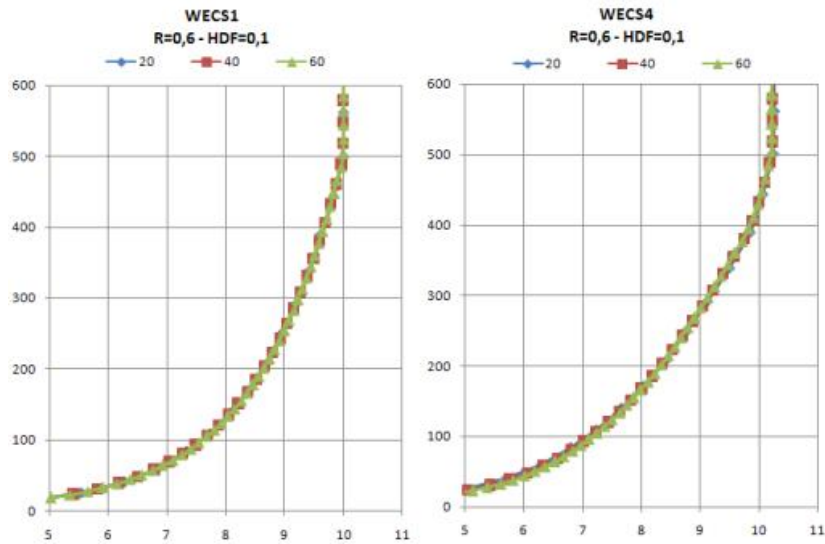


Figure 2.3: Vertical profiles in different stations for R=0.6 and HDF=0.1.

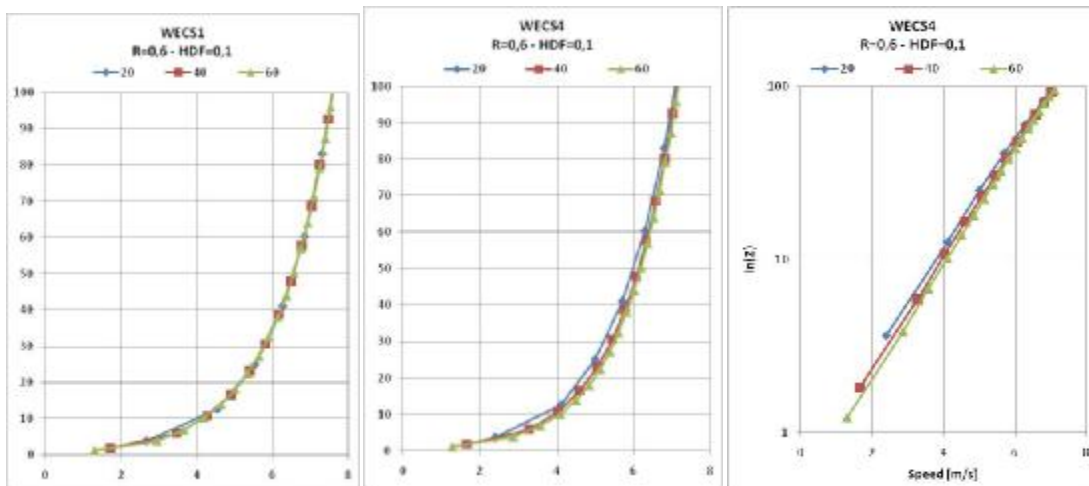


Figure 2.4: Zoom of the figure 2.3.

Figures 2.3 show the results obtained for a high roughness value ($R=0.6$). Also for a such high value of roughness the influence of the vertical number of cell seems to be not important. Particular attention must be done when analyzing the vertical profiles in the first 100 meters height, where 40 seems to be a more accurate value for the number of vertical cells (zoom of figure 2.4 in log scale).

The conclusions valid for low roughness simulations can then be drawn also for a high value of roughness height, and the minimum number of cells can be still maintained to run the simulations.

2.4 Height distribution factor sensitivity

In the following paragraph the sensitivity of the results to the height distribution factor has been studied for a high roughness value ($R=0.6$). After having established the influence of the vertical cell number, the boundary layer thickness discretization has been left at the default value (20 cells). Then, two series of simulations have been run: the first one studies the influence of the increase of the height distribution factor, from the default value to 0.4; the second one studies the reduction from the default value to 0.01. The vertical profiles have been extracted at the inlet and at the outlet of the geometrical model and the results are shown in the figures 2.5 and 2.6.

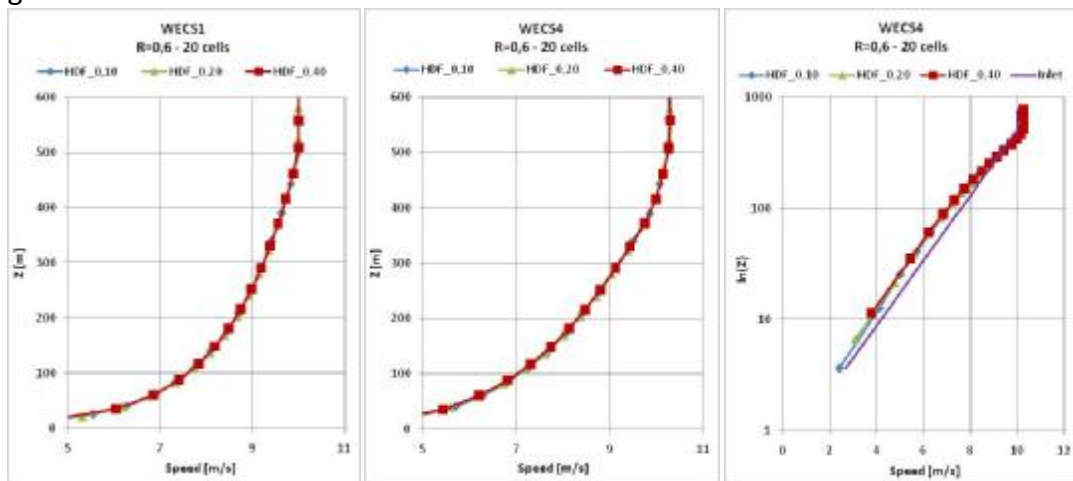


Figure 2.5: Vertical profiles in stations 1 and 4 for $R=0.6$, increasing the height distribution factor.

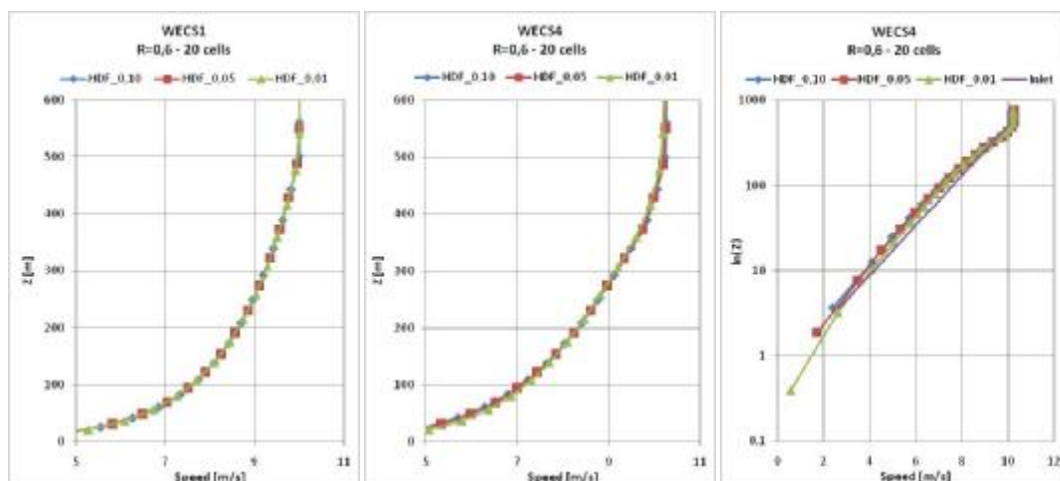


Figure 2.6: Vertical profiles in stations 1 and 4 for $R=0.6$ and decreasing the height distribution factor.

As it can be seen from figures number 2.5 and 2.6, the increase of the height distribution factor seems to be not responsible of the variation of the vertical profiles. Anyhow, a better accuracy of the results is obtained, as a general rule, for low values of the HDF, which moves the vertical central cell points in a region closer to the ground, where velocity gradients are higher.

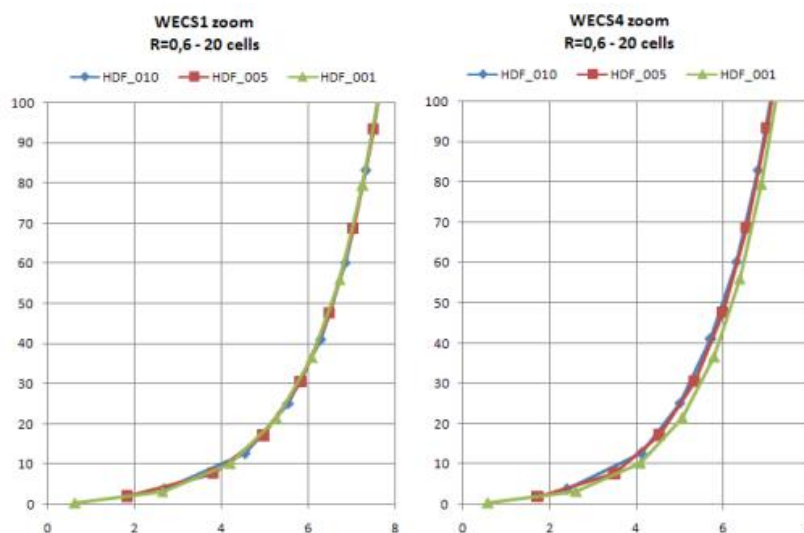


Figure 2.7: Zoom of the figure 2.5.

Figure 2.6 shows that the three curves don't overlap completely, as it can be seen also from the zoom in the further figure 2.7.

The phenomenon shown in figure 2.7 is very important, and must be taken into account when working with wind data measured at low height respect to the ground level. So it's reasonable to run simulations with low values of height distribution factor when a particular precision near the ground level is needed, while in situations regarding the evaluation of the wind resources at high levels this phenomenon becomes less marked. The experience shows, moreover, that the height distribution factor cannot be decreased indefinitely, but its value is strictly connected to the resolution of the horizontal grid. Then, if it is necessary to have good precision in such low heights, also the dimension of the horizontal cells must

be decreased, to avoid to build thin and long cells near the ground, and also avoid connected divergence problems of the solver.

Finally, another consideration must be done when varying the vertical discretization grading: according to Crasto G. [1], the height of the first cell near the ground should be at least double the height of the roughness height.

2.5 Influence of the top boundary condition

To determine the most realistic boundary condition at the top of the domain, able to describe the atmospheric boundary layer development over a flat terrain, two series of simulations have been run. In the first series the *zero relative pressure* was assigned, while in the second series the *wall with no friction* condition was assigned. The zero relative pressure boundary condition is an outflow boundary condition where the pressure is set to the atmospheric value and the velocity is evaluated from the internal side of the calculation domain.

The wall with no friction BC fixes a null value for the vertical component of the velocity while the other velocity component and the pressure are calculated. This BC behaves like a physical wall where the viscous tangential forces are equal to zero.

Each series of simulations have been run for 3 different roughness values, i.e. $R=0.01$ (corresponding to a flat area with very short grass), $R=0.3$ (a countryside with few trees and farms) and finally $R=0.6$ (typical value for a forest). For all the simulations the number of cells in the vertical direction was set to 20.

The following figures show the results of the simulations, particularly the wind profile of the atmospheric boundary layer.

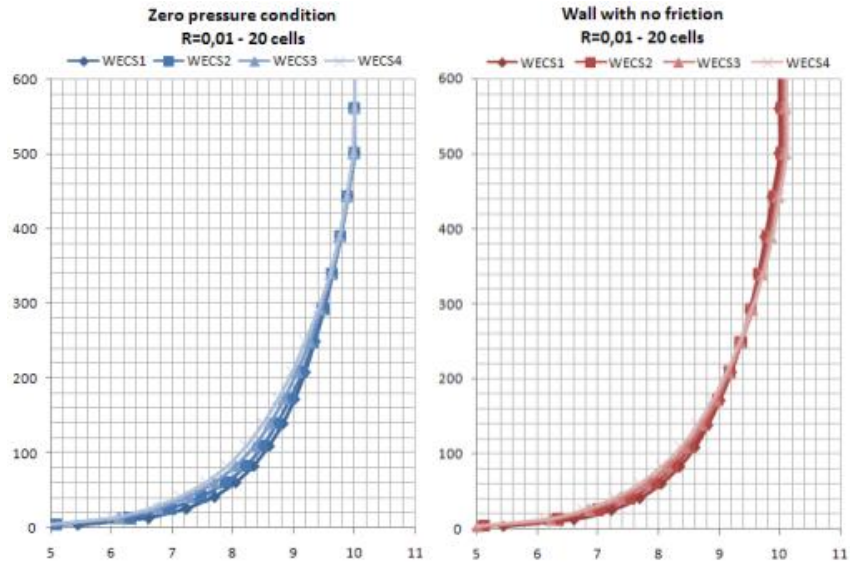


Figure 2.8: Vertical profiles for $R=0.01$ and 20 vertical cells.

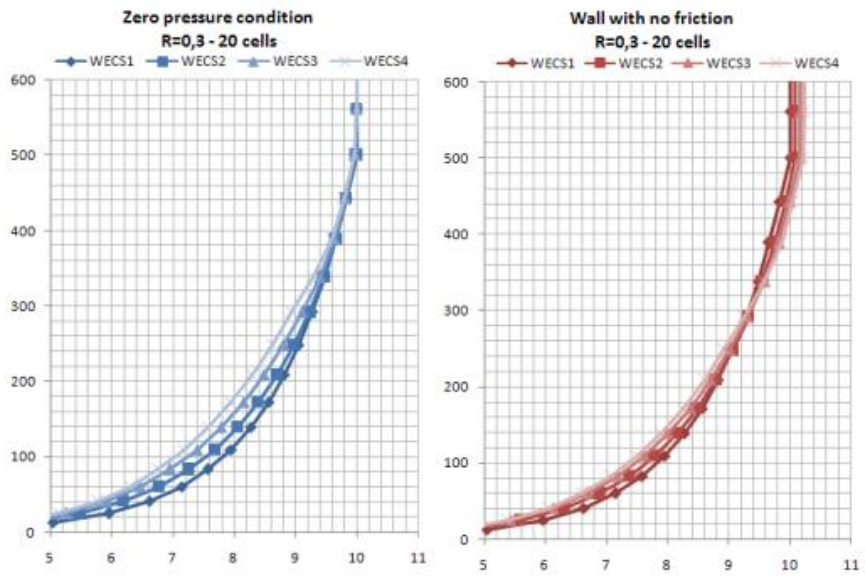


Figure 2.9: Vertical profiles for $R=0.3$ and 20 vertical cells.

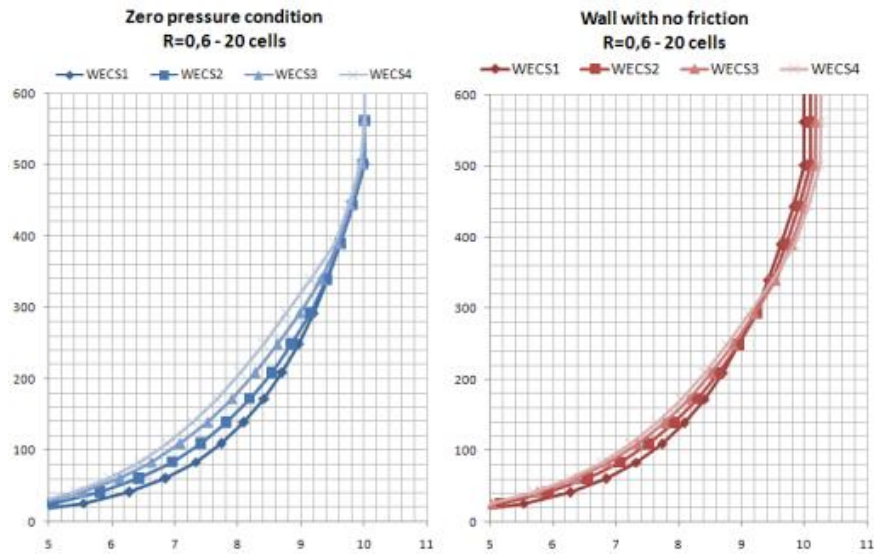


Figure 2.10: Vertical profiles for $R=0.6$ and 20 vertical cells.

As can be easily seen, the two boundary conditions at the top level produce different results in the shape of the velocity profile. The most important effect is the velocity variation during the flow development on the domain. According to the neutral ABL, the velocity profile should not change along the domain, and this is the case of figure 2.8 where a low value of roughness was imposed. The effect of profile variation is amplified by the growth of the roughness value. This effect can be due to numerical models used to close the RANS equation (turbulence model) and to simulate the terrain roughness. In all the cases showed before, the effect of the velocity deficit is more evident in simulations with zero pressure condition than in simulations with the wall with no friction condition.

The second effect is the deformation of the velocity profile near the geostrophic height. The zero pressure boundary condition produces in the exit station a velocity profile that shows a maximum velocity similar to that of the inlet station. This fact, together with the strong decrease of the axial velocity below the geostrophic height leads to a diminution of the volume flow in the vertical stations along the domain. As a results, the vertical component of the velocity increases and parte of the volume flow exits from the top boundary condition, in contrast with the definition of neutral ABL.

A different behavior is obtained when the wall with no friction BC is used. According to the development of the boundary layer, the effects of the roughness slow down along the domain and the velocity deficit between two stations decreases when going toward the end of the flat terrain, confirming the general behavior of a turbulent neutral ABL. Due to the particular boundary conditions the same volume flow is maintained for all the vertical stations and the streamlines remain horizontal.

2.6 Remarks

Flat terrains, like the model considered in this document, are not sensitive to the increase of the vertical number of cells and then the default value (20 cells) seems to be a good compromise between the goodness of the final results and the time taken to run the simulation. An increase of the vertical cell number could be useful in terrain with complex features.

The height distribution factor needs to be taken into account when, in region of the flow attached to the terrain, a good precision is needed, and it should be decreased slowly together with the increase of the horizontal resolution to avoid divergence problems. From this analysis a limit value of the Height distribution factor was found (HDF= 0.01).

The boundary condition to be assigned to the top level of the geometrical model has been analyzed, showing that mass flow rate can be conserved only assigning the wall with no friction boundary condition. This BC produces the effect of the shift of the free stream velocity respect to the value superimposed in the WindSim module, but the effect of velocity deficit due to the roughness value is strongly reduced, mostly for high values of the roughness height.

CHAPTER 3. Simulations over 2D hilly terrains.

3.1 Introduction

The present chapter concerns a series of numerical simulations on the behavior of an air flow over a single hill, characterized by a roughness value on the ground. All simulations have been run using a geometrical model in which a 3D “ \cos^2 ” hill is built: this fact causes the presence of a finite Y-dimension, generating 3D effects. To avoid these errors can affect the results, all the simulations have been run considering a very thin section of the 3D hill domain in Y-direction and the velocity vertical profiles, needed for further analysis, have been extracted in the central cell of the y-direction.

The results obtained in this way are totally equivalent to the results that could be extracted by a “flatted” hill in y-direction, as illustrated in the next paragraph (figure 3.2), and they are not affected by tridimensional flow effects.

The shape of the speed vertical profiles along the hill has been also studied, and the speed-up effect due to the variation of the terrain’s height has been highlighted. Moreover, a sensitivity study of the numerical grid for hilly terrain has been carried out, both for horizontal and vertical discretization.

Finally, a test case concerning the behavior of an air flow over a hilly terrain as been reproduced: this case was good enough to put in evidence the separation phenomena of the flow downstream the hill, and set some important parameters of the WindSim software, to be used in complex terrains.

3.1.1 Similitude between a 3D “ \cos^2 ” and “flatted” hill

The following figure 3.1 shows the shape of the geometrical model in the X-Z plane used to study the speed-up effect. The measurement stations, where speed vertical profiles have been extracted, are also visualized.

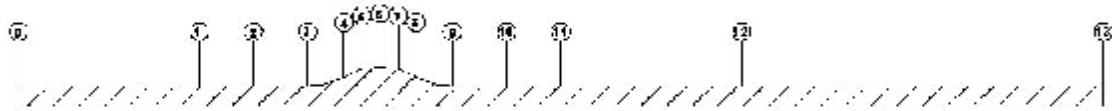


Figure 3.1: Measurement stations of the speed vertical profiles.

The length of the model is 6000 m in X-direction, while the width is 75 m in Y-direction. The number of cells is 240 X 3, for an equally spaced horizontal discretization of 25 m. The height of the model is 1000 m and the vertical Z-direction contains 20 cells; the Height Distribution Factor is 0,1 that means the first cell height is equal to 9 m.

The hill shape has height $H=100$ m and half-length $L=400$ m, while the top is centered at a distance of 2000 m from inlet section.

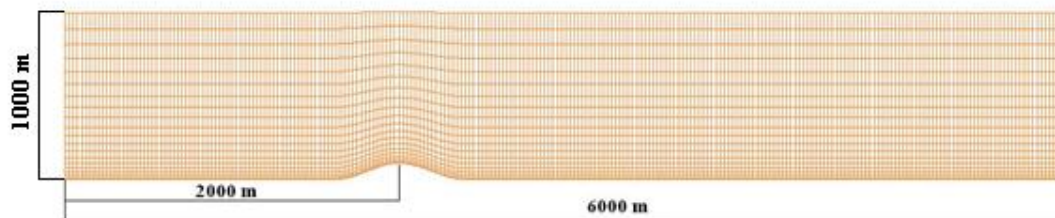


Figure 2: Sketch of the numerical domain of the 2D hill.

In the following table 3.1 the coordinates of the measurement points are reported.

Table 3.1: Measurement coordinates

Station	X [m]	Y [m]
0	0	37,5
1	1000	37,5
2	1300	37,5
3	1600	37,5
4	1800	37,5
5	1900	37,5
6	2000	37,5
7	2100	37,5
8	2200	37,5
9	2400	37,5
10	3700	37,5
11	3000	37,5
12	4000	37,5
13	6000	37,5

The height of the model is 1000 m from the ground level and the top surface contain the wall with no-friction boundary condition, studied in the previous chapter. The geostrophic wind is standing at 500 m with a free stream velocity equal to 10 m/s. The description of the model is completed with the following figure 3.2, showing the difference in the profile of the hill obtained with a section of the Y-Z plane, between a “ \cos^2 ” and a “flatted” hill.

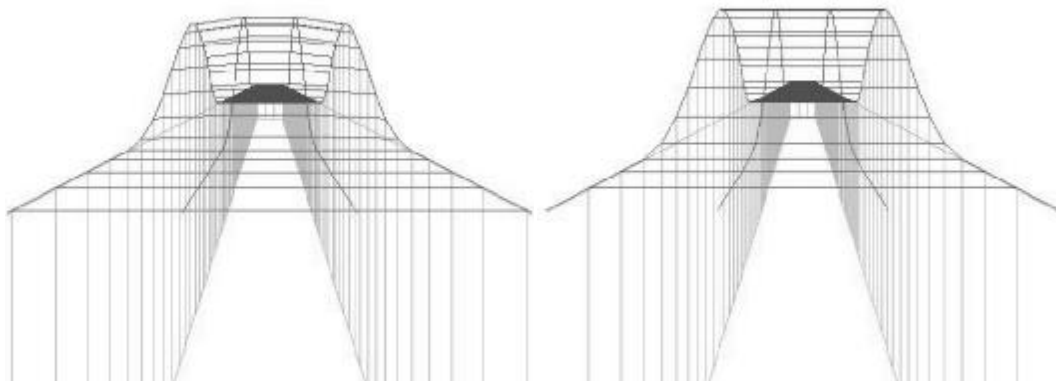


Figure 3.2: View of the geometrical models “ \cos^2 ” and “flatted” hill from the inlet sections.

The “ \cos^2 ” hill has been built through the automatic procedure of the WindSim software, which draws the hill shape using the following equation.

3.1.1) _____ ,

while the flattened one has been created manually by the same 3D equation projected in the X-Z plane.

The simulations have been run using the standard κ - ϵ turbulence model and the coupled solver. The velocity vertical profiles have been calculated in some of the stations shown in figure 3.1 and the results of the comparison are reported in the following figures 3.3.

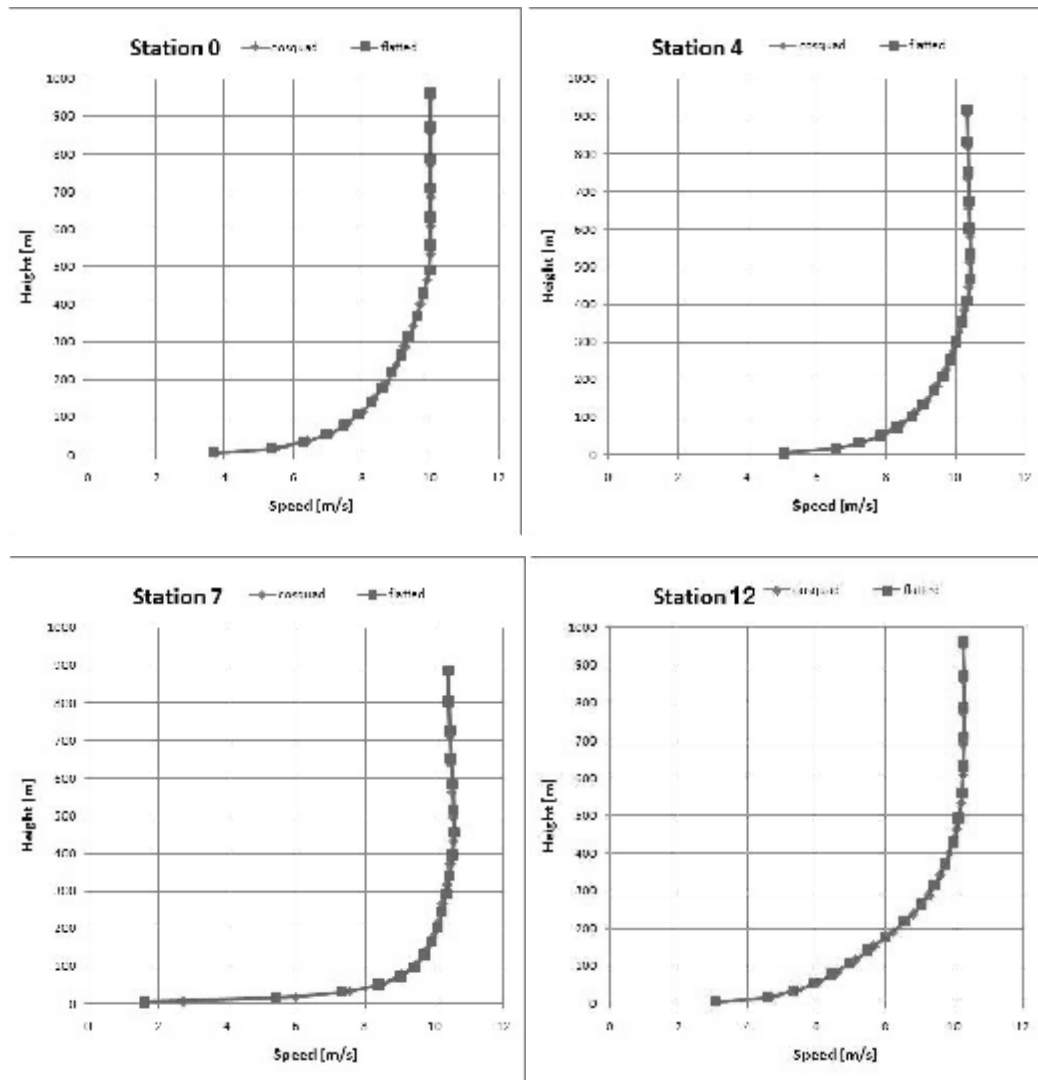


Figure 3.3: Overlay of the vertical profiles in some measurement stations.

The results shown above demonstrate that the vertical profiles extracted from the central cell in the Y-direction of a 3D “cos²” hill are the same as those extracted from a “flatted” Y-direction hill, which doesn’t introduce therefore any tri-dimensional effect on the flow. This is a striking result, and it is useful for further simulations, because it’s demonstrated that a very thin 3D “cos²” hill, built up with the automatic procedure of the WindSim software, can be used to understand the behavior of an air flow over a 2D hill.

3.2 Speed-up effect

The speed-up effect has been studied for the case shown in figure 3.1, varying the value of the roughness height. Particularly, 3 tests have been done, for R=0,01 m , R=0,3 m and R=0,6 m.

The following picture shows the development of the speed vertical profiles measured in the various stations, upstream and downstream the hill top (located in station number 6).

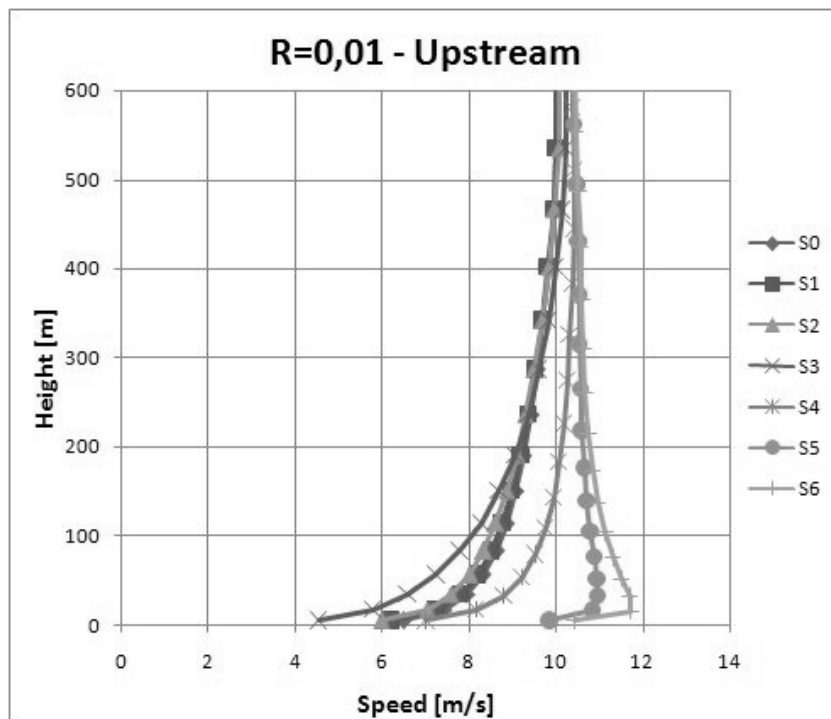


Figure 3.4: Development of the vertical profiles upstream the hill-top.

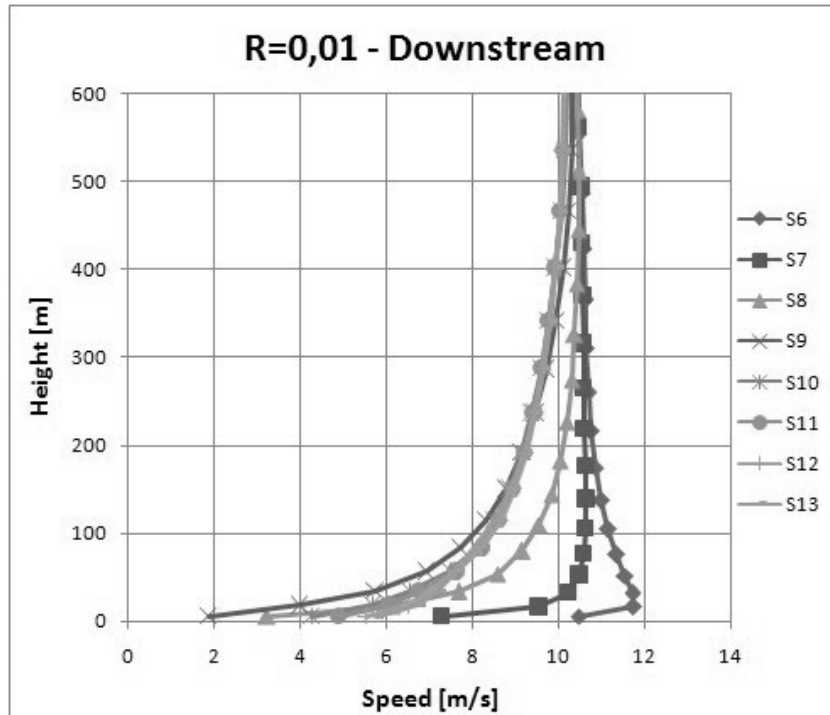


Figure 3.5: Development of the vertical profiles downstream the hill-top.

Starting from station 0, the flow is slowing down until station 2: this behavior is caused by the effect of the roughness (as explained in chapter 2). Between station 2 and station 3 the slowing effect is stronger, due to the presence of the obstacle (the hill). When the flow passes over station 3, the hill geometry accelerates the flow, like in a subsonic convergent duct. The speed increase is registered until the top of the hill, where the horizontal speed gets the maximum value.

When the flow passes over station 6 (hill top), it sees a subsonic divergent duct and the effect of slow down begins and continues until station 9. The divergent channel tends to increase the boundary layer more than the starting value. Then gradually, the flow starts to accelerate again, in order to get the initial conditions.

A similar behavior of the flow over a hilly terrain can be found in [2]

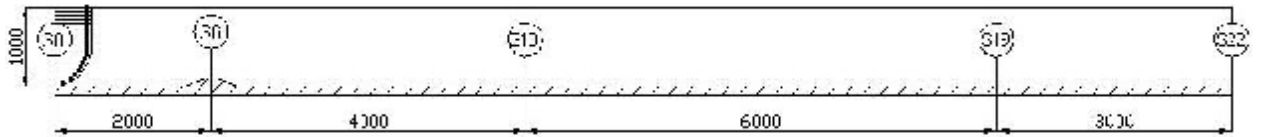


Figure 3.6: Sketch of the 15 km long hilly terrain model.

A longer model described in figure 3.6 has been studied to highlight the behavior of the vertical profiles in the far region downstream the hill top. As it can be seen from figures 3.7 the vertical profile of the flow seems to be fully developed in station 22, in a distance 13 km away downstream the hill top.

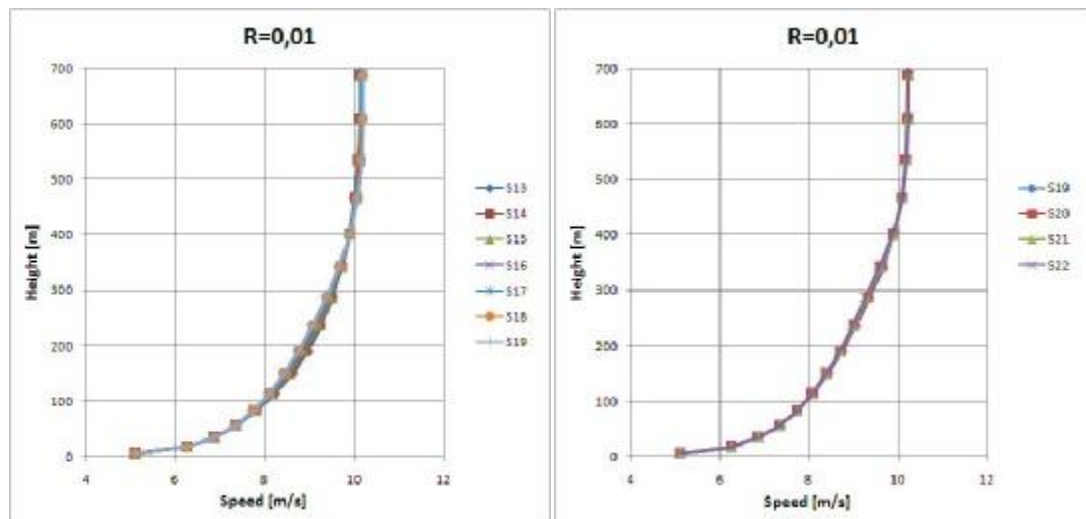


Figure 3.7: Vertical profiles extracted in the downstream flat area of the model.

The simulations run for $R=0,3$ m and for $R=0,6$ m show similar behavior of the speed up effect. The effect of the roughness height is stronger than that of the first simulation, and this fact induces a different growth of the boundary layer thickness, as shown in the figure 3.8.

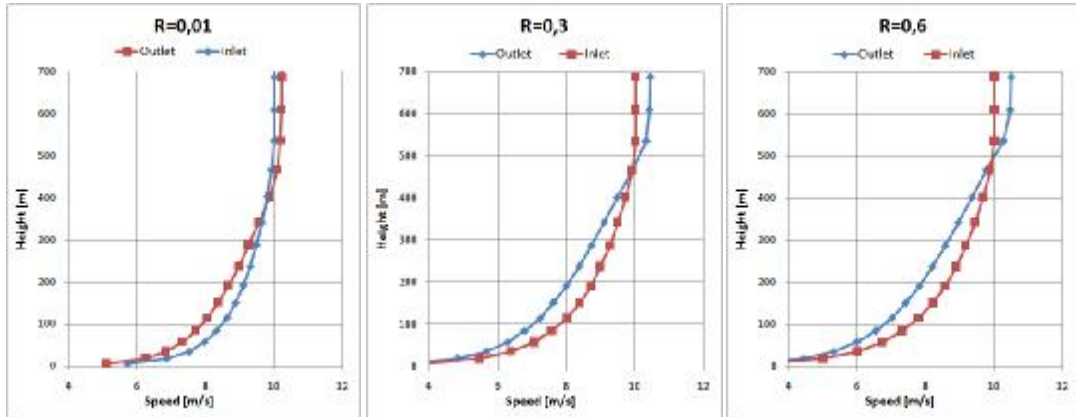


Figure 3.8: Different growth of the boundary layer thickness at the outlet position.

3.3 Sensitivity to the first cell height

The model described in figure 3.1 with a roughness value equal to 0,01 m has been studied in order to understand the influence of the vertical discretization. The number of cells has been fixed to 20, but the value of the height distribution factor has been changed from the default value 0,1 to lower values, that means the height of the first cell attached to the ground is reaching lower values.

The following pictures show the velocity profiles while the height of the first cell is changing.

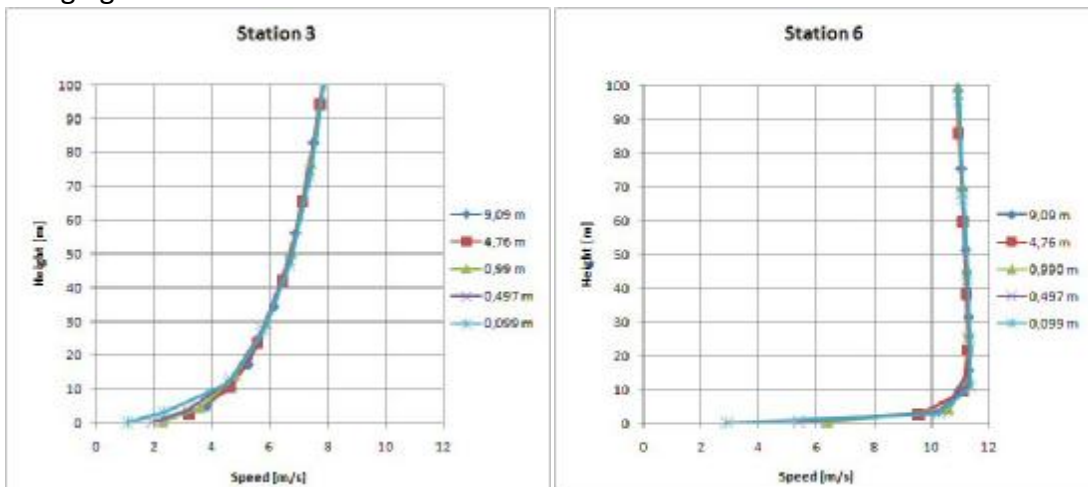


Figure 3.8: Influence of the height distribution factor variation in stations 3 and 6.

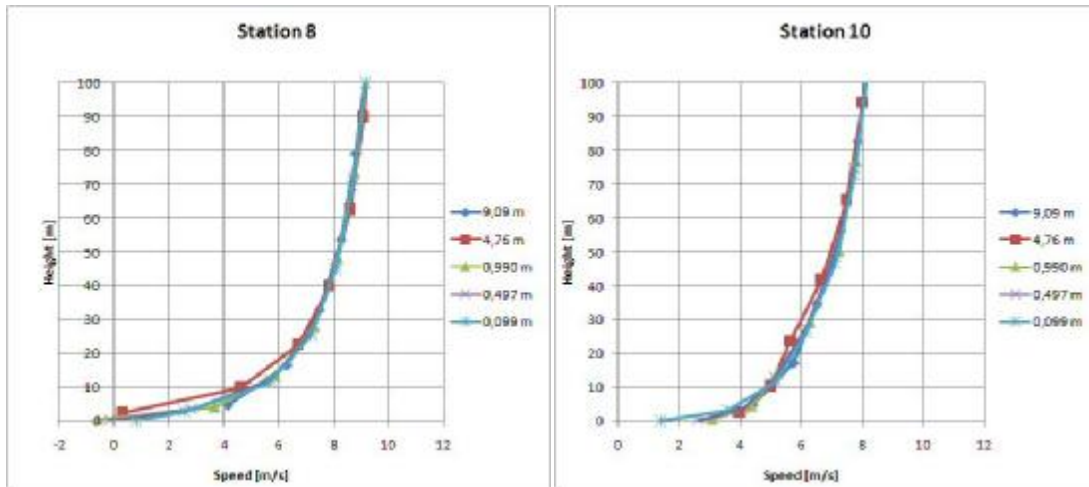


Figure 3.9: Influence of the height distribution factor variation in stations 8 and 10.

The pictures 3.8 and 3.9 show how the velocity profiles reach a different shape when the height distribution factor is varying. Particularly, the value $HDF=0,01$, which produces the first cell height equal to 0,99 m, seems to be the one when the vertical grid independence is obtained.

In the picture 3.10 with semi logarithmic axis for vertical height, valid in station 8 where the flow can incur in a recirculation bubble, is shown in good manner the trend of the vertical profiles, which in the case mentioned has the best alignment.

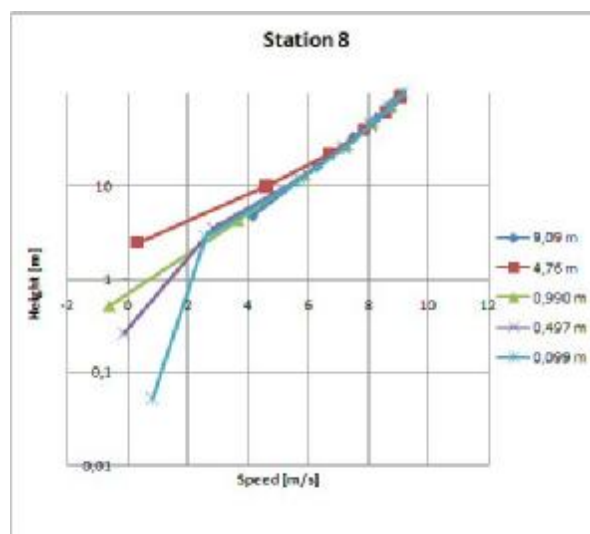


Figure 3.10: Influence of the height distribution factor variation in station 8.

3.4 Sensitivity to the horizontal discretization

Another group of simulations has been run, in order to determine the influence of the horizontal discretization. The model used for the test is the same of the previous paragraph, but the number of cells in x-direction has been varied to change the grid spacing in a range from 100 meters to 10 meters. For 20 m and 10 m resolution the refinement option has been used.

As it can be easily seen from pictures 3.11 , 3.12 and 3.13 the grid independence is obtained for a horizontal resolution of 20 m

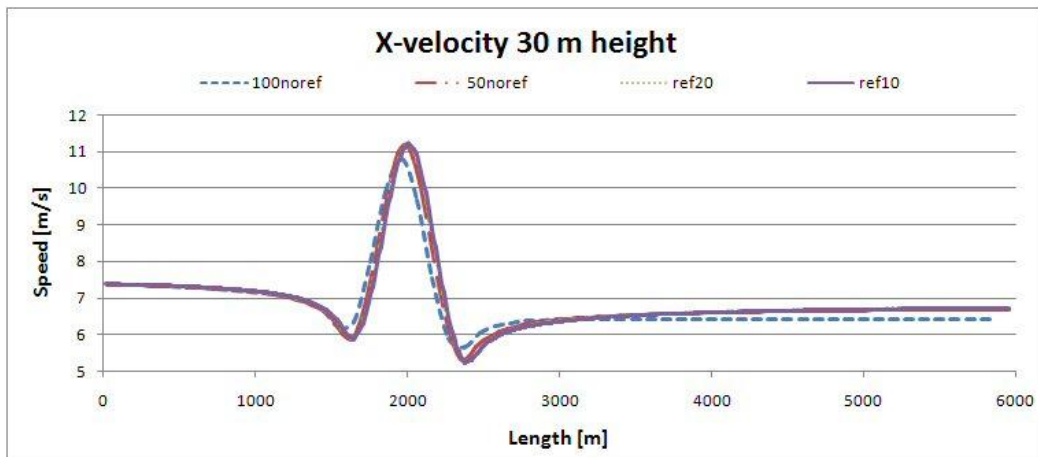


Figure 3.11: X-velocity at 30 m height.

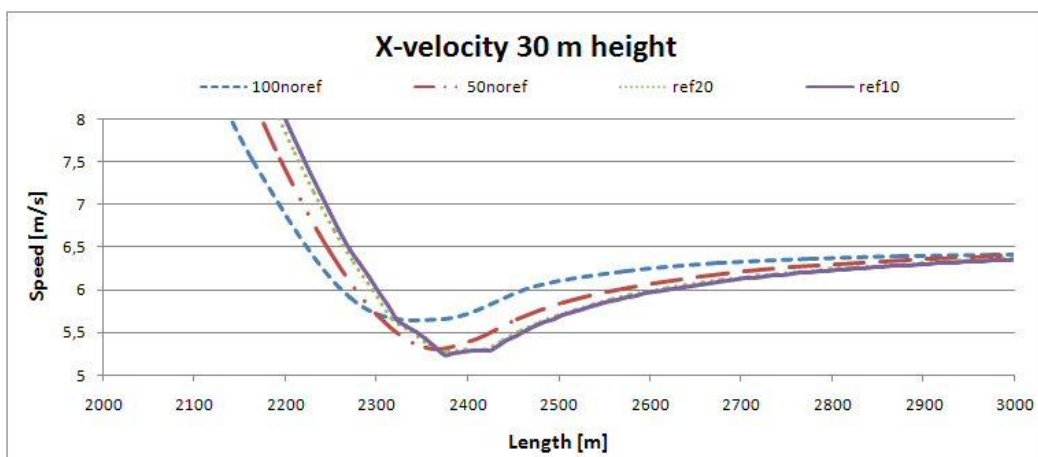


Figure 3.12: Zoom in a region downstream the hill top.

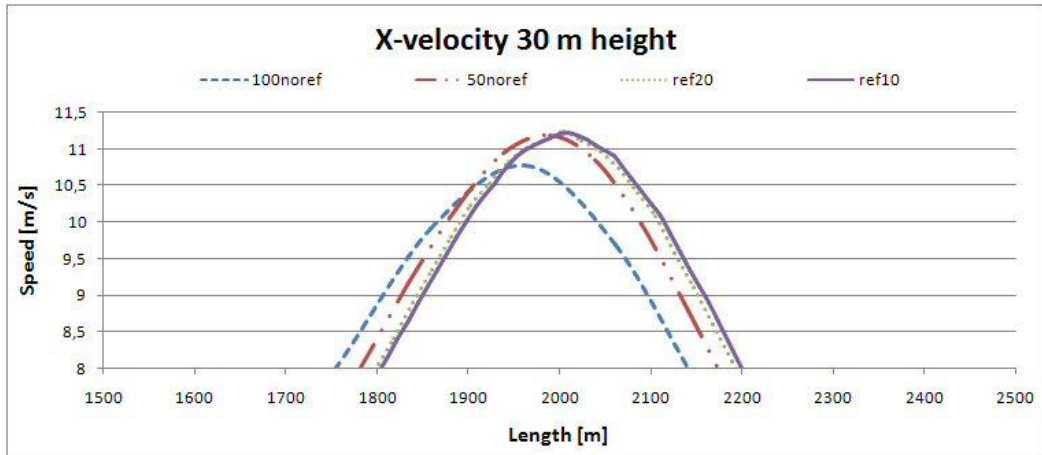


Figure 3.13: Zoom of the hill top region.

3.5 Investigation of a flow over a 2D roughly hilly terrain

This paragraph presents an investigation of a flow over a two-dimensional roughly hilly terrain. Two cases have been realized, to highlight the behavior of an attached flow and a separated flow, by varying the slope of the hill. A comparison between experimental [3] and numerical data has been done, showing good agreement.

The numerical simulations are based on the solution of Reynolds Averaged Navier-Stokes (RANS) equations, by the software WindSim, with the $\kappa - \varepsilon$ turbulence model. The realized numerical grid is in both cases non orthogonal body fitted type, while the solver which better produces a good level of convergence in the solution is the segregated, enriched with the “yap” correction of the $\kappa - \varepsilon$ turbulence model.

3.5.1 Description of the test case

The experimental data comes from the tests conducted in an open-circuit boundary-layer wind tunnel, having a test section of width, height and length of 1.2, 1.2 and 6 meters respectively. A neutrally stratified boundary layer is developed using equally spaced 0.25 m height triangular spire-type vortex generators and artificial grass of 5 mm height. In these experiments, a fully-developed turbulent boundary layer of 0.25 m height is generated, and has a typical wind profile over an open flat area (fitting to a log-law profile).

The shape of the hill is described by the following equation

$$3.5.1) \quad Z = \frac{H}{2} \cdot \left\{ 1 + \cos \left[\left(\frac{\pi}{2} \right) \cdot \left(\frac{x}{L_1} \right) \right] \right\} \quad ,$$

where:

H [m] is the height of the hill

L_1 [m] is the half-length of the hill at the upwind mid-height of the hill

The slope of the hill S is defined as the average slope for the top half of the hill upstream of the crest, through the equation

$$3.5.2) S = \frac{H}{2 \cdot L_1}$$

The two models for attached and separated flow are identified as S3H4 and S5H4, where S3 and S5 stand for the hill slope of 0.3 and 0.5 respectively, and H4 stand for the hill height of 4 centimeters.

A sketch of the studied hill is presented in the following figure 3.5.1

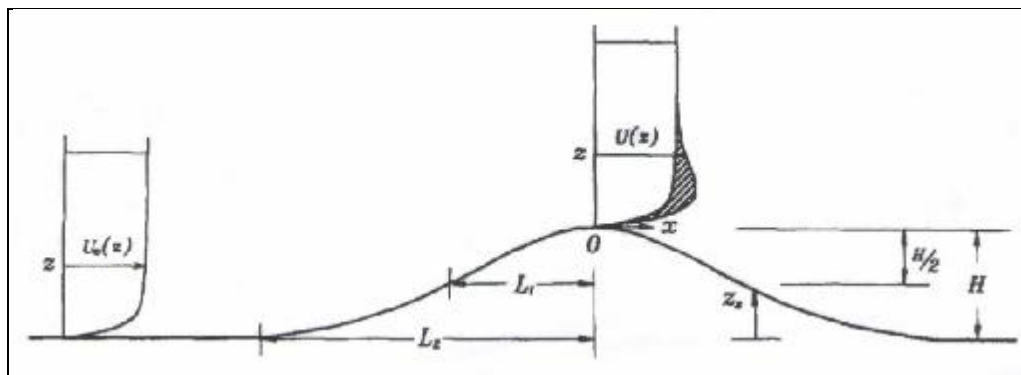


Figure 3.14: Schematic diagram of the wind flow over a hill.

The experimental data, relative to the flow at the inlet of the wind tunnel, have been curve-fitted with the following logarithmic function

$$3.5.3) U_0(z) = \frac{U_*}{k} \cdot \log\left(\frac{z}{z_0}\right)$$

with:

$$U_* = 0,33 \quad \text{m/s}$$

$$z_0 = 0,05 \quad \text{mm}$$

$$K = 0,41 \quad \text{Von Karman constant}$$

The following figure 3.15 shows the interpolation of the velocity profile.

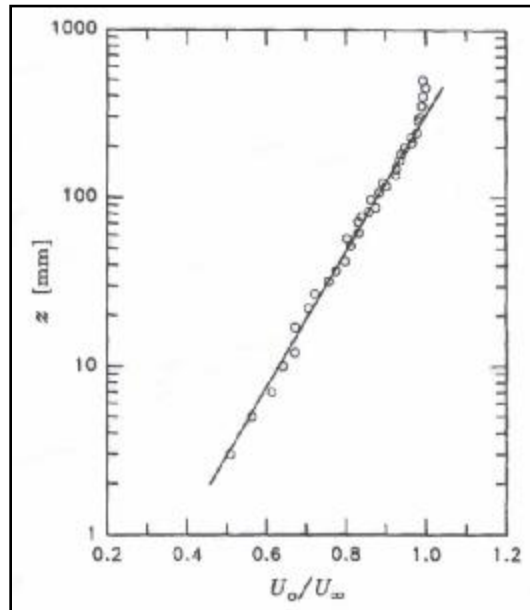


Figure 3.15: Mean velocity profile over the flat floor at inlet position.

3.5.2 Numerical method

The WindSim software solves the RANS equations (Reynolds Averaged Navier Stokes), using the finite volume method through a graphical interface and the CFD solver Phoenics. The continuity equation and the three momentum equations are solved, using the standard $\kappa - \varepsilon$ turbulence model to introduce the equations needed to solve the Reynolds stress tensor.

Also the “Yap” correction for $\kappa - \varepsilon$ turbulence model is needed to help the convergence process in the separated flow. In addition to the turbulence model, the wall function is imposed to the ground.

The fluid blown by the wind tunnel is air, in incompressible flow condition: in this way, no thermal effects are present. The characteristic of this flow is expressed by the Reynolds Number, defined as

$$3.5.4) \quad Re_H = \frac{U_\infty \cdot H}{\nu} = 1,87 \times 10^4$$

where:

U_∞ [m/s] free stream velocity

H [m] hill height
v [m²/s] air cinematic viscosity

The numerical grid is built, in both S3H4 and S5H4 cases, with a non-orthogonal grid as shown in the further paragraphs.

3.5.3 Case S3H4

The case S3H4 has been studied in a geometrical model extending from streamwise coordinates -1,5 m to 1,5 m, for a total length of 3 m. The width is 0,1 m while the total height is 2 m; thus, the blockage factor is 2%.

The grid contains 130 cells in x-directions, 3 cells in y-direction and 60 cells in vertical direction; the horizontal resolution has been increased in the hill profile, in order to describe in a sufficient accurate way the behavior of the flow in the critical zone. The height of the first cell is 2 mm, while the roughness height imposed in the wall function formula is 0,05 mm. Then, the condition stating that the first cell height must be at least double of the roughness value is satisfied [1].

The following figure 3.16 shows a particular of the numerical grid.

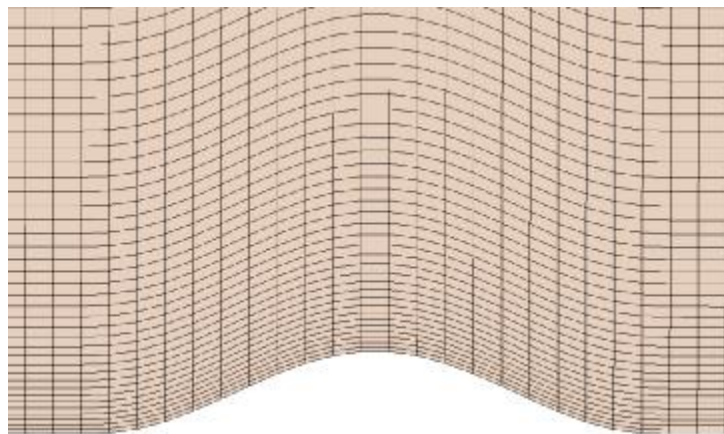


Figure 1.16: Particular of the non-orthogonal regular grid for case S3H4.

The boundary condition imposed on the top of the model is “wall with no-friction”, while at the bottom the terrain contains the roughness information all over the surface.

At the inlet, the vertical profile of velocity is imposed, with the log law reaching a speed of 7 m/s, to keep the same Reynolds Number of the experiments, at 0,28 m height.

To run the simulation, a segregated solver (SIMPLEST) has been used, with a number of iterations equal to 1000.

The picture below shows the residual values of the variables solved by WindSim.

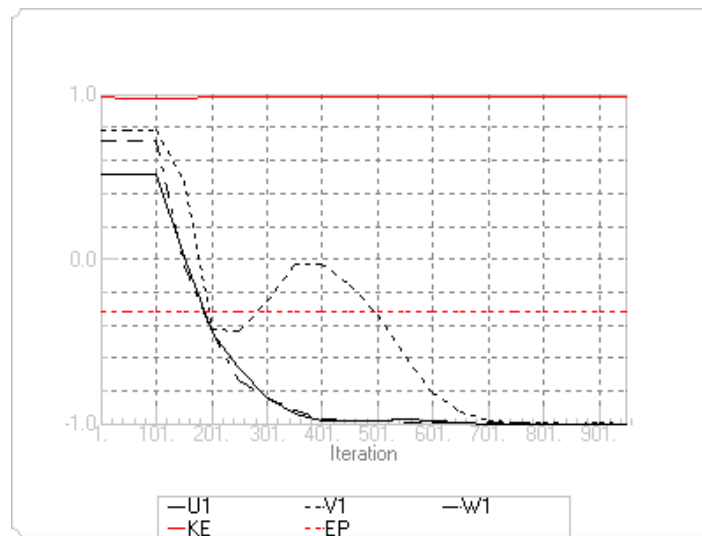


Figure 3.17: Residual values for case S3H4.

3.5.4 Case S5H4

The case S5H4 has been created keeping the same original dimensions of the case S3H4.

Due to the shape of the hill, steeper than case S3H4, the number of horizontal cells has been grown up to 150, to obtain a total number of cell equal to 2700. The height of the first cell is still 2 mm, while the roughness value of the terrain is 0,05 mm. Also in this case, the height of the first cell is higher than the two times the value of the roughness height. Concerning the boundary conditions, the top of the model is set as “wall with no-friction”.

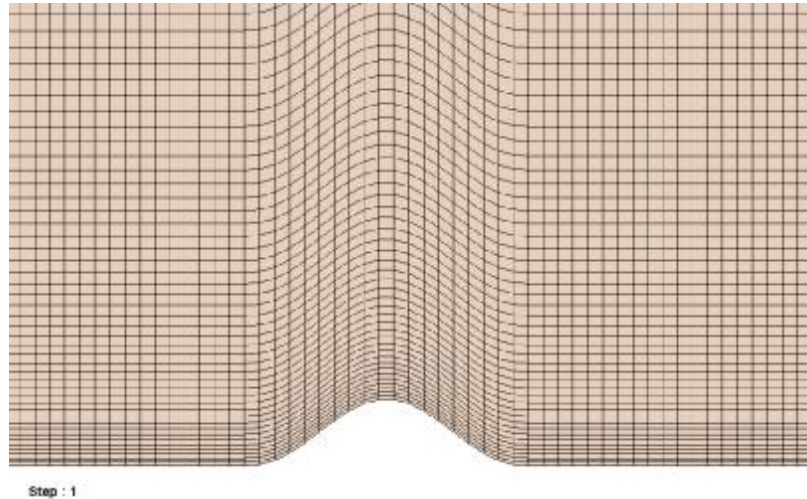


Figure 3.18: Particular of the numerical grid for case S5H4.

To run the simulation, the segregated solver option has been used, with a number of iterations equal to 2000. The “Yap” correction [5] for turbulence model has also been used in this case, to help the convergence process in a flow supposed to be separated.

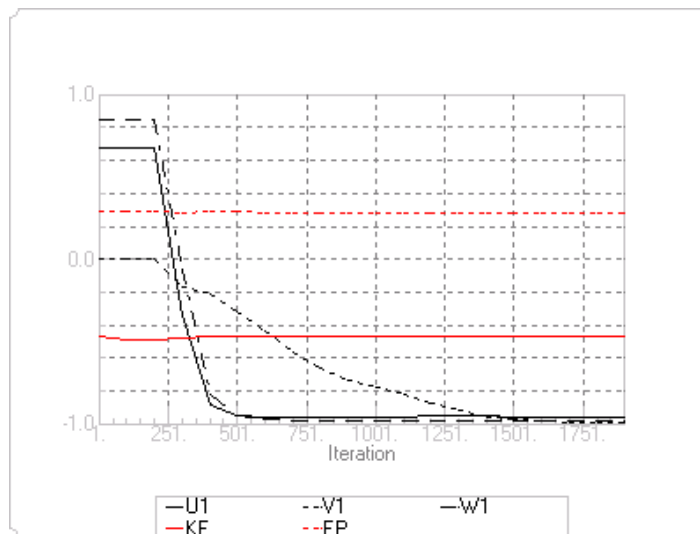


Figure 3.19: Residual values for case S5H4.

3.5.5 Results

Case S3H4

The following images show the solution of the flow field for the case S3H4. Figure 3.20 show the 2D speed scalar value calculated all over the model, while figure 3.21 shows the vertical velocity component, due to the shape of the terrain.

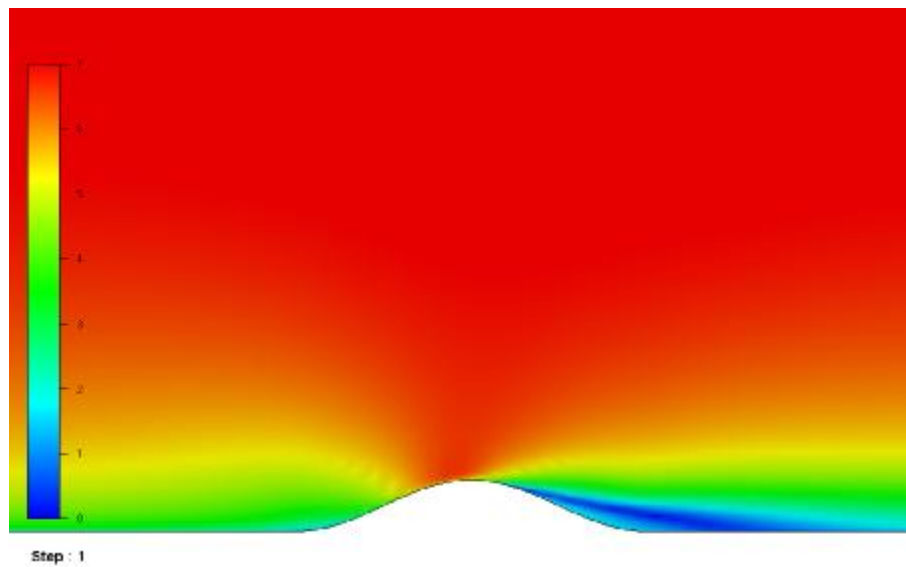


Figure 3.20: Speed 2D.

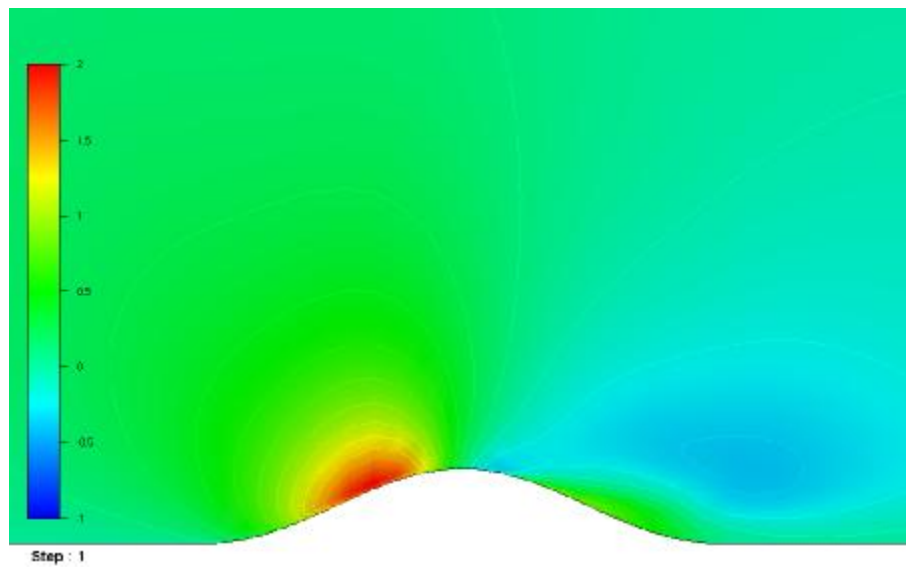


Figure 3.21: Z velocity component.

Finally, the horizontal x velocity component has been shown in the figure 3.22, in order to show the region of the model in which the recirculation of the flow can occur.

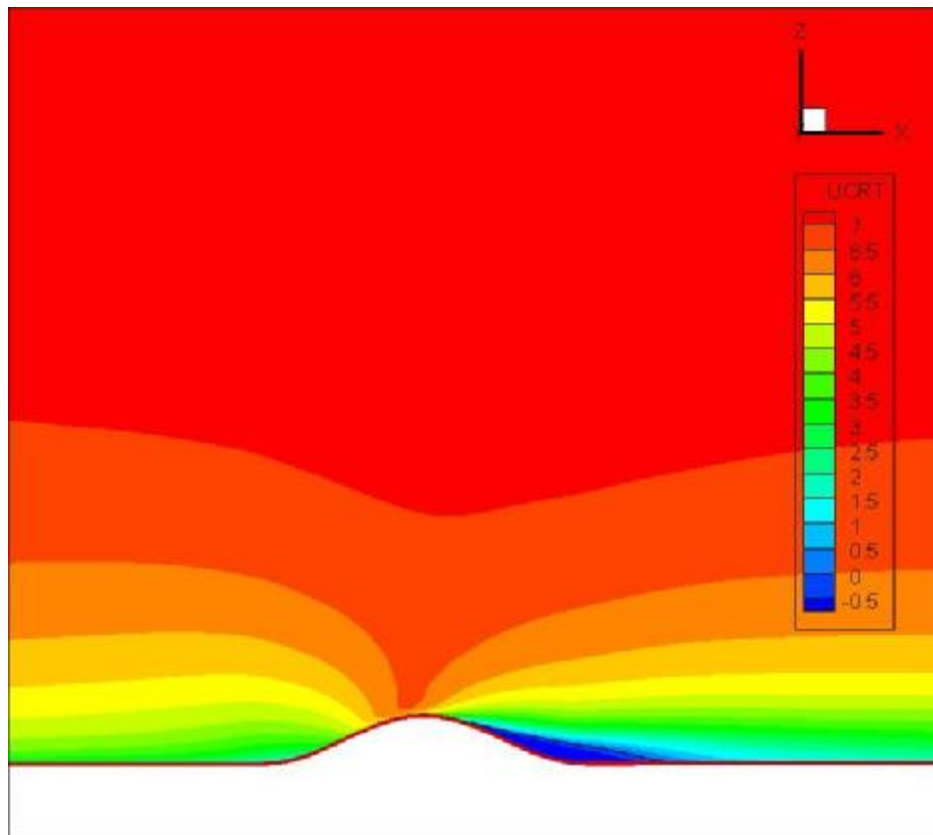


Figura 3.22: X velocity component.

The following figure 3.23 shows the comparison between experimental data, marked with the circle, and the results of the simulation, marked with the solid line. The solid circles laying on the ground line refer to the measurement points.

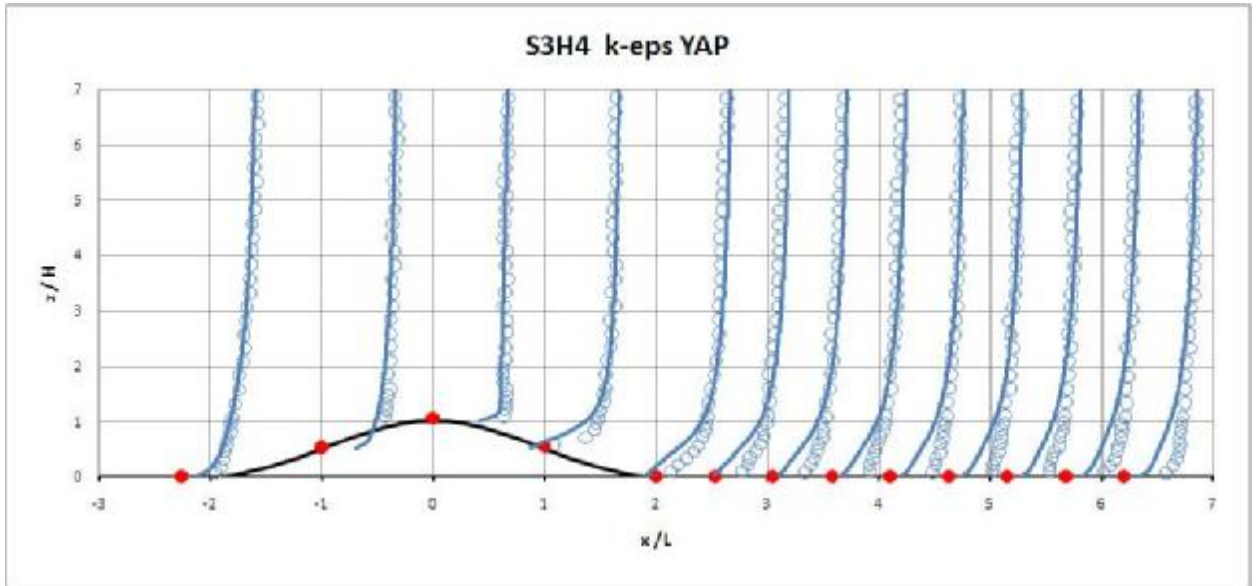


Figure 3.23: Comparison between experimental and numerical data.

Case S5H4

As well as in the previous case, the pictures below show the 2D speed scalar distribution, the vertical velocity component and the horizontal x velocity.

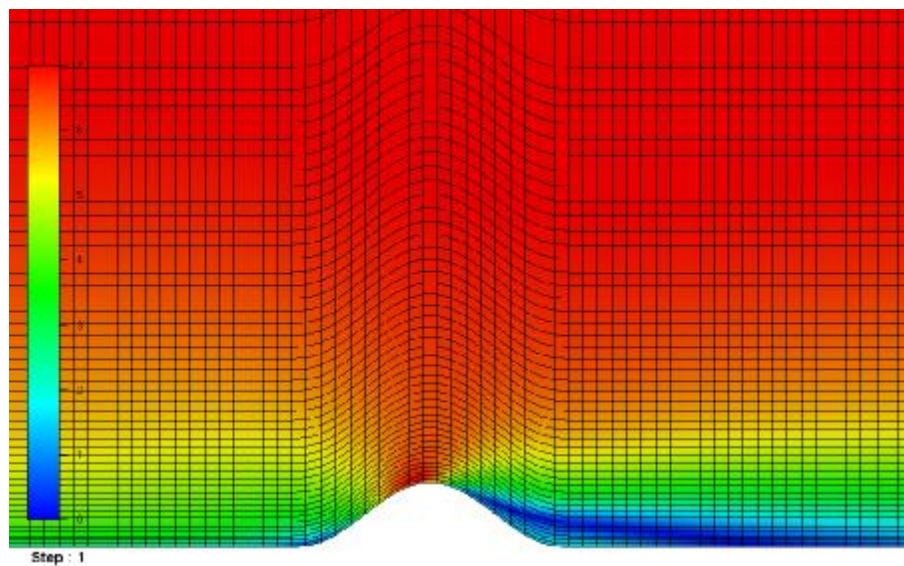


Figure 3.24: Speed 2D.

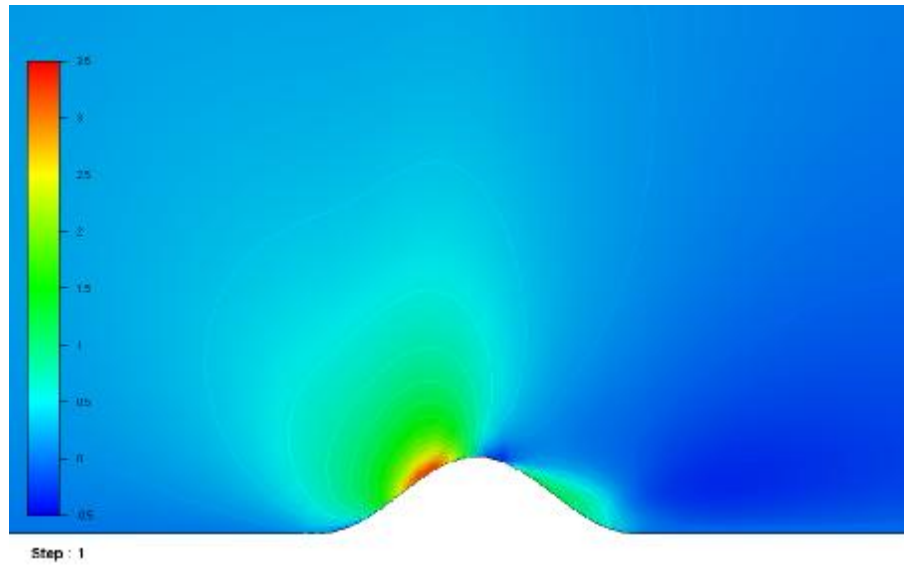


Figure 3.25: Z velocity component

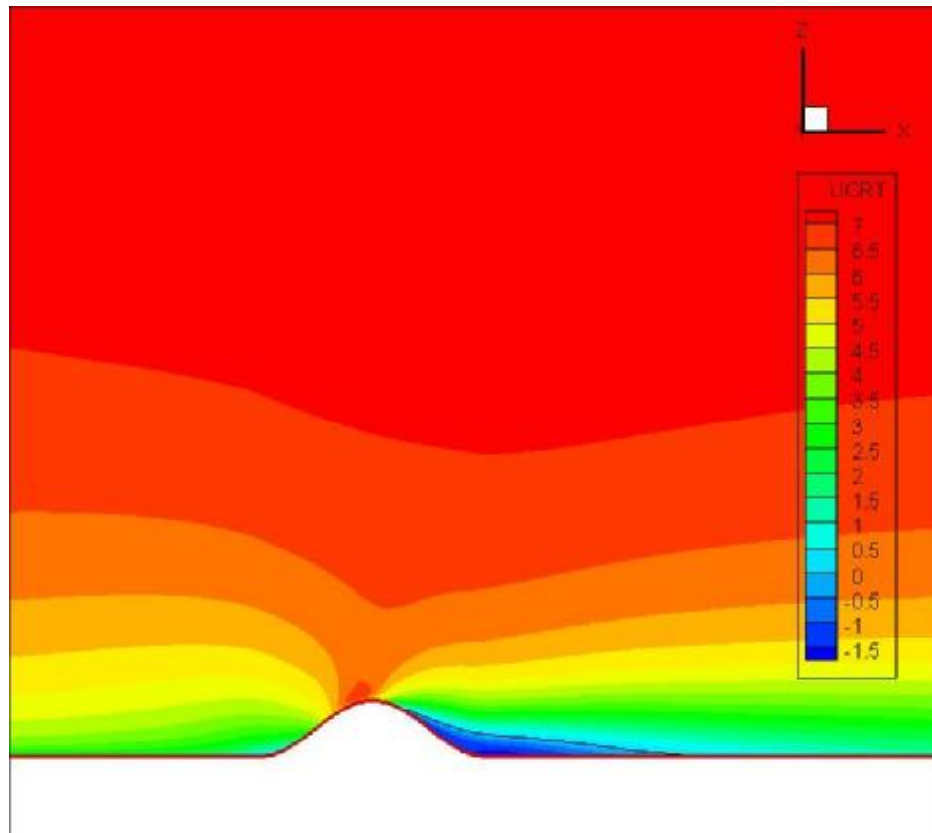


Figura3.26: X velocity component.

In the figure 3.27 a comparison between experimental data and results of the simulation is shown. Moreover, the flow recirculation area profile downstream the hill is represented by iso lines of x-velocity equal to 0.0, relative to both CFD with a low-Reynolds model by Kim et al. (1997) (thin line) and the presented results with k-ε YAP model (thick line).

The recirculation zone observed in the experiments is slightly thicker and shorter than the one calculated.

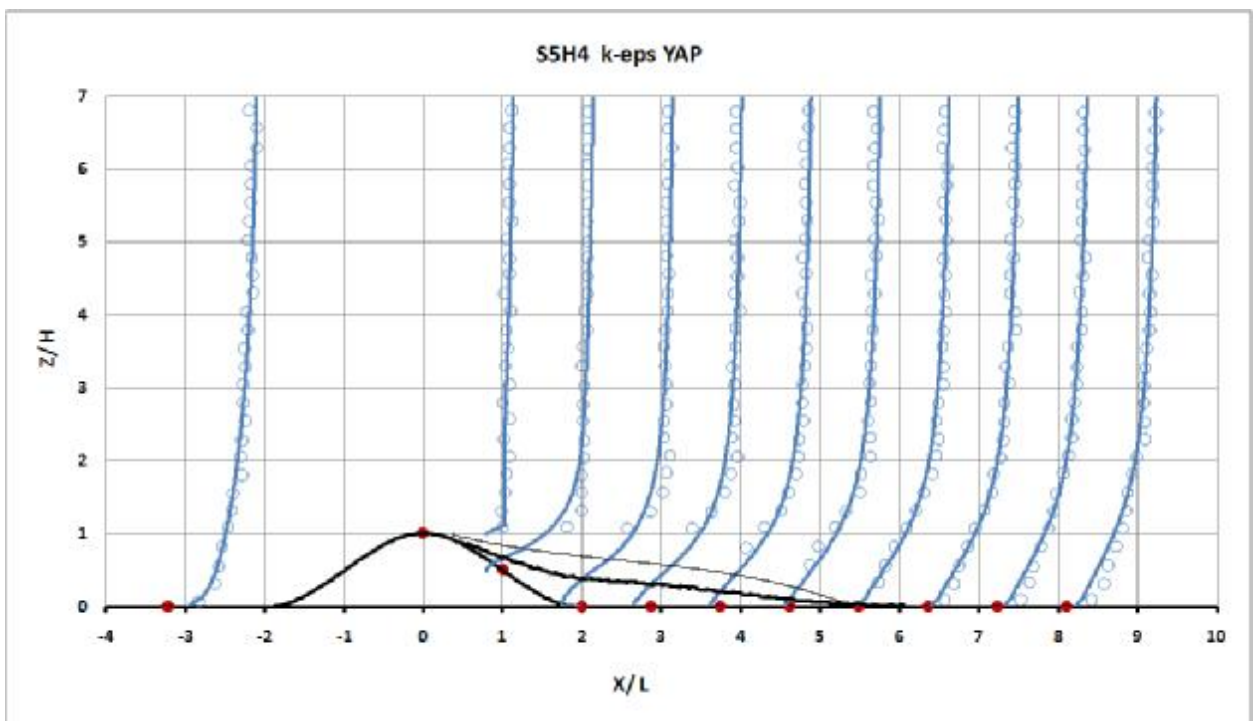


Figure 14: Comparison between experimental and numerical data.

Table 3.2: Value of the reattachment point.

	Kim et al.	CFD
	-	k-ε YAP
x/L	5,35	6,05

3.6 Remarks

A CFD study of a boundary layer flow over 2D rough hills has been performed. The aim of the presented simulations was to reproduce the experimental wind tunnel results by Kim et al. (1997), and set the parameters of WindSim to run in presence of hilly terrains. Both the experimental and numerical results show the different behaviors of the flow depending on the shape of the hill. Moreover, both the hills with slope $S 0,3$ and $S 0,5$ cause a separation of the flow, generating a recirculation area downstream of the hill.

The case S5H4 shows a good agreement between experimental and numerical results. The performed simulation over predicts the length of the separated area by 13%; the overestimation of the recirculation zone is probably due to the usage of the YAP correction [5] for the $k-\varepsilon$ closure. The $k-\varepsilon$ model coupled with the hybrid discretization scheme on a BFC grid would estimate a recirculation bubble smaller than the one of the experiments, according to the conclusions by Kim et al. (1997) . To solve this case, the segregated solver was good enough to reproduce the behavior of the flow. The turbulence has been modeled with a standard $k-\varepsilon$ model and the “Yap” correction has been used.

Concerning the case S3H4, in which separation has not been observed in the experiments, a very small recirculation zone appeared in the numerical results, in the region downstream of the hill.

This fact should be linked to the turbulence level of the flow set at the inlet of the tunnel: the modification of the turbulence intensity is not allowed in WindSim when setting the inlet boundary condition; then, it's possible that in this simulation the turbulence level wasn't high enough to maintain the flow attached to the terrain. Anyway, no specification on the turbulence level was found in the reference [2] for this very complex case, which has been observed to be at the boundary between the attached flow and occurrence of separation.

In all cases a segregated solver with the “Yap” correction for $k-\varepsilon$ turbulence model was also needed to produce a good convergence of the field variables.

CHAPTER 4. Characteristics of the site.

4.1 Introduction

The site studied in this document is located in a region of Central Italy and is visible in the following satellite image.



Figure 4.1: Aerial picture of the Antrodoco site.

It can be seen that the region is in a mountain area, covered by thick forest for a wide zone. Rocks and stones are also present with some agricultural plowed areas. Two roads are signed in the maps, running along the valley colored in yellow. Moreover, some urban areas are found in the zone, mostly small mountain villages; one of this is an important town, in proximity of the conjunction of the two roads.

Some experimental wind data of this site are available, due to the presence of two measurement masts, placed in the plateau at the center of the map. The positions of the two masts are signed by the red marker and specified by the UTM32 (ED50) coordinates.

4.2 Digital Terrain Model (DTM)

4.2.1 Orographic characteristics of the site

The orographic and roughness data of the terrain can be set in WindSim through a digital map, in a ****.gws* format. The map was prepared by Enel S.p.a., Italy, and contains information regarding both the elevation and the value of the roughness height. The following figure 4.2 shows the contour x-y map of the site.

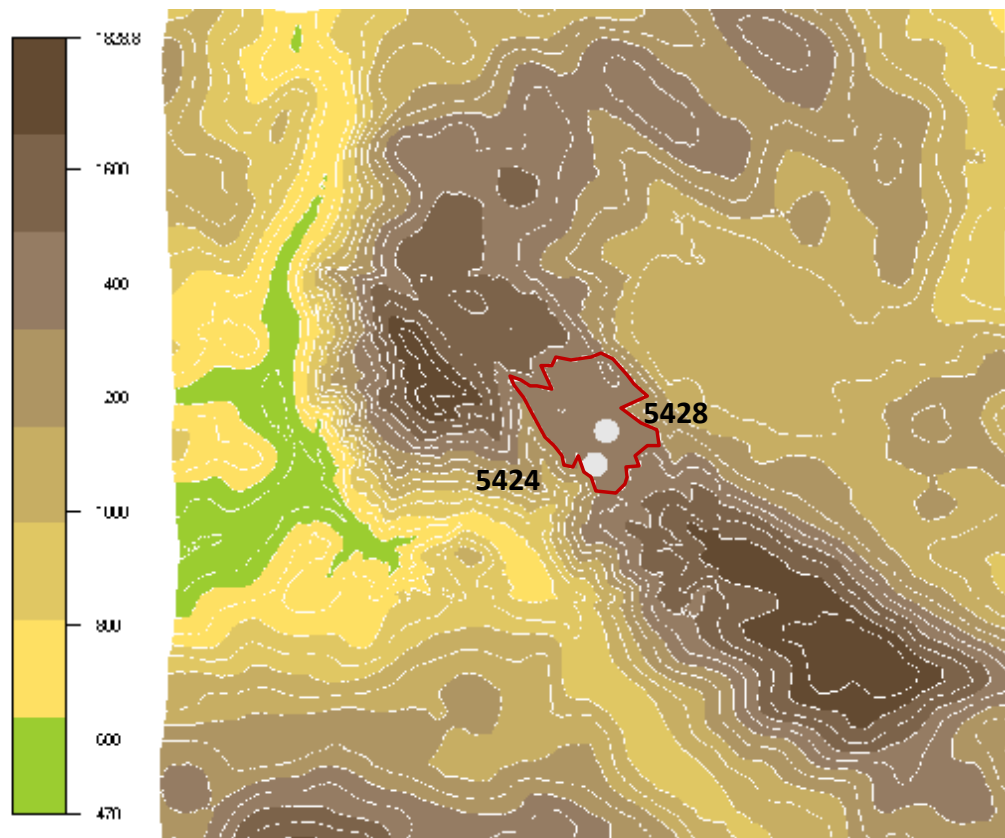


Figure 4.2: Orography map of the site.

The map highlights the plateau in the center of the whole area(enclosed in a red curve), where the two measurement masts are placed. A tridimensional view can also be obtained in WindSim, in order to offer a better interpretation of the Orography to the user.

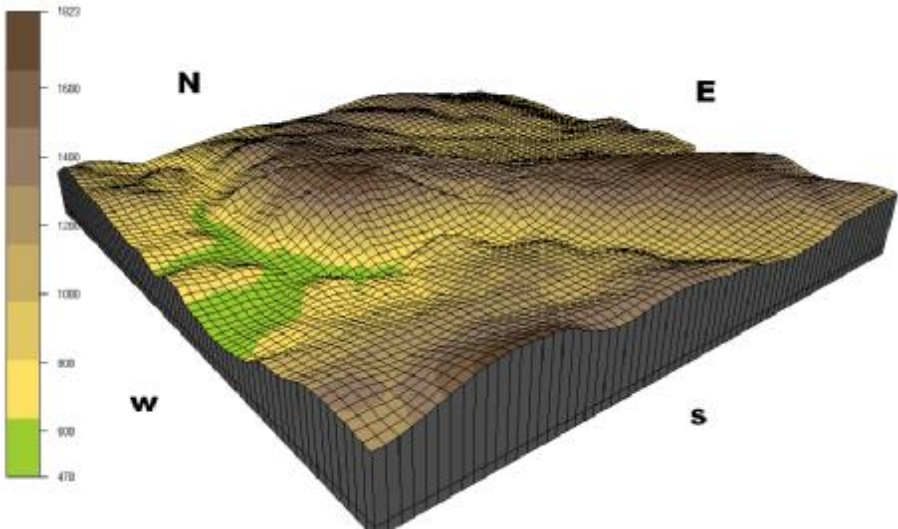


Figure 4.3: Tridimensional view of the site from South West.

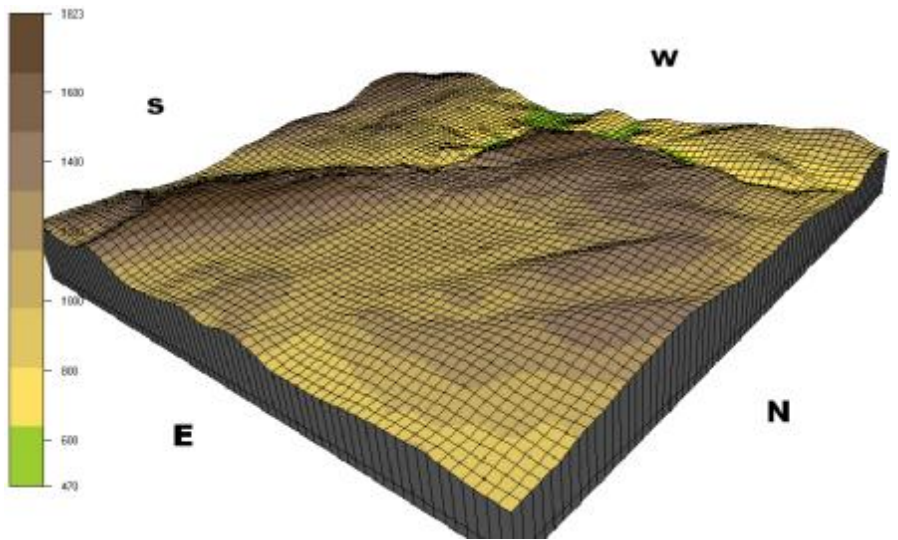


Figure 4.4: Tridimensional view of the site from North East.

Both the pictures 4.3 and 4.4 have been built and colored maintaining the same height scale which ranges from 470 m to 1823 m.

Notice that in direction NE-SW the quote of the terrain grows very slowly to the top of the rocky mountain. Over the top, a deep valley is present, where the step between the maximum and the minimum height is in order of 1400 meters.

4.2.2 Roughness characteristics

The whole area has been subdivided into three main roughness regions, characterized by three different values of roughness heights. The roughness classes are explained in the table 4.1.

Table 4.1: Roughness values assigned to the site.

DESCRIPTION	Z_0 [m]
Rocks, stones an plowed terrain	0,03
Forest	0,65
Urban areas	0,80

The figure 4.5 represents the roughness contour map of the region.

The plateau is all characterized by a roughness value of $Z_0 = 0,03$ m ; in fact no trees and other kind of vegetation is present in this region.

The Windsim software allows the comparison of the actual contour map of roughness with a satellite aerial picture of the site. The result of this procedure is reported in the further figure 4.6.

To allow the overlay, in the roughness colors with paints that contrast the typical colors of a satellite picture are used in the figure and listed in table 4.2.

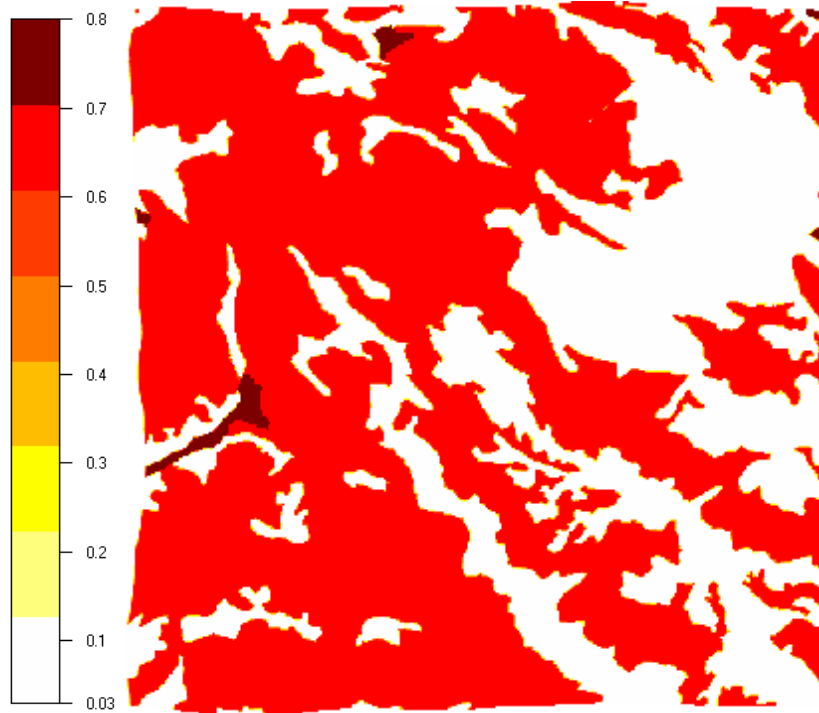


Figure 4.5: Roughness map of the site.

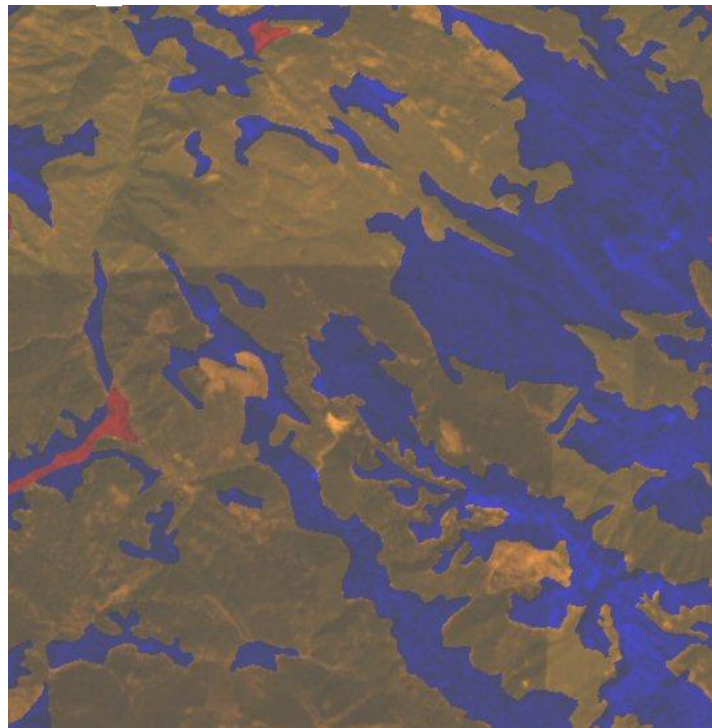


Figure 4.6: Overlay between roughness map and aerial picture of the site.

Table 4.2: Description of the colors used in the overlaid map.

DESCRIPTION	COLOR
Rocks, stones and plowed terrain	Blue
Forest	Transparent
Urban areas	Red

The analysis of the previous picture shows a good agreement between the roughness map and the satellite picture, demonstrating that the whole area is well described. Anyhow, there is evidence of some particular zones which are not exactly described by the contour lines. In fact, the yellow color of the picture (relative to the satellite image) should be totally covered by the blue color (relative to the roughness map), equal to $Z_0 = 0,03$ m.

Some disagreement in the zones far from the plateau is not important and errors in the roughness should not influence in a strong way the final results in the area of interest. Nevertheless, the better the area very close to the plateau is described the more precise the result are.

A more detailed comparison between the two maps can be accomplished by the image 4.7, showing the roughness contours drawn upon the satellite image at the maximum possible resolution (30 meters).

It is confirmed that some zones are not covered by vegetation as conversely represented in the roughness map. Despite of this slight disagreement, the roughness map has not been modified.

Finally, a further consideration should be done: the satellite map, extracted from Google Earth software, can contain irregularities and be some years old.

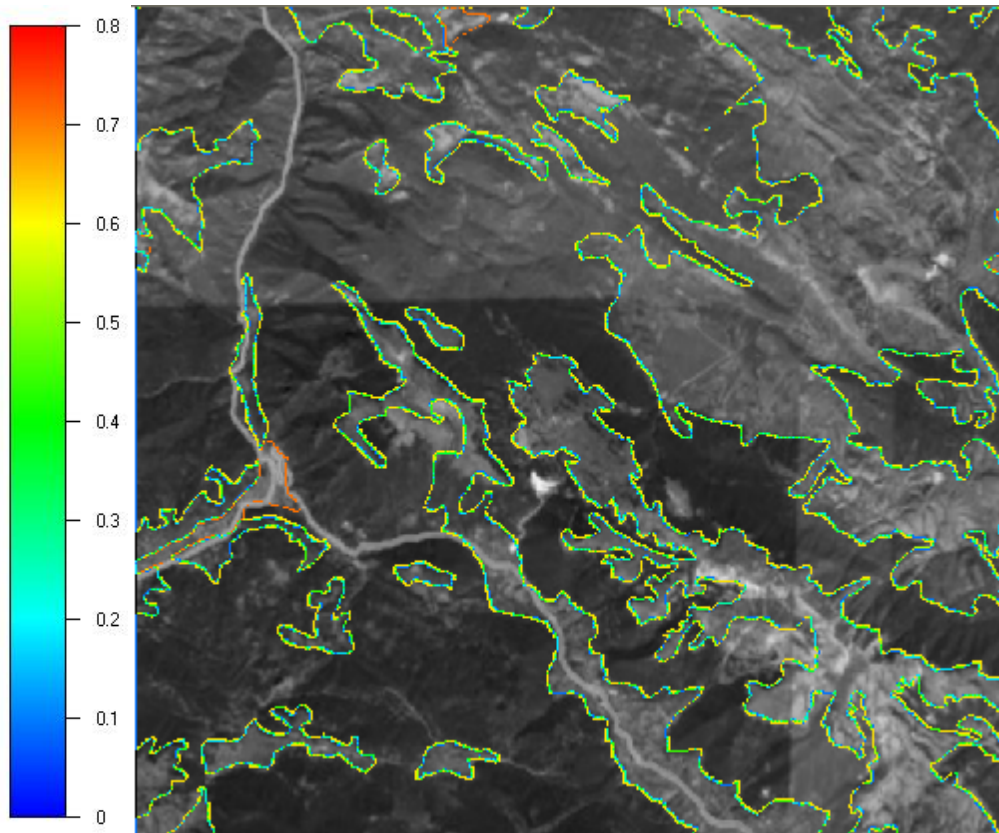


Figure 4.7: Roughness contours drawn upon satellite picture of the site.

4.3 Wind data

Two measurement masts are available in the plateau, both containing multiple points of measurement. Particularly, station number 5424 has two stations, one of them 15 meters height and the other one 30 meters height above ground level. Concerning station number 5428, it has 3 stations respectively at 30, 40 and 50 meters height referred to the ground level.

Available wind data are not continue in time along the year, but are relative to two measurement periods, precisely the period November-December 2006 and April-May 2007.

In this periods, data are recorded by the anemometer's data logger, and stored in a data file like the one shown in table 4.3, where data are averaged in a 10 minutes period.

The figure 4.8 show the 3 cup NRG #40 anemometer and the NRG #200P wind direction vane used to measure the wind characteristics. Detailed informations about their technical specifications can be found in bibliography [8].



Figure 4.8: NRG #40 anemometer and NRG #200P wind direction vane.

Table 4.3: Example of a wind data file.

<i>Day</i>	<i>Month</i>	<i>Year</i>	<i>Hour</i>	<i>Minute</i>	<i>Mean speed</i>	<i>Max speed</i>	<i>Standard deviation</i>	<i>Direction</i>	<i>Standard deviation</i>
-	-	-	[H]	[min]	[m/s]	[m/s]	[m/s]	[degree]	[degree]
1	4	2007	0	30	0,3	2,0	0,5	74	2
1	4	2007	0	40	0,8	1,5	0,5	75	0
1	4	2007	0	50	0,9	1,5	0,3	75	0
1	4	2007	0	60	1,3	2,0	0,2	75	0
1	4	2007	1	10	1,6	2,3	0,3	70	8
1	4	2007	1	20	1,8	2,3	0,3	73	18

Figures 4.9 and 4.10 show a 3D representation of the two measurement stations, the position and display the orography of the terrain close to them.

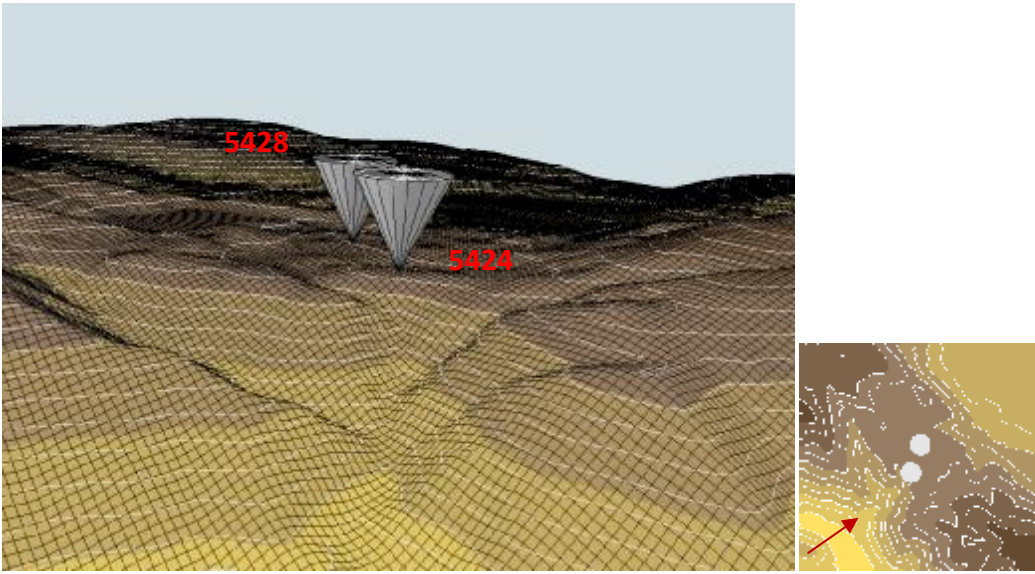


Figure 4.9: Site position of the mast measurement. View from South-West.

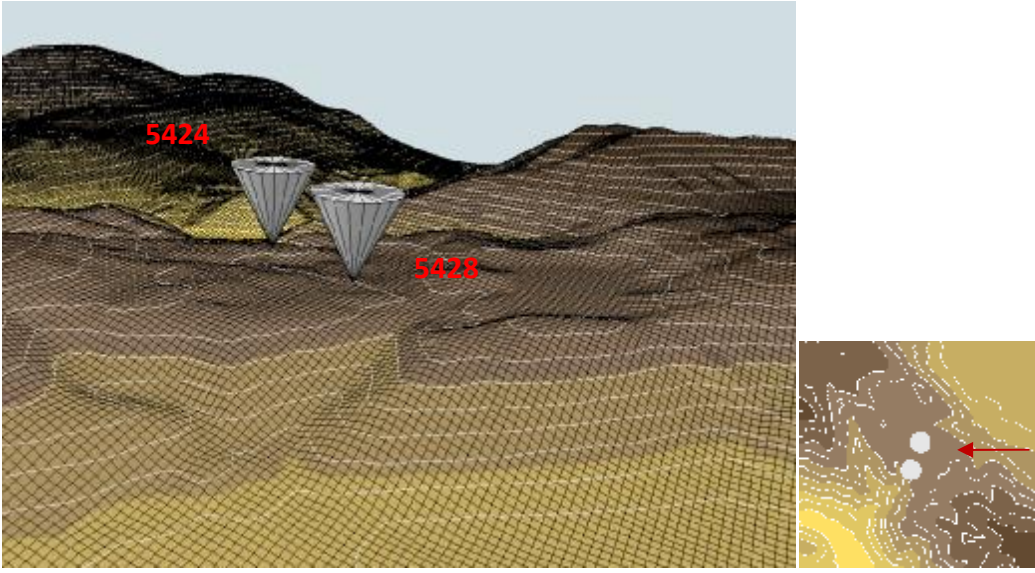


Figure 4.10: Site position of the mast measurement. View from East.

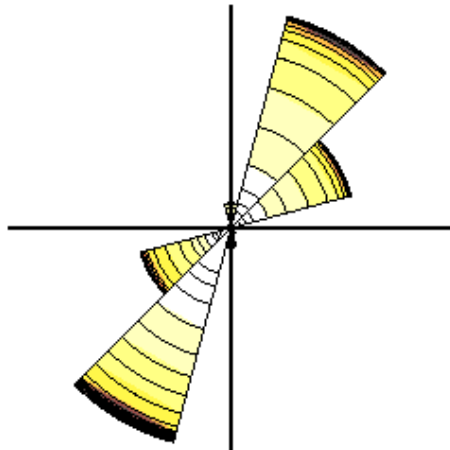
In the following pages, a short description of every single data set is presented, reporting the wind frequency rose and the main results which describe the Weibull fitting of the recorded data.

All experimental data have been filtered, eliminating the errors recorded by the data logger during bad working of the anemometer.

Moreover, the dominant sectors have been highlighted in yellow; referring to the other sectors, the ones which frequency is below 1,5 % have been highlighted in green, in order to facilitate the understanding of the processed WindSim results, presented in chapter 6.

In these sectors the calculated statistic parameters are not good enough because the number of samples used to analyze them is low.

4.3.1 Mast 5424-30 m Period April-May 2007



RESULTS

Total records	8614
Average wind speed	4,79 m/s
Shape parameter K	1,84
Scale parameter A	5,77 m/s
Median	4,73 m/s
Mode	3,77 m/s

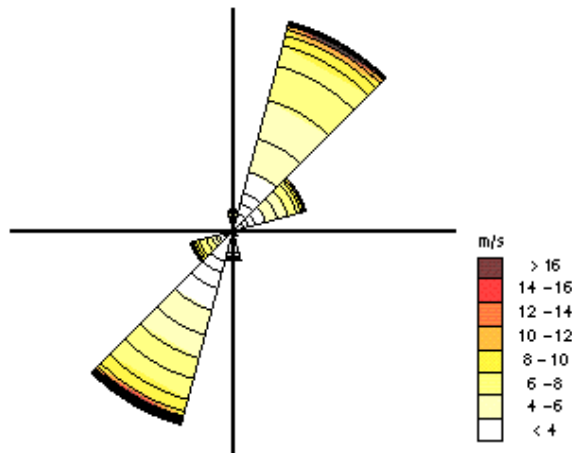
Table 4.4: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5424-30 April-May												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,4	2,42	2,59	1,32	1,08	0,77	0,91	1,4	2,57	0,84	0,64	1,51
A	4,1	5,95	6	2,6	1,31	0,94	1,57	5,46	7,29	1,09	0,57	3
freq	0,034	0,299	0,17	0,011	0,01	0,009	0,026	0,303	0,128	0,006	0,003	0,005
mean	3,39	5,12	5,12	2,33	1,27	1,33	1,73	4,63	5,93	1,34	1,15	2,59



Figure 4.11: Frequency of records and Weibull fitting.

4.3.2 Mast 5424-15 m Period April-May 2007



RESULTS

Total records	8614
Average wind speed	4,59 m/s
Shape parameter K	1,93
Scale parameter A	5,58 m/s
Median	4,61 m/s
Mode	3,82 m/s

Table 4.5: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5424-15 April-May												
	0	30	60	90	120	150	180	210	240	270	300	330
k	1,95	2,6	2,23	0,85	0,87	0,69	0,72	1,75	2,4	1,26	0,72	1,8
A	3,87	6,15	5,1	1,03	0,77	0,56	0,76	5,82	6,84	1,28	0,77	2,73
freq	0,034	0,354	0,123	0,012	0,01	0,01	0,05	0,328	0,072	0,003	0,003	0,003
mean	3,18	5,23	4,27	1,27	0,92	0,93	1,15	4,86	5,63	1,23	1,2	2,3

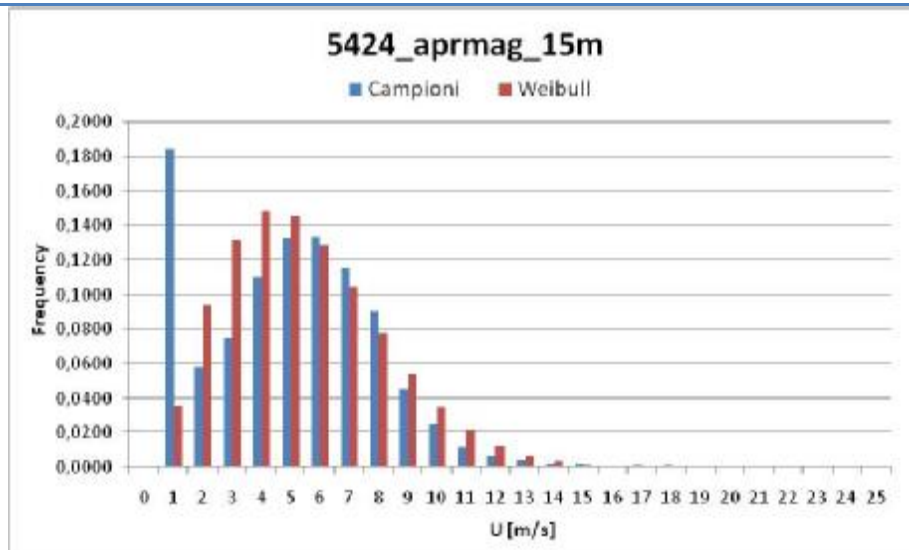
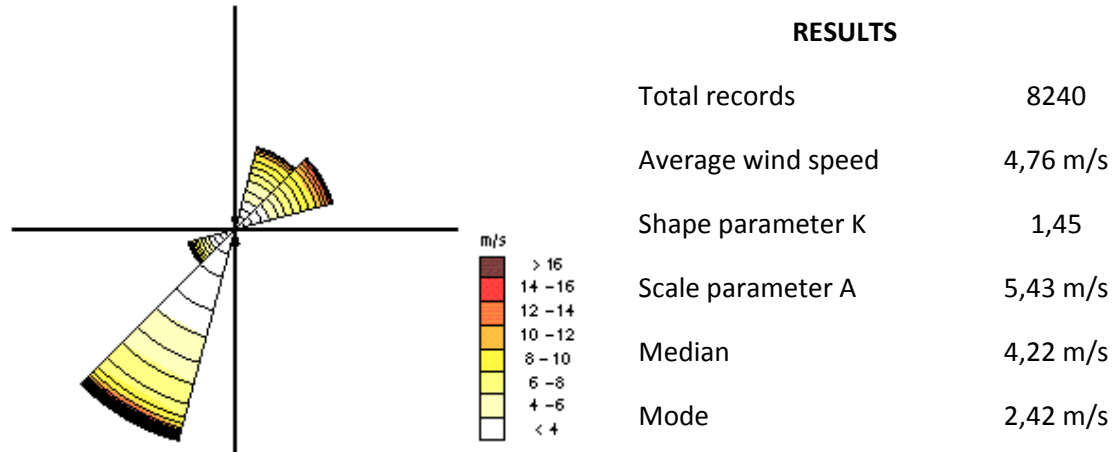


Figure 4.12: Frequency of records and Weibull fitting.

4.3.3 Mast 5424-30 m Period November-December 2006



RESULTS	
Total records	8240
Average wind speed	4,76 m/s
Shape parameter K	1,45
Scale parameter A	5,43 m/s
Median	4,22 m/s
Mode	2,42 m/s

Table 4.6: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5424-30 November-December												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,15	2	1,92	0,75	1,82	1,04	1,1	1,31	1,32	1,13	0,71	0,74
A	4,04	6,32	7,08	0,89	0,71	1,05	2,69	5,11	4,71	0,73	0,63	0,67
freq	0,026	0,17	0,202	0,01	0,01	0,007	0,032	0,436	0,097	0,007	0,004	0,006
mean	3,35	5,47	6,01	1,2	0,63	1,15	2,61	4,59	4,24	0,68	1,05	1,03

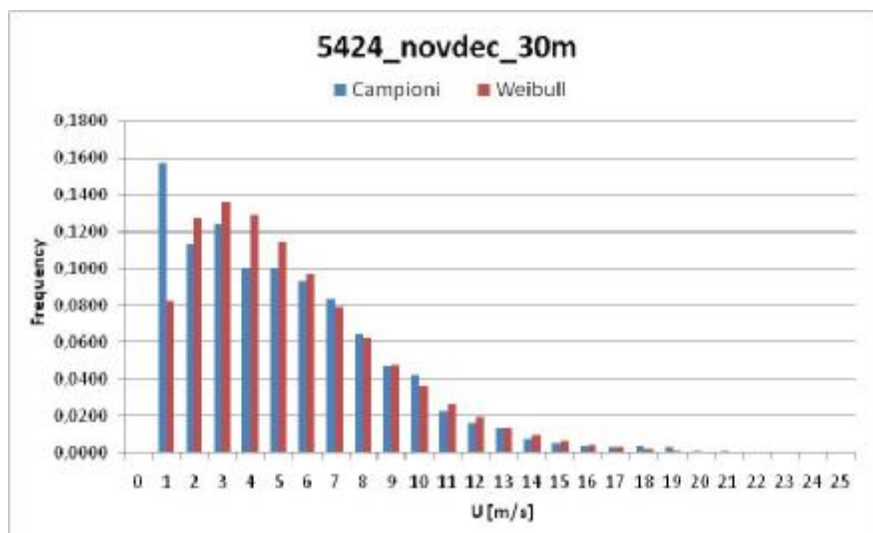
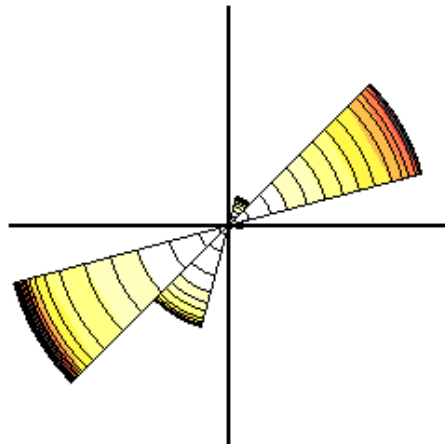


Figure 4.13: Frequency of records and Weibull fitting.

4.3.4 Mast 5424-15 m Period November-December 2006



RESULTS

Total records	8240
Average wind speed	4,93 m/s
Shape parameter K	1,54
Scale parameter A	5,52 m/s
Median	4,35 m/s
Mode	2,80 m/s

Table 4.7: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5424-15 November-December												
	0	30	60	90	120	150	180	210	240	270	300	330
k	1,34	1,88	2,06	1,29	1,93	1,17	1,27	1,51	1,58	1,19	0,98	1,04
A	2,45	4,81	7,22	2,61	0,68	1,14	1,62	3,85	5,76	1,95	0,99	1,03
freq	0,009	0,05	0,332	0,022	0,01	0,006	0,01	0,174	0,369	0,012	0,003	0,005
mean	2,29	4,18	6,26	2,37	0,59	1,15	1,44	3,56	5,14	1,81	1,07	1,08

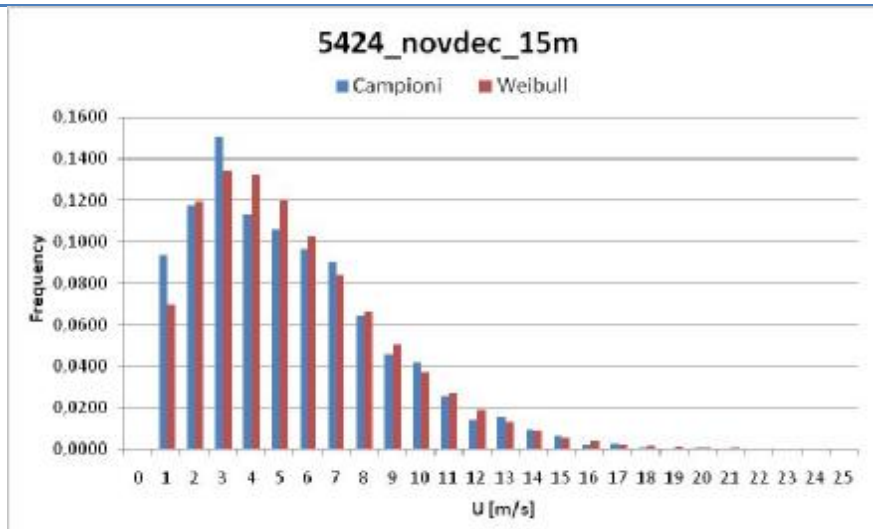


Figure 4.14: Frequency of records and Weibull fitting.

4.3.5 Mast 5428-50 m Period April-May 2007

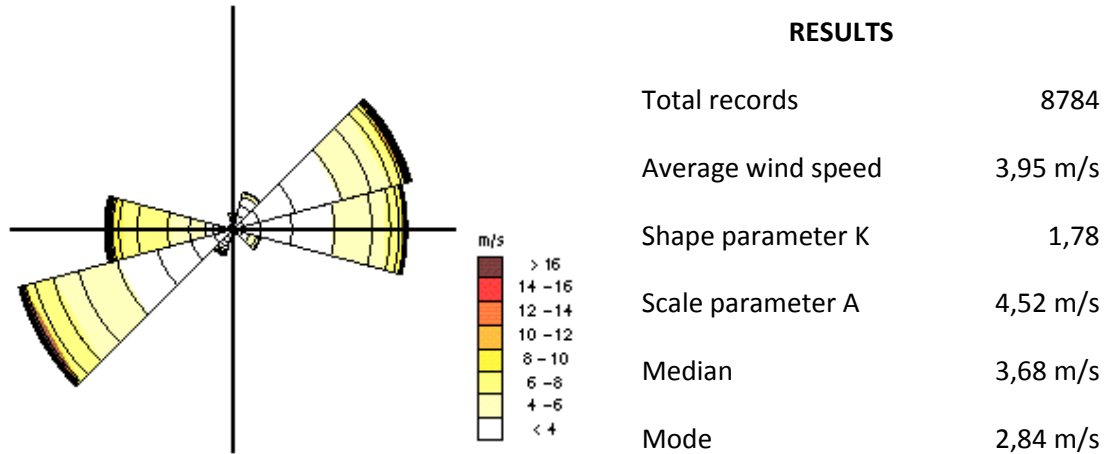


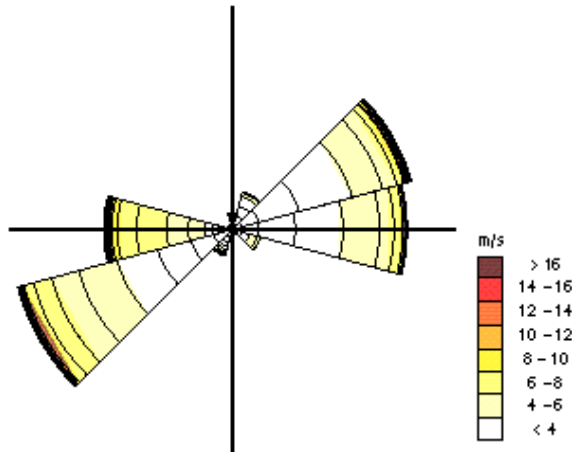
Table 4.8: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5428-50 April-May												
	0	30	60	90	120	150	180	210	240	270	300	330
k	1,85	2,12	1,68	2,25	3,02	1,33	1,41	1,28	1,85	2,77	1,44	1,31
A	2,83	2,98	4,26	4,3	4,09	2,03	1,96	2,87	4,59	6,8	2,12	1,92
freq	0,019	0,046	0,219	0,206	0,03	0,006	0,009	0,033	0,263	0,152	0,007	0,006
mean	2,36	2,53	3,87	3,75	3,49	1,78	1,74	2,63	3,96	5,71	1,88	1,72



Figure 4.15: Frequency of records and Weibull fitting.

4.3.6 Mast 5428-40 m Period April-May 2007



RESULTS

Total records	8784
Average wind speed	3,86 m/s
Shape parameter K	1,80
Scale parameter A	4,40 m/s
Median	3,59 m/s
Mode	2,80 m/s

Table 4.9: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5428-40 April-May												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,05	2,27	1,83	2,39	3,22	1,43	1,5	1,29	1,91	2,82	1,51	1,18
A	3,12	3,05	4,04	4,09	3,98	1,92	1,87	2,88	4,6	6,76	2,26	1,66
freq	0,019	0,046	0,219	0,206	0,03	0,006	0,009	0,033	0,263	0,152	0,007	0,006
mean	2,59	2,61	3,64	3,58	3,4	1,7	1,63	2,65	3,96	5,68	1,93	1,59

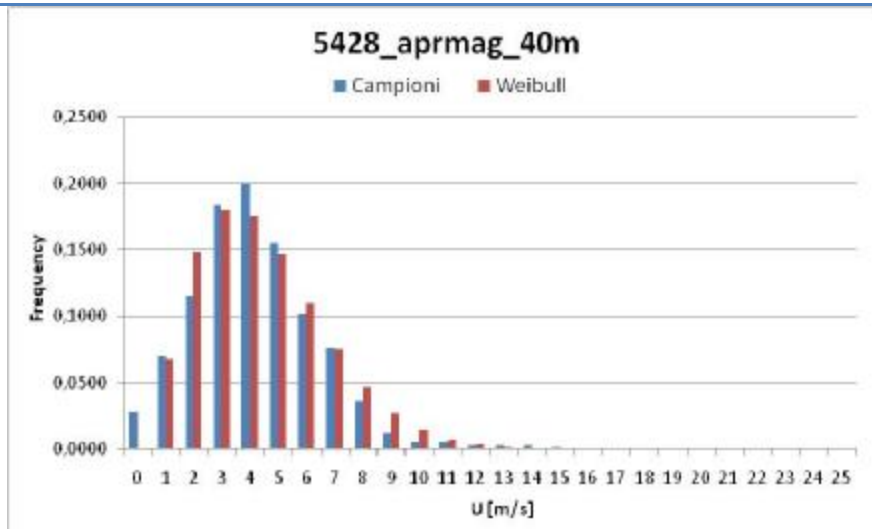
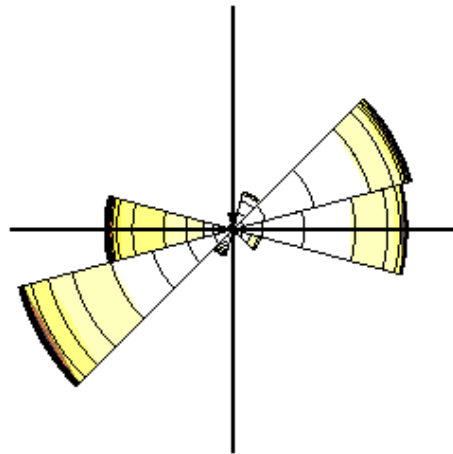


Figure 4.16: Frequency of records and Weibull fitting.

4.3.7 Mast 5428-30 m Period April-May 2007



RESULTS

Total records	8784
Average wind speed	3,62 m/s
Shape parameter K	1,78
Scale parameter A	4,13 m/s
Median	3,36 m/s
Mode	2,60 m/s

Table 4.10: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5428-30 April-May												
	0	30	60	90	120	150	180	210	240	270	300	330
k	1,87	2,15	2,03	2,57	3,33	1,3	1,55	1,32	1,92	2,85	1,57	1,19
A	2,66	2,62	3,7	3,82	3,85	1,72	1,82	2,71	4,4	6,56	2,25	1,66
freq	0,019	0,046	0,219	0,206	0,03	0,006	0,009	0,033	0,263	0,152	0,007	0,006
mean	2,24	2,24	3,26	3,33	3,28	1,56	1,6	2,46	3,79	5,52	1,91	1,57

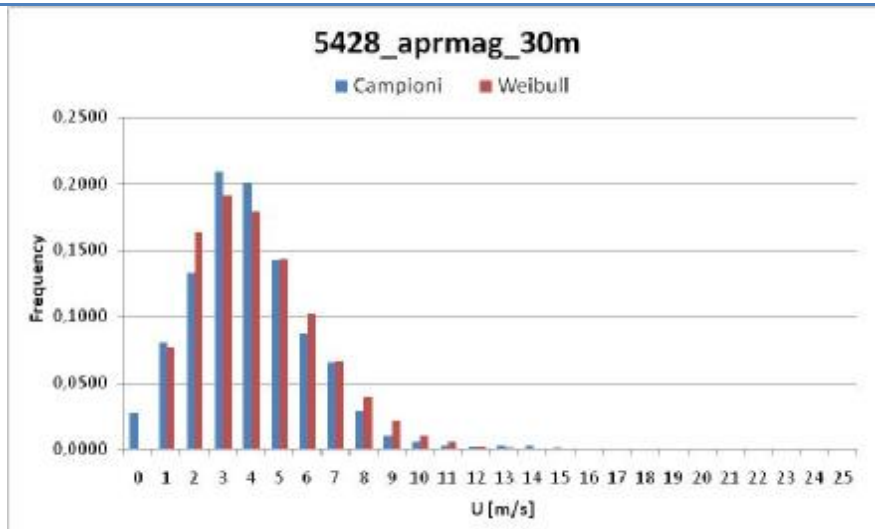
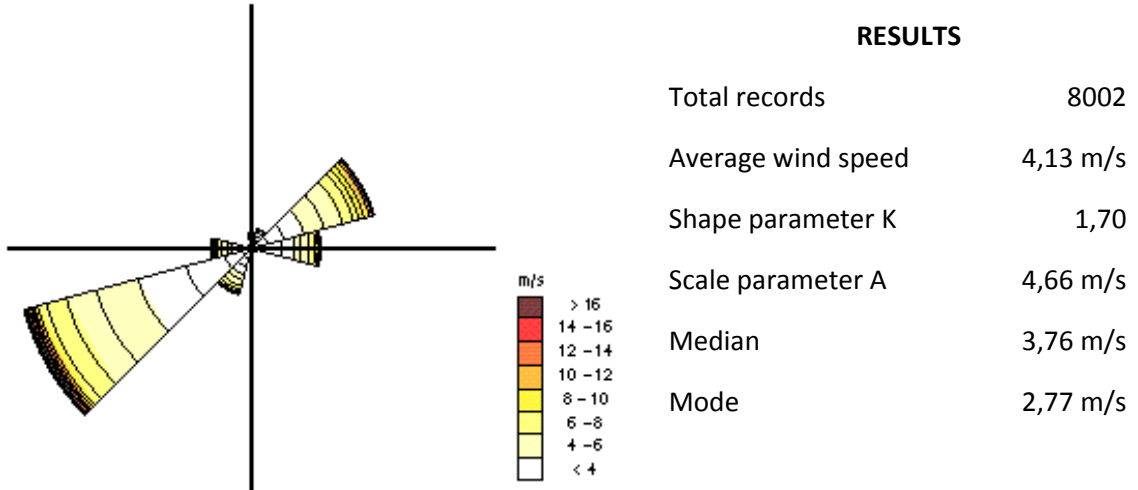


Figure 4.17: Frequency of records and Weibull fitting.

4.3.8 Mast 5428-50 m Period November-December 2006



RESULTS

Total records	8002
Average wind speed	4,13 m/s
Shape parameter K	1,70
Scale parameter A	4,66 m/s
Median	3,76 m/s
Mode	2,77 m/s

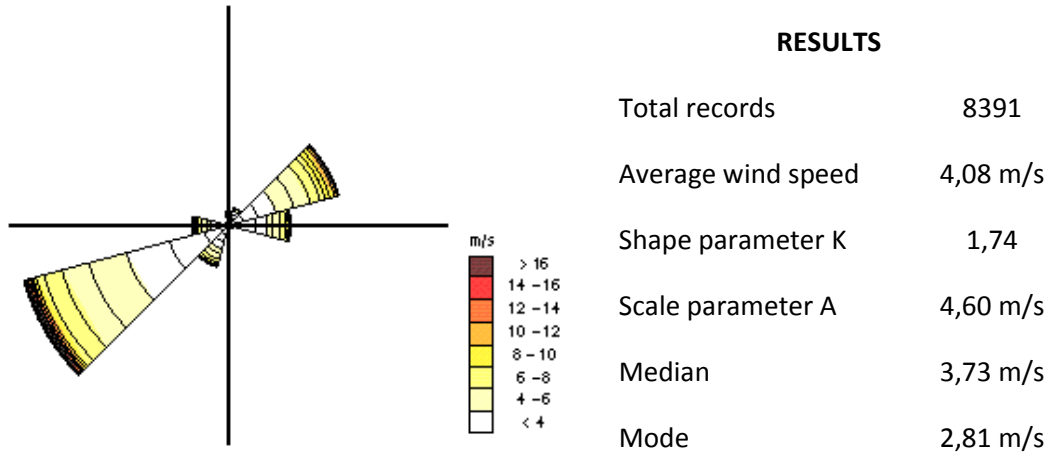
Table 4.11: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5428-50 November-December												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,37	2,06	1,94	1,73	1,13	1,35	1,15	1,87	1,69	1,72	1,03	1,05
A	2,85	2,87	5,31	4,16	1,72	0,76	1,27	4,46	4,86	5,07	1,54	1,2
freq	0,026	0,035	0,221	0,121	0,01	0,003	0,006	0,083	0,411	0,071	0,005	0,007
mean	2,4	2,39	4,7	3,77	1,69	0,69	1,29	3,92	4,39	4,4	1,53	1,21



Figure 4.18: Frequency of records and Weibull fitting.

4.3.9 Mast 5428-40 m Period November-December 2006



RESULTS	
Total records	8391
Average wind speed	4,08 m/s
Shape parameter K	1,74
Scale parameter A	4,60 m/s
Median	3,73 m/s
Mode	2,81 m/s

Table 4.12: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

5428-40 November-December												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,4	1,98	2	1,72	1,29	1,37	1,15	1,92	1,76	1,74	1,07	1,12
A	2,99	2,95	5,26	4,14	1,93	0,79	1,22	4,3	4,79	4,96	1,6	1,24
freq	0,026	0,037	0,23	0,119	0,01	0,003	0,006	0,08	0,408	0,07	0,004	0,007
mean	2,54	2,46	4,66	3,77	1,76	0,72	1,23	3,76	4,31	4,31	1,55	1,23

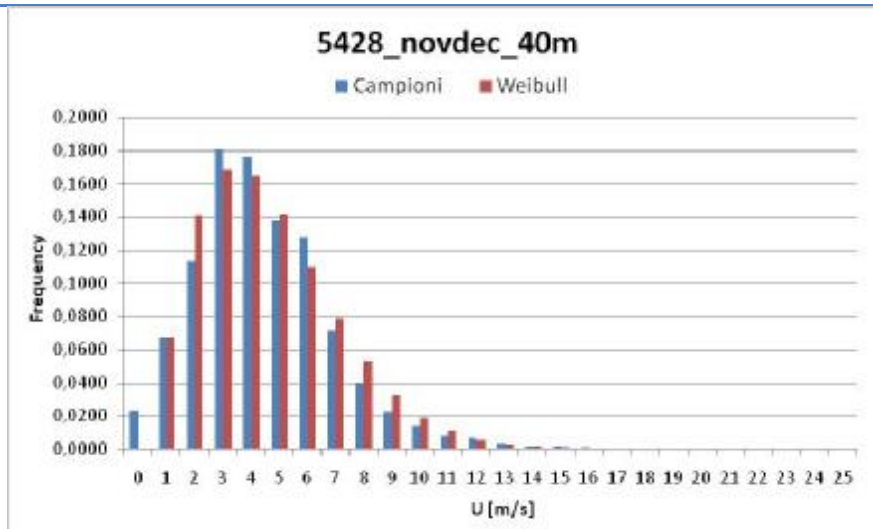
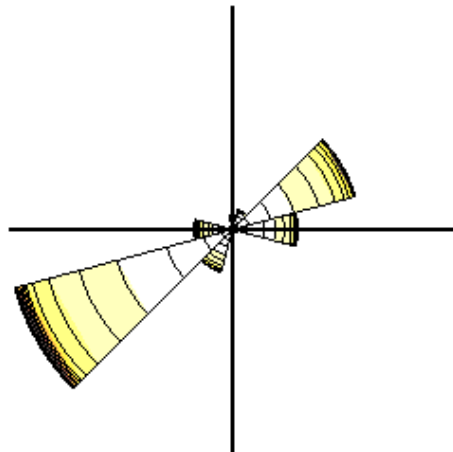


Figure 4.19: Frequency of records and Weibull fitting.

4.3.10 Mast 5428-30 m Period November-December 2006



RESULTS

Total records	8391
Average wind speed	3,84 m/s
Shape parameter K	1,81
Scale parameter A	4,37 m/s
Median	3,57 m/s
Mode	2,80 m/s

4.13: Weibull (k,A), frequency (-) and average wind speed (m/s) versus sector.

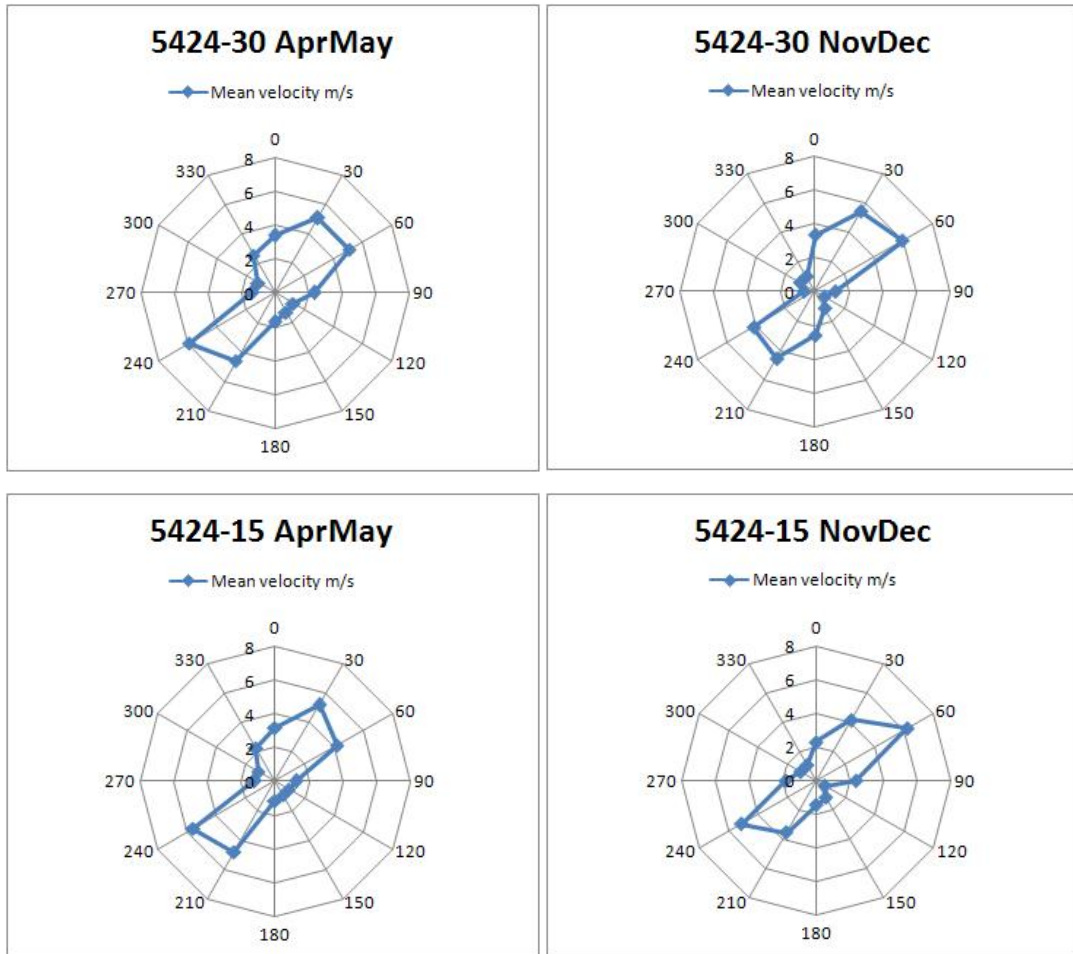
5428-30 November-December												
	0	30	60	90	120	150	180	210	240	270	300	330
k	2,24	1,98	2,23	1,86	1,23	1,73	1,17	1,94	1,77	1,73	1,17	0,89
A	2,58	2,81	4,69	3,87	1,71	0,77	1,12	4,18	4,67	4,84	1,74	0,91
freq	0,026	0,037	0,23	0,119	0,01	0,003	0,006	0,08	0,408	0,07	0,004	0,007
mean	2,18	2,34	4,13	3,45	1,61	0,69	1,13	3,62	4,2	4,17	1,58	1,11



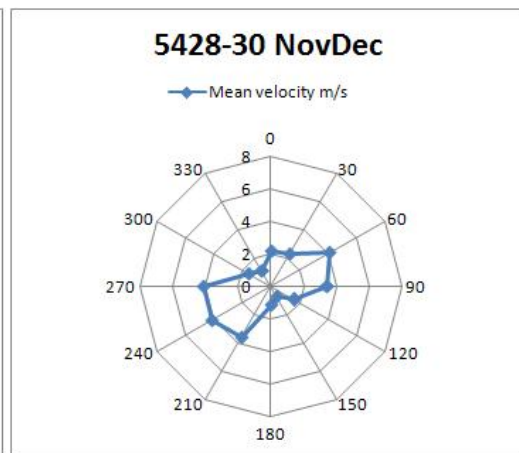
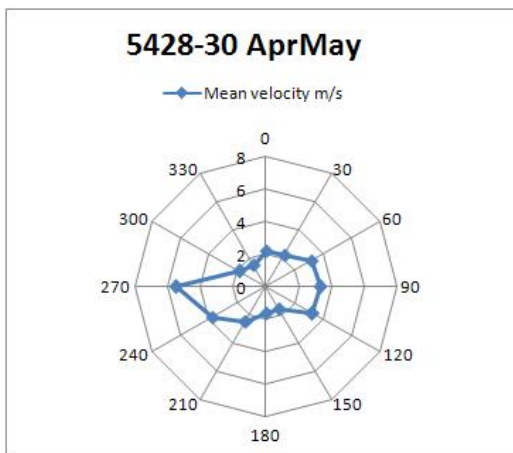
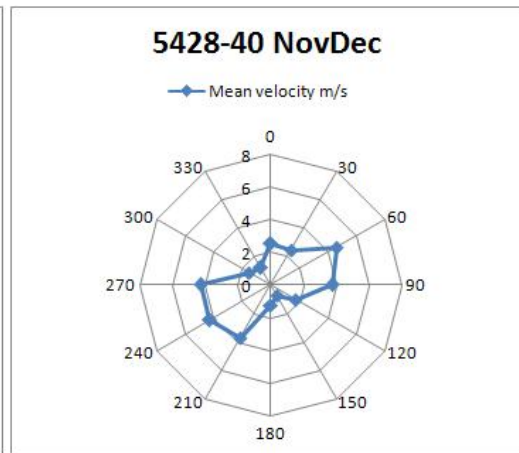
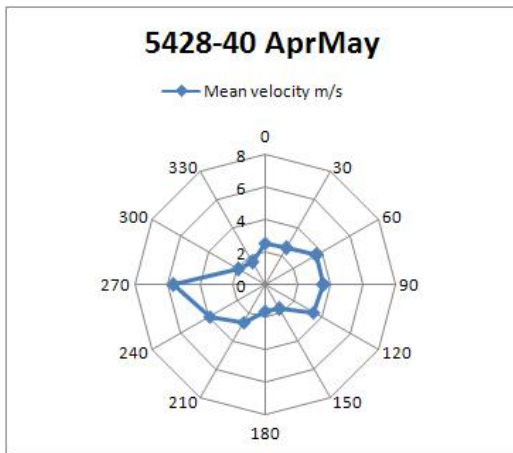
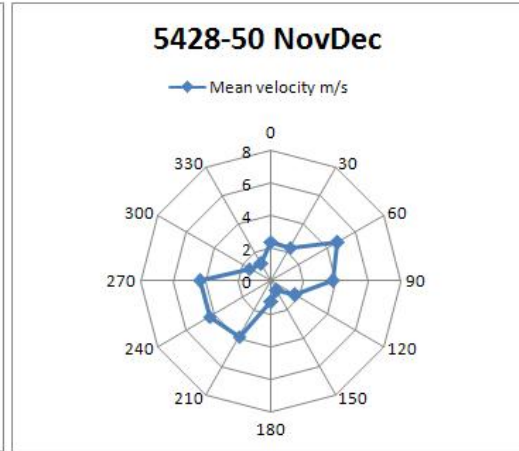
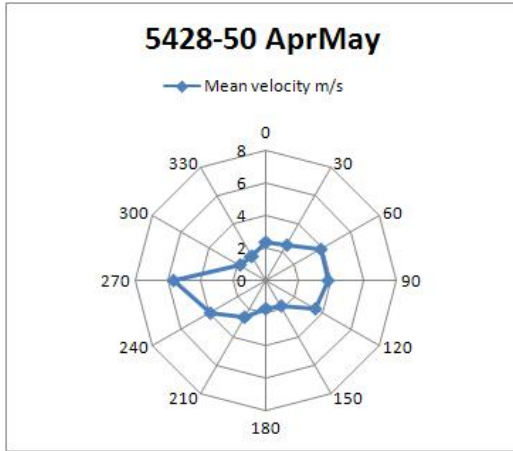
Figure 4.20: Frequency of records and Weibull fitting.

4.4 Mean Speed per sector tables

4.4.1 Mast 5424

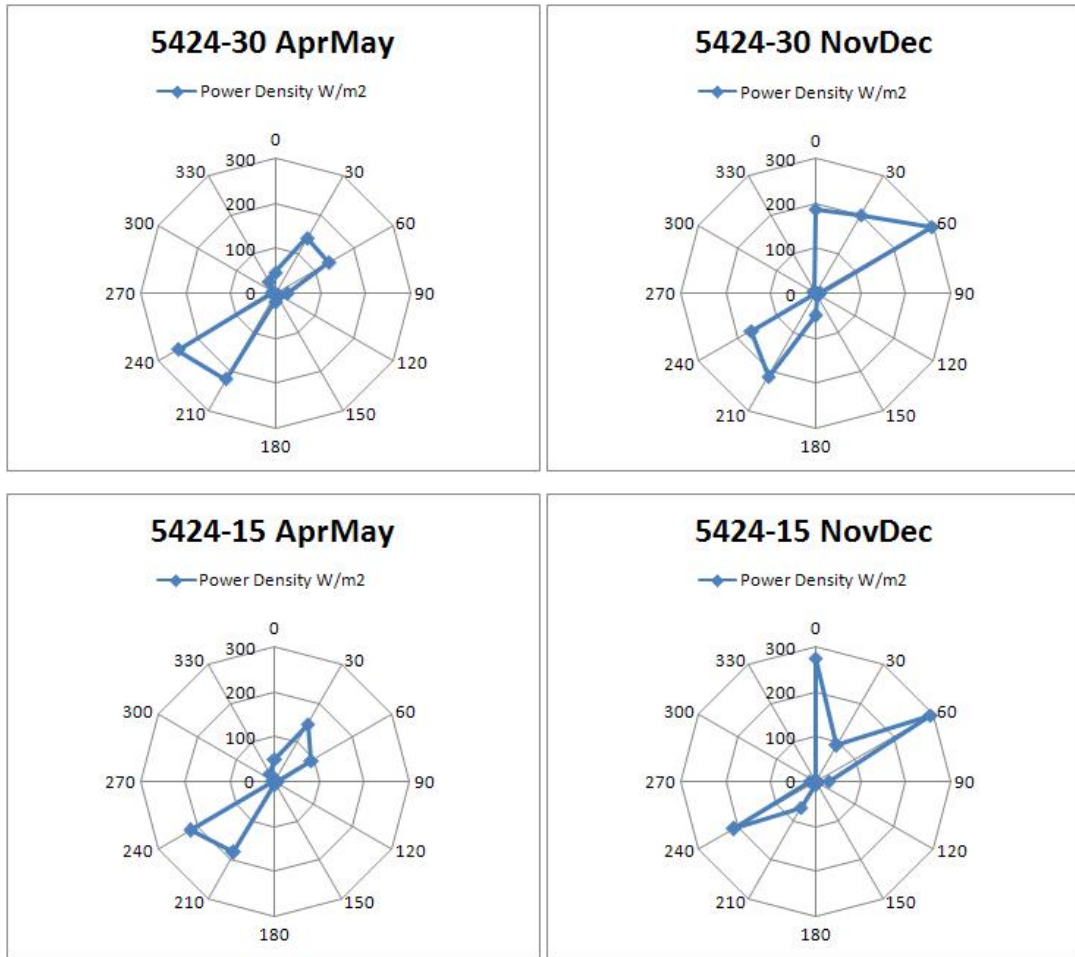


4.4.2 Mast 5428

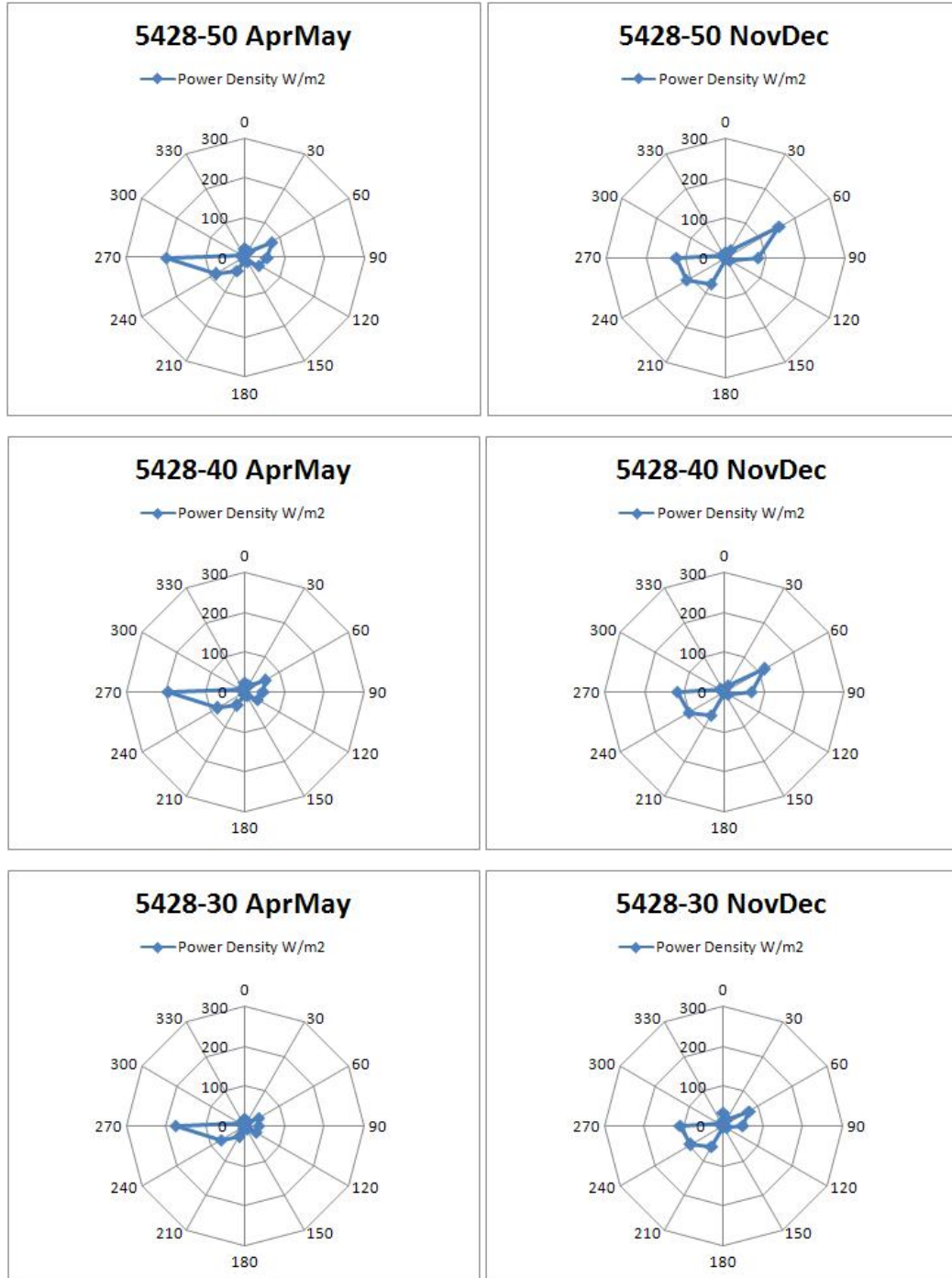


4.5 Power Density per sector tables

4.5.1 Mast 5424



4.5.2 Mast 5428



4.6 Remarks

The terrain characteristics of the site and wind experimental data have been presented.

The performed analysis allowed to highlight the dominant wind direction sectors, corresponding to 30 and 210 degrees, for the tower 5424, while 60 and 240 degrees, for the tower 5428.

The tower 5424 has shown a different behavior of the flow between the two measurement station, respectively 15 and 30 meters, in the period November-December. This fact can be related to some kind of thermal seasonal effects which can influence the flow in a region very close to the ground. It must be noticed, in this case, the position of the tower 5424 in proximity of a cliff, which can induce orographic effects in the flow.

According to the selected dominant sectors, a preview of the position of the wind turbines for a plant layout proposal can be done by considering to put them in a series of rows perpendicular to the main wind directions.

The choice should be confirmed by the results of the simulations which will be carry out in chapter 5 and 6.

CHAPTER 5. Test of the Free Stream Velocity (FSV)

5.1 Introduction

The purpose of this chapter is to find the right settings in order to describe the speed vertical profile of the flow given as a boundary condition. Particularly, the free stream velocity of the geostrophic wind has been studied. In fact, the scalar value of the free stream velocity can be set in WindSim by the user in the Windfield module.

The analysis of this input parameter has been done in three steps: first of all a mesoscale model of the considered site has been created, keeping a coarse grid resolution; then, as a second step, a smaller microscale model with medium grid resolution has been studied, in order to put in evidence the recirculation effect of the flow; finally, another microscale model with fine grid resolution has been built, finding as a result the optimum value of the free stream velocity to set in WindSim for further simulations.

The microscale models have been studied using the “nesting technique”: this procedure consists in running a simulation of the windfield in the mesoscale model and using this one to extract boundary conditions to be applied to the microscale models, which have to be created therefore in a zone inside the mesoscale site.

5.2 Mesoscale Model

The mesoscale model is presented in this section. The horizontal velocity field has been calculated; using data of the tower 5428 at 50 meters height to scale the results, the relative maps obtained with free stream velocities of 10 and 20 m/s are respectively shown in Figure 5.1 and 5.2.

It must be noticed that in proximity of the tower position the maps show a value equal to 1, demonstrating that the scale procedure is correct.

The maps have been extracted for 4 wind directions: 0, 90, 180, 270 degrees, at 30 and 50 meters height.

The test of the free stream velocity has been carried out for 4 scalar values: 1, 5, 10 and 20 m/s.

The general parameters set for the simulations are displayed in the following tables 5.1 and 5.2.

Geometrical model

Table 5.1: Input parameters of the geometrical model

INPUT PARAMETER	VALUE
Maximum number of cells	100000
Real number of cells	98532
X-resolution	210 m
Y-resolution	210 m
Number of cells in X-direction	67
Number of cells in Y-direction	60
Number of cells in Z-direction	20
Height distribution factor	0,1
Height of the model	4938 m
First cell minimum height	44,8 m
First cell maximum height	57,2 m

Numerical model

Table 5.2: Settings for the numerical model

BOUNDARY CONDITIONS	
Top	wall with no friction
Geostrophic height	500 m
Free stream wind speed	1-5-10-20 m/s
Wind direction	0-90-180-270 degrees
Turbulence model	standard k- ϵ
SOLVER SETTINGS	
Solver	Coupled
Number of iterations	50

Any important variation has been recognized, mostly in the zone of the plateau containing the two mast measurements.

Anyhow, it can be seen that there is a directional sector, corresponding to 90 degrees, in which some small variations in the wind field are more evident. In this

case the wind is blowing from the east sector, from the plateau perpendicular to the valley.

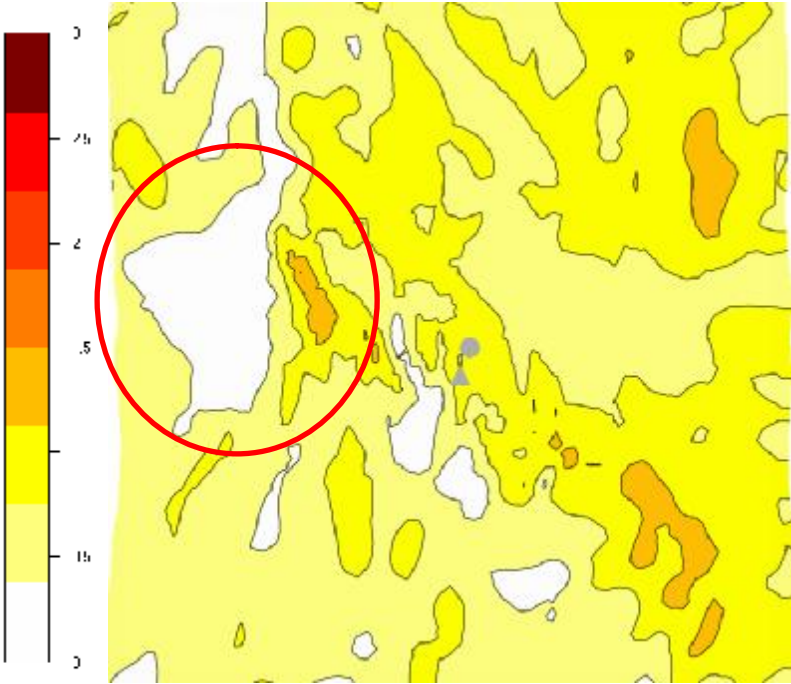


Figure 5.1: Speed-2D at 30 m height for sector 90 and velocity 10 m/s.

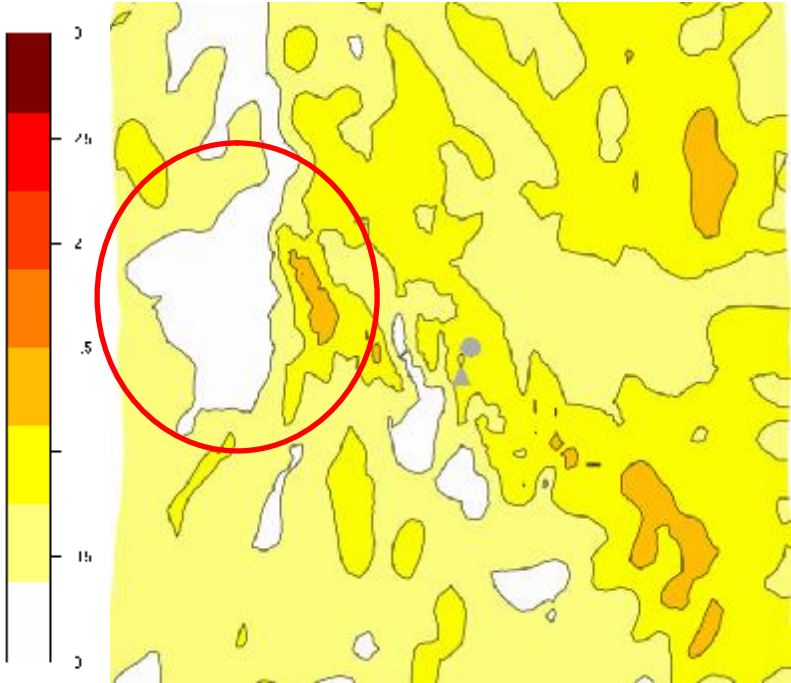


Figure 5.2: Speed-2D at 30 m height for sector 90 and velocity 20 m/s.

5.3 Microscale model – evidence of recirculation

The small variations highlighted in the mesoscale model for sector 90, placed in the west region of the site, led to a need for a deeper study in order to understand the causes that generate the differences in speed-up. A wise preview of the phenomenon can be done considering the specified wind direction and the orography map, shown in chapter 4. In fact, a deep valley is present in the west zone of the site, corresponding to the path of the roads signed in the map. In this region a recirculation effect due to high steepness of the west-side of the mountain can be generated when the wind is blowing from the east direction.

A smaller model has been created and the orography map is shown in the following picture 5.3.

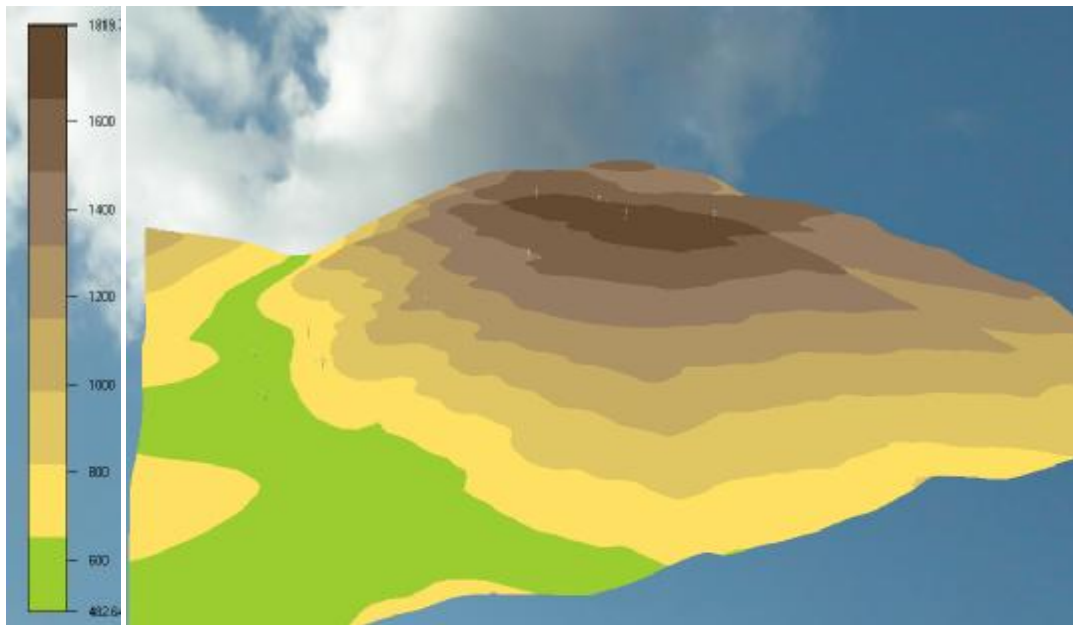


Figure 5.3: Orographic 3D map of the microscale model.

The microscale model covers an area of dimensions 6 X 6 km and has a grid resolution finer than the mesoscale model previously described. Boundary conditions are superimposed by the outer model (nesting technique).

The general parameters applied in the simulation, run for wind direction 90 degrees and wind speed 10 m/s, are exposed in the tables 5.3 and 5.4.

Geometrical model

Table 5.3: Input parameters for geometrical model in the mesoscale model.

INPUT PARAMETER	VALUE
Maximum number of cells	500000
Real number of cells	499392
X-resolution	26 m
Y-resolution	26 m
Number of cells in X-direction	135
Number of cells in Y-direction	135
Number of cells in Z-direction	25
Height distribution factor	0,1
Height of the model	5000 m
First cell minimum height	17,5 m
First cell maximum height	22,2 m

Numerical model

Table 5.4: Settings for the numerical model.

BOUNDARY CONDITIONS	
Top	Wall with no friction
Geostrophic height	From mesoscale model (500 m)
Free stream wind speed	From mesoscale model (10 m/s)
Wind direction	90 degrees (East)
Turbulence model	Standard k- ϵ
SOLVER SETTINGS	
Solver	Coupled
Number of iterations	500

Some vertical profiles have been extracted in the positions described in the picture 5.4 below to understand the behavior of the flow above the microscale refined model.

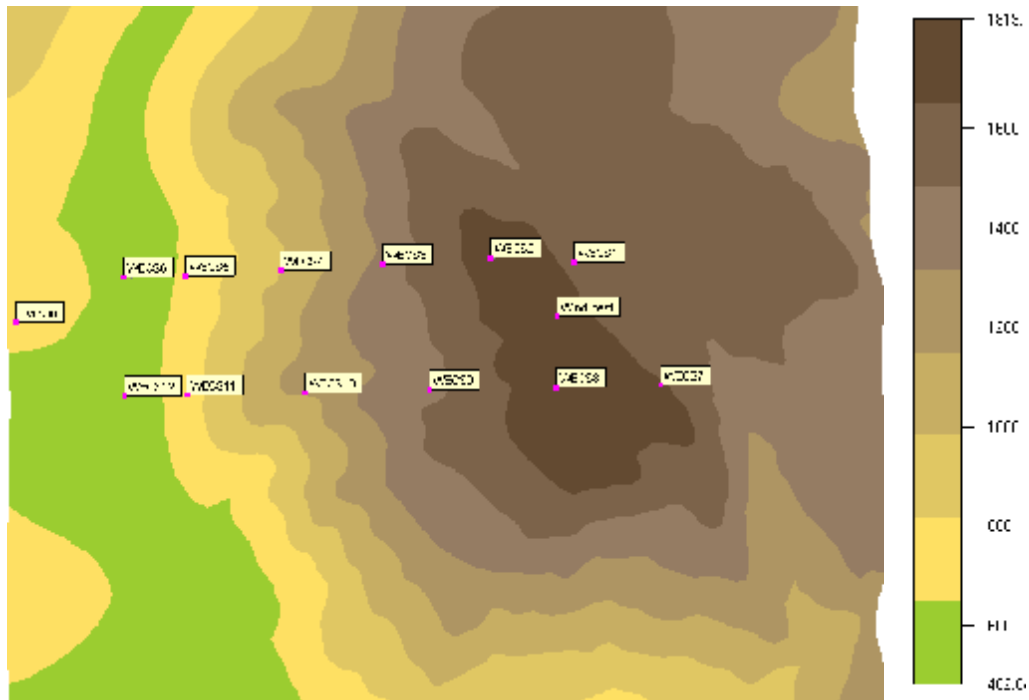


Figure 5.4: Site of extraction of vertical profiles.

The analysis of the microscale model has shown the presence of a big recirculation bubble, downstream of the mountain with wind direction of 90 degrees, extending for 500 meters above ground level, as shown in pictures 5.5 and 5.6.

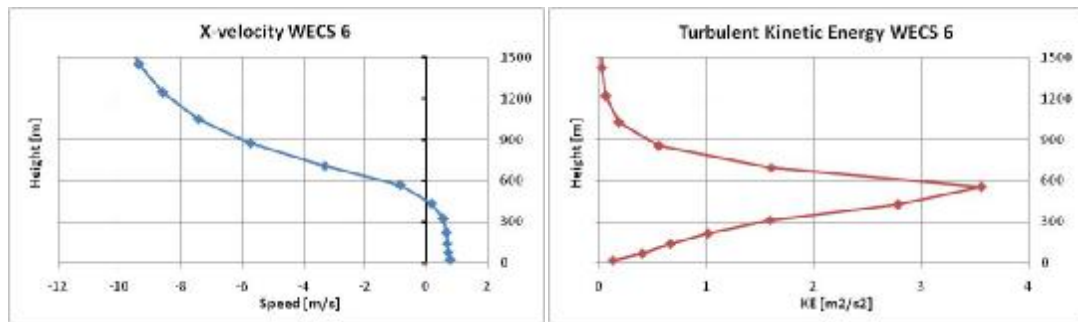


Figure 5.5: Vertical profiles in position 6.

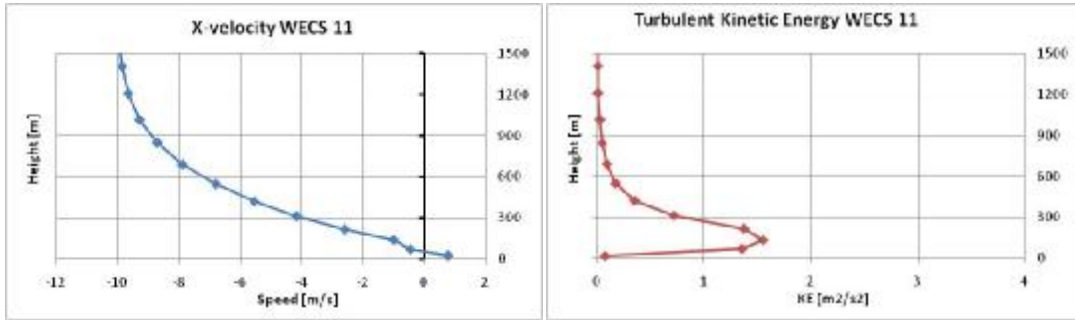


Figure 5.6: Vertical profiles in position 11.

The presence of the flow separation region can be connected to the variations in speed-up maps recognized in the mesoscale model, the separated flow obtained for the 90 degrees sector run is highlighted by means of streamlines colored by velocity magnitude reported in the figure 5.7 .A more accurate analysis of the microscale model has been also done, this will be presented in the next paragraph.

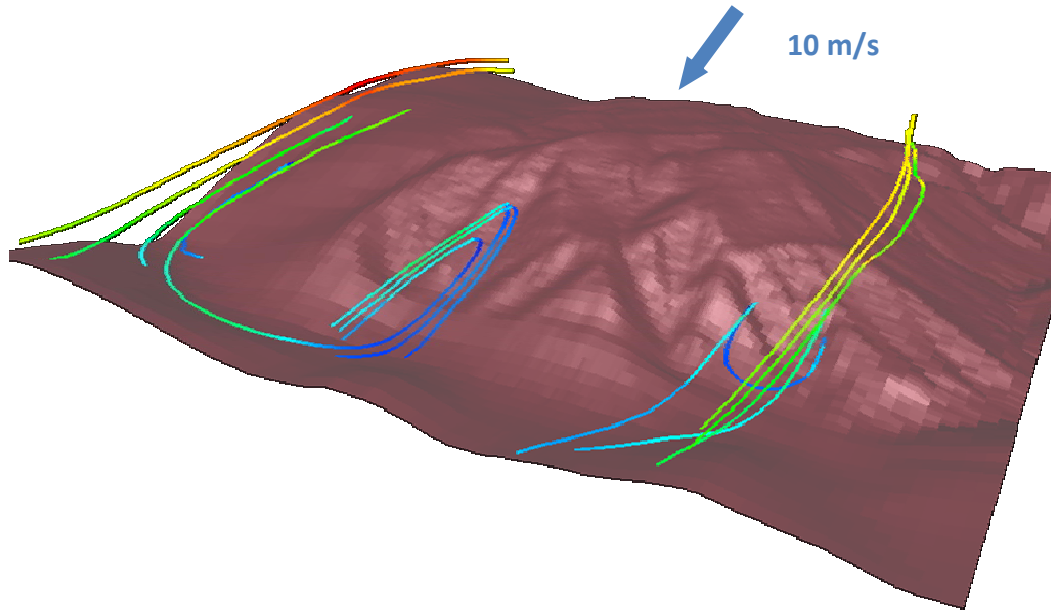


Figure 5.7: Particle tracing of the flow over the microscale model.

5.4 Microscale model – final speed-up maps

The third step of the free stream velocity test have been run in the same microscale model used to highlight the recirculation region, by changing the grid configuration in order to have a finer grid.

An inner layer has been built all over the terrain for 25 meters height, in order to fix the height of the cells attached to the ground. A sketch of the numerical grid is presented in the figure 5.8.

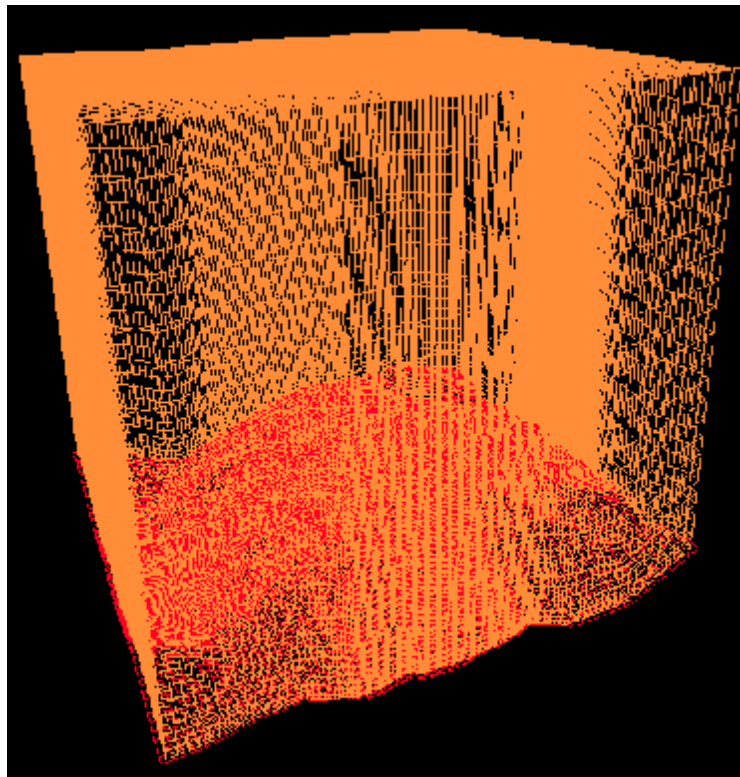


Figure 5.8: Numerical grid of the microscale model.

The test has been run by varying the wind speed, respectively for values equal to 5 – 10 – 20 m/s, and the wind direction, from 0 – 90 – 180 and 270 degrees.

In order to obtain the convergence of the solution, the CONWIZ function of Phoenix [5] has been used; this function helps the solver in improving the stability of the numerical method, but can also slow down the convergence rate.

The orthogonalization of the grid has also been tested only in the case of wind speed equal to 10 m/s; creating less skewed cells in the mesh improves the convergence of the iterative process.

The other geometrical and numerical settings are displayed in the following tables.

Geometrical model

Table 5.5: Input parameters for geometrical model in the microscale model.

INPUT PARAMETER	VALUE
Maximum number of cells	800000
Real number of cells	795328
Resolution in X-direction	25,6 m
Resolution in Y-direction	25,6 m
Number of cells in X-direction	136
Number of cells in Y-direction	136
Number of cells in Z-direction	43
Height distribution factor of Inner Layer	0,6
Height distribution factor Outer Layer	0,02
Height of the Inner Layer	25 m
Total Height of the model	5000 m
Minimum height of the first cell	2,6 m
Maximum height of the first cell	2,6 m

Numerical model

Tabella 5.6: Settings for the numerical model.

BOUNDARY CONDITIONS	
Top	Wall with no friction
Geostrophic height	From linked model (500 m)
Free stream wind speed	From linked model (variable)
Wind direction	0 - 90 - 180 - 270
Turbulence model	standard k- ϵ
SOLVER SETTINGS	
Solver	Coupled
Number of iterations	500 - 800 - 400

The high number of cells built up in the microscale model required some time to arrive to a good convergence level, which has been kept constant in all the simulations. Anyhow, few critical runs have been found, mostly in the simulation for 10 m/s: in the speed-up maps some areas in which divergence is growing up are well evident; this tendency to diverge has been reduced by increasing the free stream velocity, therefore the tests with 20 m/s are the ones that showed less problems of convergence. This fact let to think that in such a complex terrain and in such fine grid resolution the default value of the free stream velocity is not good enough. A presentation of this phenomenon can be seen in the figure 5.9, where in the left zone of the picture is shown a critical area; the problems are mostly relevant when wind is blowing from East sector, confirming the hypothesis by which the small variation in speed-up maps can be strictly connected to flow separation effects.

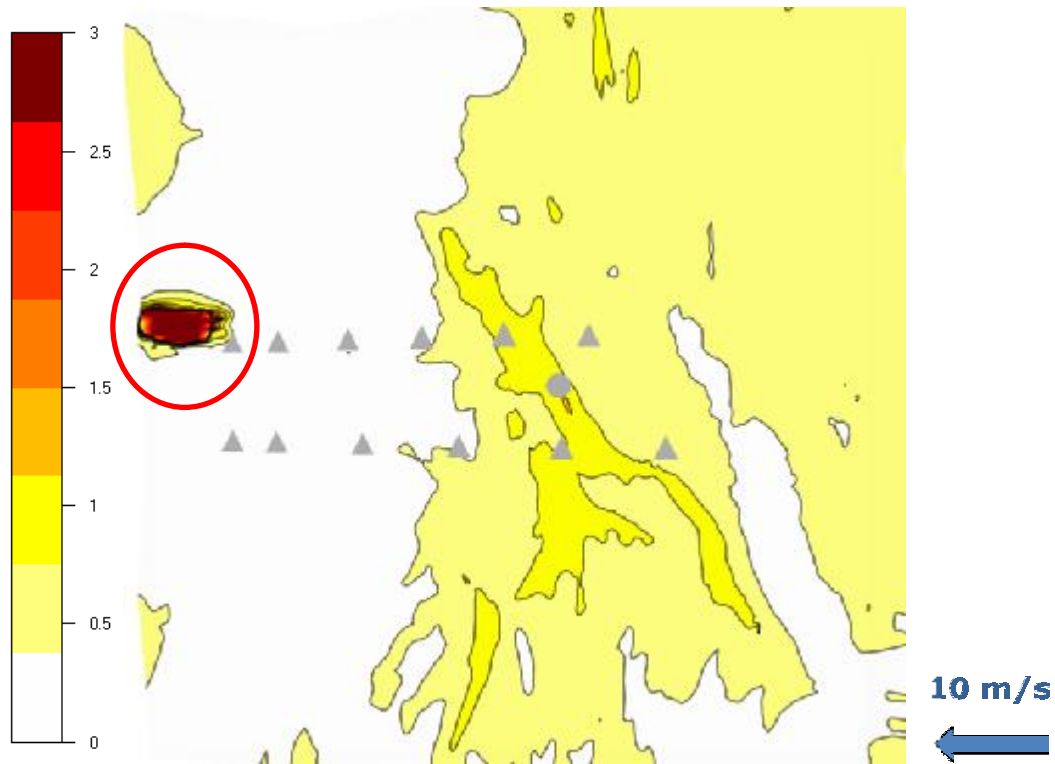


Figure 5.9: Map of speed-ups for directional sector 90, Speed 10 m/s, Height 30 m.

5.5 Remarks

The following pages show the horizontal speed maps for the microscale model, extracted at 30 meters height above ground level for a wind speed of 5, 10 and 20 m/s.

As it can be easily seen, the velocity of the free stream seems to be not very important in determining the speed-up maps. In fact, the pictures show good agreement and the variations are very small. The variations are smaller in the center of the maps, where the plateau is sited and the two mast measurements are placed. The value of 20 m/s for the free stream velocity seems to be the best value tested in order to reduce this variations and convergence problems; so this, value has been chosen to run the final simulations over the site.

To improve the results of this test, some other studies can be done: first of all a difference in speed-up value between various maps can be calculated, in order to have a concrete value to read and understand; then, a further turbulence model like the $k-\varepsilon$ RNG can be tested and a different solver, the segregated solver. Finally, also the diagonal wind directions should be tested, at least for the dominant sectors of the experimental data shown in chapter 4.

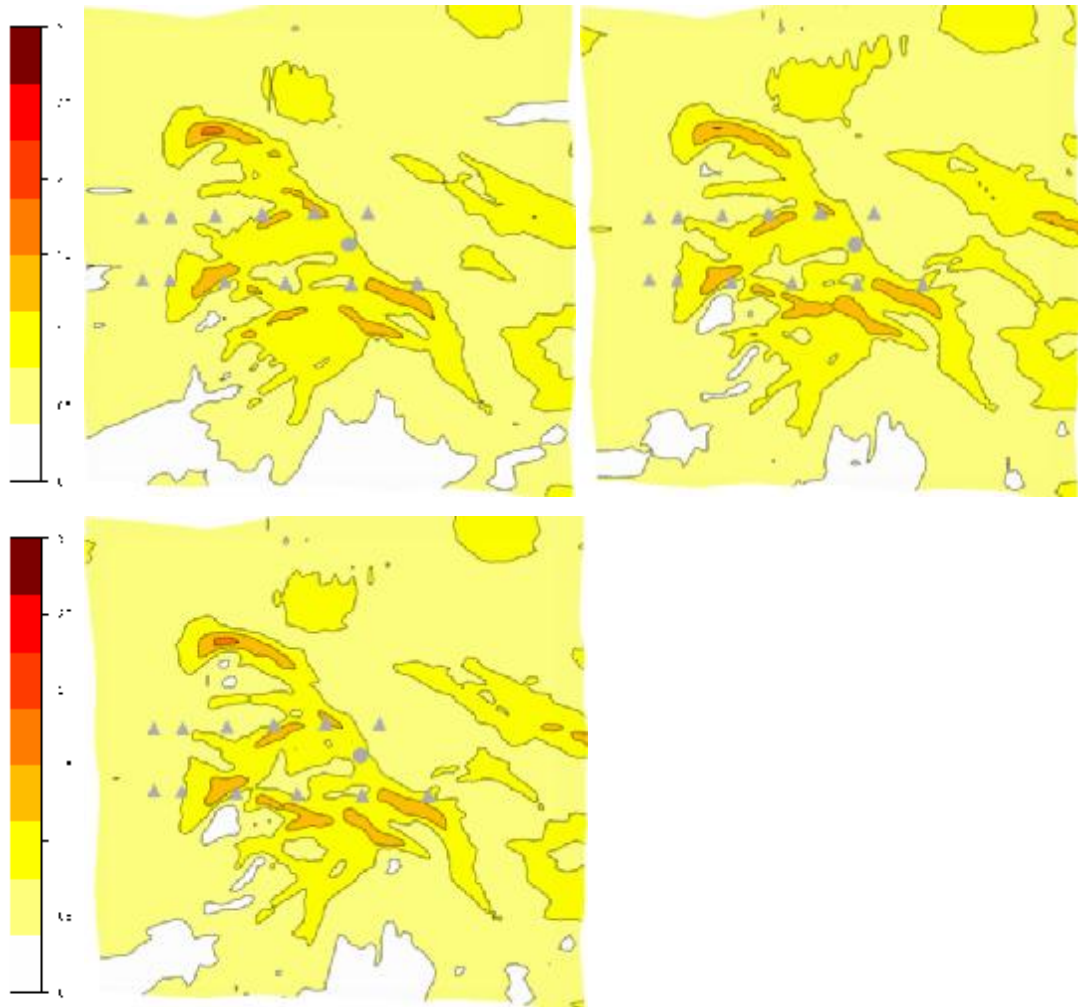


Figure 5.30: Speed-2D at 30 m height for sector 0 and velocity 5, 10, 20 m/s.



Figure 5.11: Speed-2D at 30 m height for sector 90 and velocity 5, 10, 20 m/s.

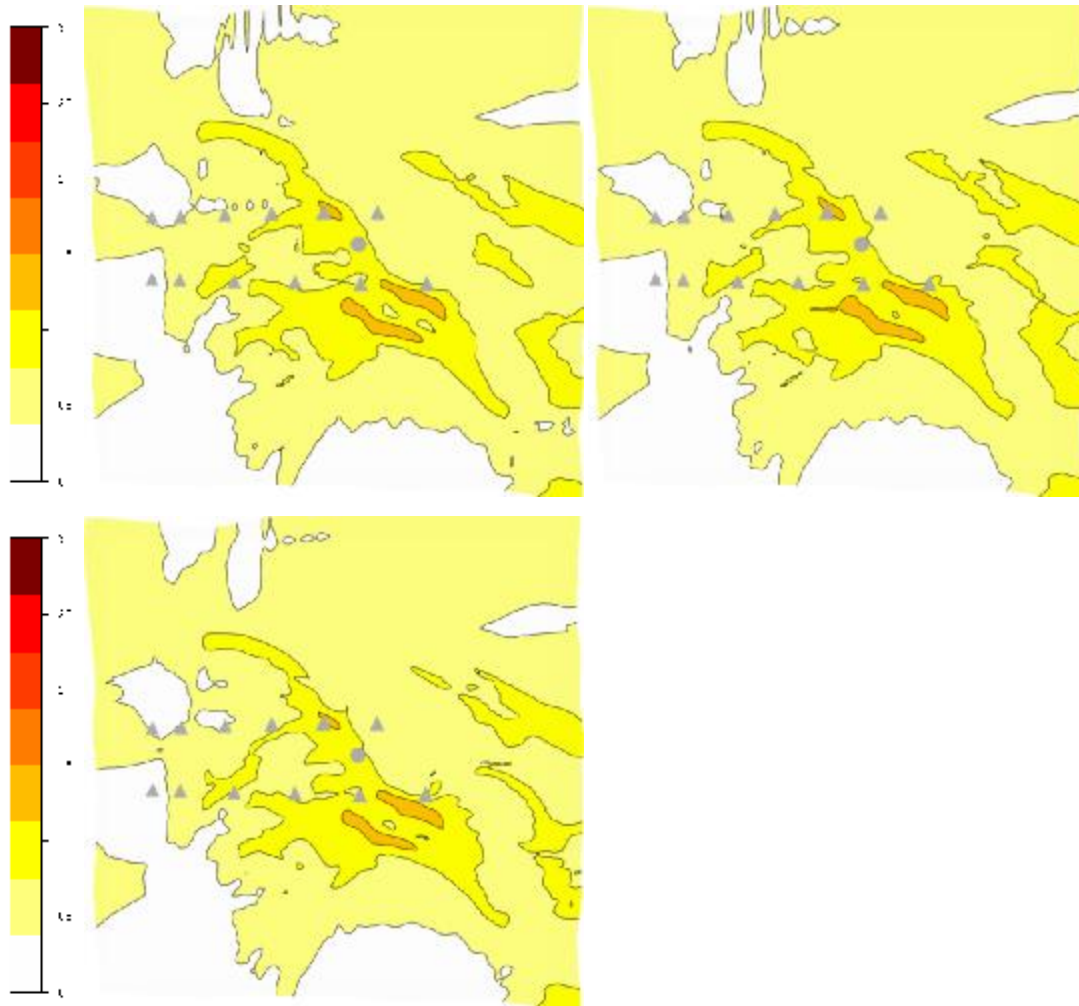


Figure 5.12: Speed-2D at 30 m height for sector 180 and velocity 5, 10, 20 m/s.

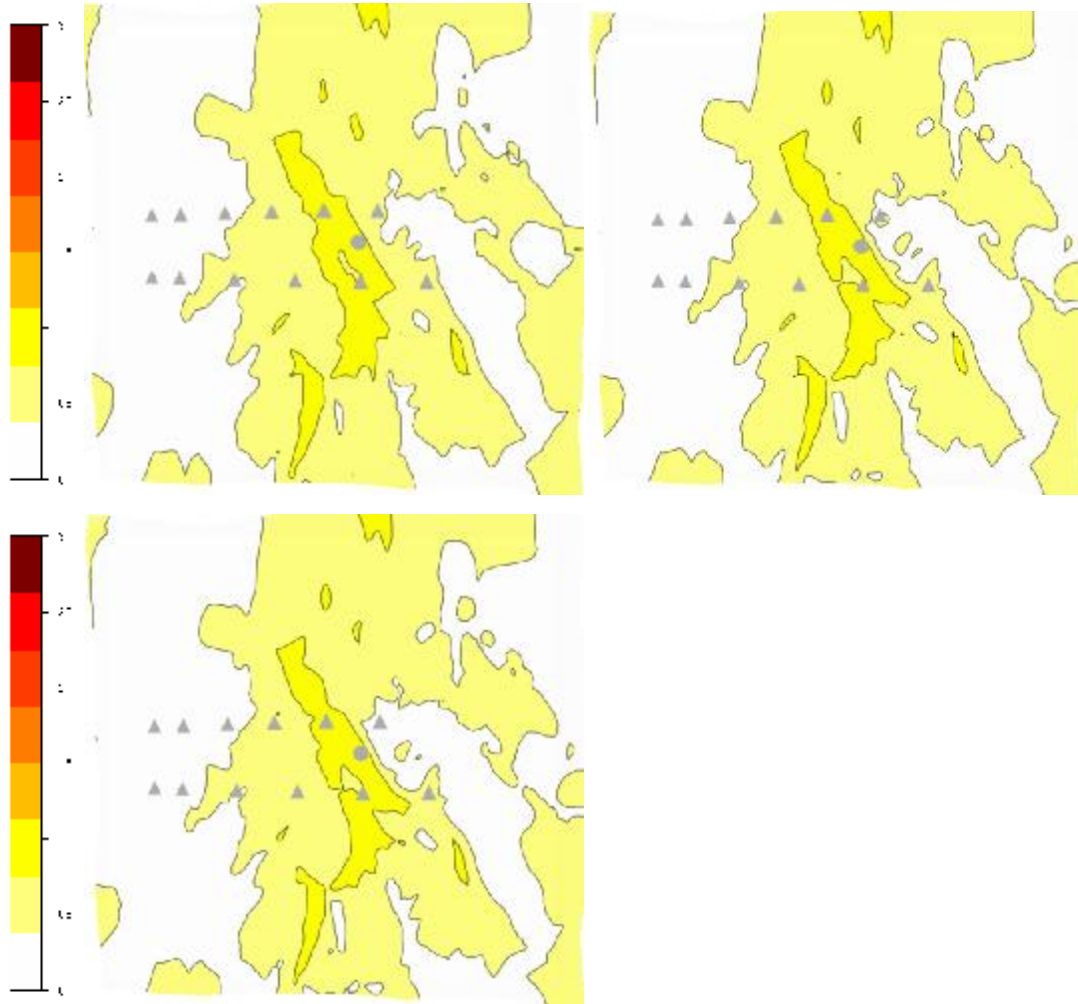


Figure 5.13: Speed-2D at 30 m height for sector 270 and velocity 5, 10, 20 m/s.

CHAPTER 6. Analysis of the wind field over the site.

6.1 Introduction

In this chapter a series of CFD simulations of the wind flow over a complex site in central Italy has been carried out. Different boundary condition, corresponding to different direction of the wind rose have been employed. The results of this CFD analysis are used to perform a cross checking procedure of the experimental wind data.

The geometrical model used in the simulations encloses the entire site and was resolved with a grid able to guarantee accurate results, as shown in chapter 5.

The cross checking procedure has been conducted by using the transferred climatology object, included in the WindSim software. As described in chapter 1, when multiple measurement masts are available in one site, it is possible to predict the wind resources in one measurement station using data available in the other measurement station. The comparison between experimental and CFD data can be done by producing graphs and calculating RMS errors to evaluate the accuracy of the results. Velocity vertical profiles and turbulence intensity vertical profiles have been also calculated and analyzed.

6.2 Geometrical model

The geometrical domain is based on a unique model to solve the Wind Field. A refinement of the major interest zone has been created, containing the plateau where the two measurement masts are located.

The characteristics of the geometrical model are shown in the table 6.1.

Table 6.1: Geometrical model characteristics.

Geometrical model	
Total number of cells	3136000
Extension	<i>Km 14,2 x 14,2</i>
x_resolution	<i>min 134 m – max 30 m</i>
y_resolution	<i>min 134 m – max 30 m</i>
Number of cells in x direction	280
Number of cells in y direction	280
Number of cells in z direction	40
Height Distribution Factor	0,01
Height of the model	5000 m
Minimum height of the first cell	2,4 m
Maximum height of the first cell	3,2 m
Number of cells in first 100 m height	6

6.3 Numerical Model

The 3D Wind Field of the site has been calculated through the WindSim software. This program is a powerful tool which solves the Reynolds Averaged Navier Stokes (RANS) equations for turbulent flows. The numerical model contains the continuity equation and the three momentum equations. The energy equation has not been solved since thermal effects have not been considered in the simulations. The closure problem of turbulent flows has been solved by introducing the standard κ - ϵ turbulence model. The boundary conditions and the settings for the solver are listed in table 6.2

Table 6.2: Boundary conditions and solver settings

Boundary conditions	
Top	wall with no friction
Geostrophic height	500 m
Free stream wind speed	20 m/s
Wind directions	12 sectors
Turbulence closure	
Turbulence model	standard k- ϵ
YAP correction	enabled
Solver settings	
Solver type	Segregated
Number of iteration	600
Convergence Wizard	enabled

6.4 Simulations

The total time needed to carry out the simulations was about 12 full days, one day per sector. The computer used for the simulations was a Dell Precision T5400, equipped by a quad-core CPU Intel Xeon with frequency 2,33 GHz each and 8 GBytes of RAM.

The residual values showing the obtained convergence level are shown in figure 6.1, for the simulations corresponding to the different directional sectors. A correct convergence level has been obtained in each sector.

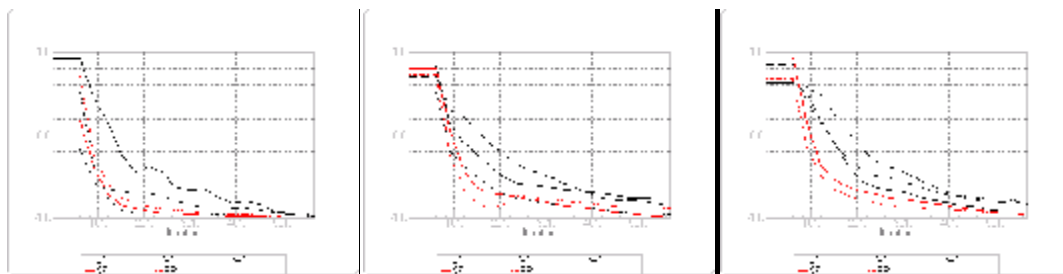


Figure 6.1 a/b/c: Residual values for sectors 0, 30 and 60 degrees.

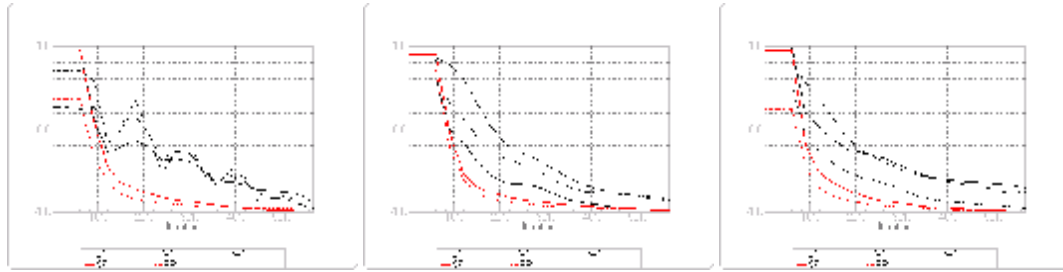


Figure 6.1 d/e/f: Residual values for sectors 90, 120 and 150 degrees.

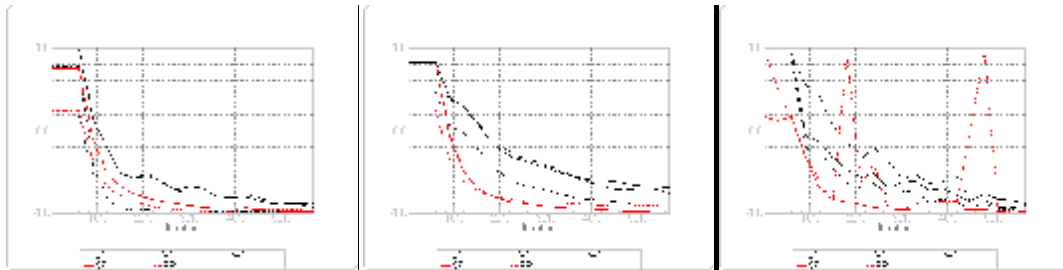


Figure 6.1 g/h/i: Residual values for sectors 180, 210 and 240 degrees.

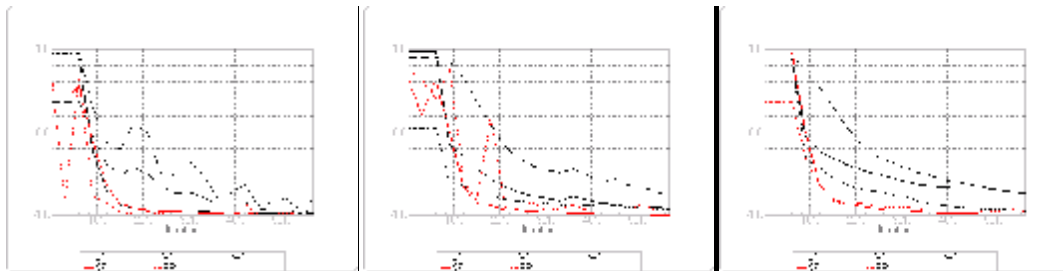


Figure 6.1 l/m/n: Residual values for sectors 270, 300 and 330 degrees.

According to the wind data analysis carried out in chapter 4, the dominant wind direction sectors have been highlighted. Moreover, some of the 12 sectors were characterized by a very low value of frequency. A threshold of 1,5 % was found to select the most frequent sectors, listed below

- 0°, 30°, 60°, 180°, 210°, 240° for tower 5424
- 0°, 30°, 60°, 90°, 210°, 240°, 270° for tower 5428

In the following pages the results corresponding to these sector, have been reported.

6.5 Vertical profiles of Velocity

Tower 5424 (CLIFF)

The sector chosen after analysis by frequency distribution of the tower 5424, placed near the cliff, are 0° , 30° , 60° , 180° , 210° , 240° and the vertical profiles of mean wind velocity relative to this sector are reported.

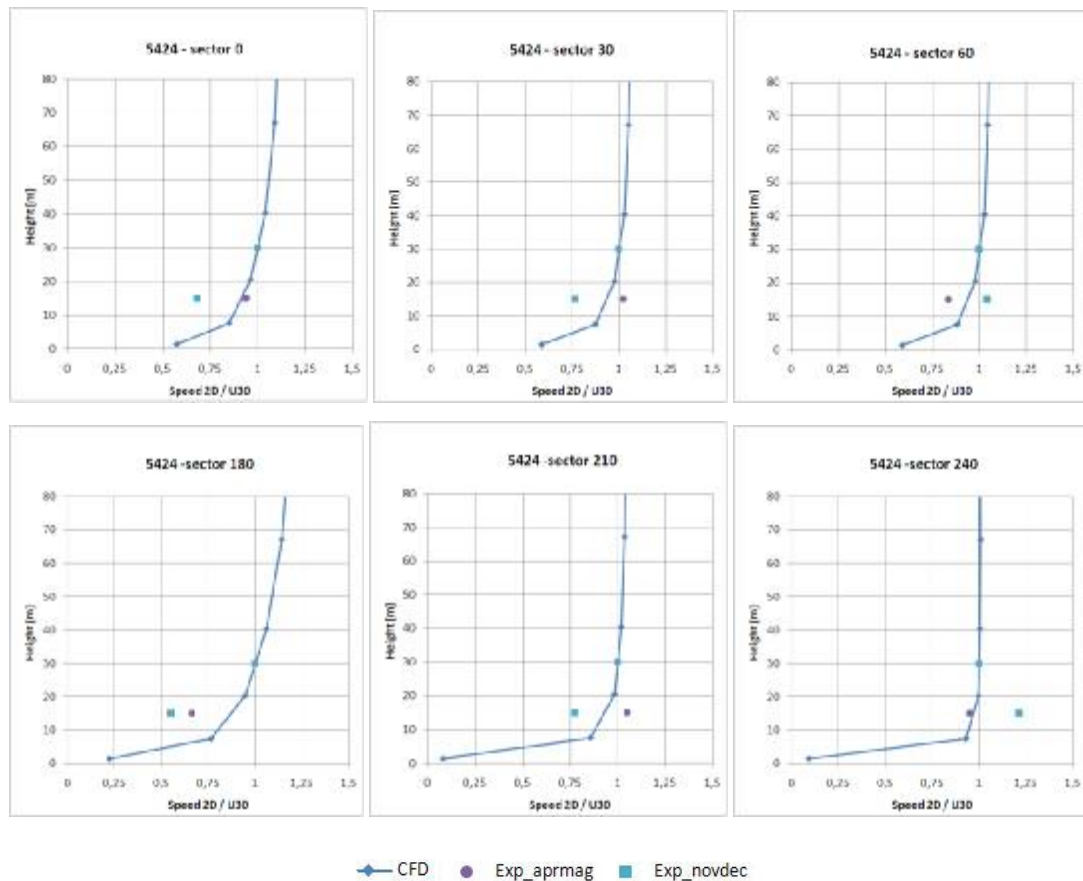


Figure 6.2 a/b/c/d/e/f: Velocity vertical profiles for tower 5424.

The values of the velocity U_{30} (experimental wind velocity at 30 meters height) have been used to calculate data in a non dimensional form, in order to make a comparison with CFD vertical profiles, also scaled with the corresponding velocity

value at 30 meters. It can be seen that every vertical profile passes through the point with coordinates (1, 30) of the graphs.

Values of U30 are reported in table 6.3, while in table 6.4 the RMS errors, calculated to evaluate the disagreement of the CFD results respect to the experimental wind data, are shown. In each table highest and lowest values are highlighted.

Table 6.3: U30 values in [m/s].

U30 - Tower 5424						
	0	30	60	180	210	240
AprMay	3,39	5,12	5,12	1,73	4,63	5,93
NovDec	3,35	5,47	6,01	2,61	4,59	4,24

Table 6.4: RMS Errors between experimental and numerical data

RMS errors - Tower 5424						
	0	30	60	180	210	240
AprMay	0,0169	0,0635	0,0732	0,1447	0,0843	0,0134
NovDec	0,1630	0,1184	0,0736	0,2246	0,1095	0,1725

The table of RMS errors has been calculated by the formulas

$$6.1) \quad E_{RMS} = \sqrt{\frac{1}{n}(e_1^2 + e_2^2 + \dots + e_n^2)} \quad \text{and}$$

$$6.2) \quad e_n^2 = (U_{EXP} - U_{CFD})_n^2$$

If the error results of the tower 5424 (CLIFF) are analyzed, worst and best sectors are deduced. The worst sector is 180°, showing highest errors in both periods April-May and November-December. The best sectors are, instead, 240° during the period April-May and 60°, during the period November-December.

It must be noticed the trend similarity between the errors and the values of the wind velocity at 30 meters height. In fact, the best sectors (240° and 60°) are

characterized by high wind speed (5,93 and 6,01 m/s), while the worst sectors are characterized by low speed (1,73 and 2,61 m/s).

A physical explanation of this phenomenon could be related to the seasonal thermal effects. It is wise to think that modification of the neutral boundary layer by thermal effects is high when the wind speed is low and vice versa. According to this, sectors characterized by high wind speed should be considered almost free of thermal effects. Then, considering that the simulations do not take into account the influence of thermal effects, the CFD results should have a better agreement with measurements in sectors where wind speed is high.

A threshold value related to the importance of thermal effects on the wind field wind velocity is difficult to be deduced from the results, but a range in which it is expected is 3-4 m/s. Data Analysis confirms that RMS errors are low when the wind is blowing with a velocity higher than a value contained in the range, resulting in a fair agreement between experimental data and result of the CFD model.

Tower 5428 (PLATEAU)

For the tower 5428, placed in the plateau, the sector chosen after the analysis for frequency distribution are 0°, 30°, 60°, 90°, 210°, 140°, 270°. In the figures, vertical distribution of mean wind speed are reported for these directions

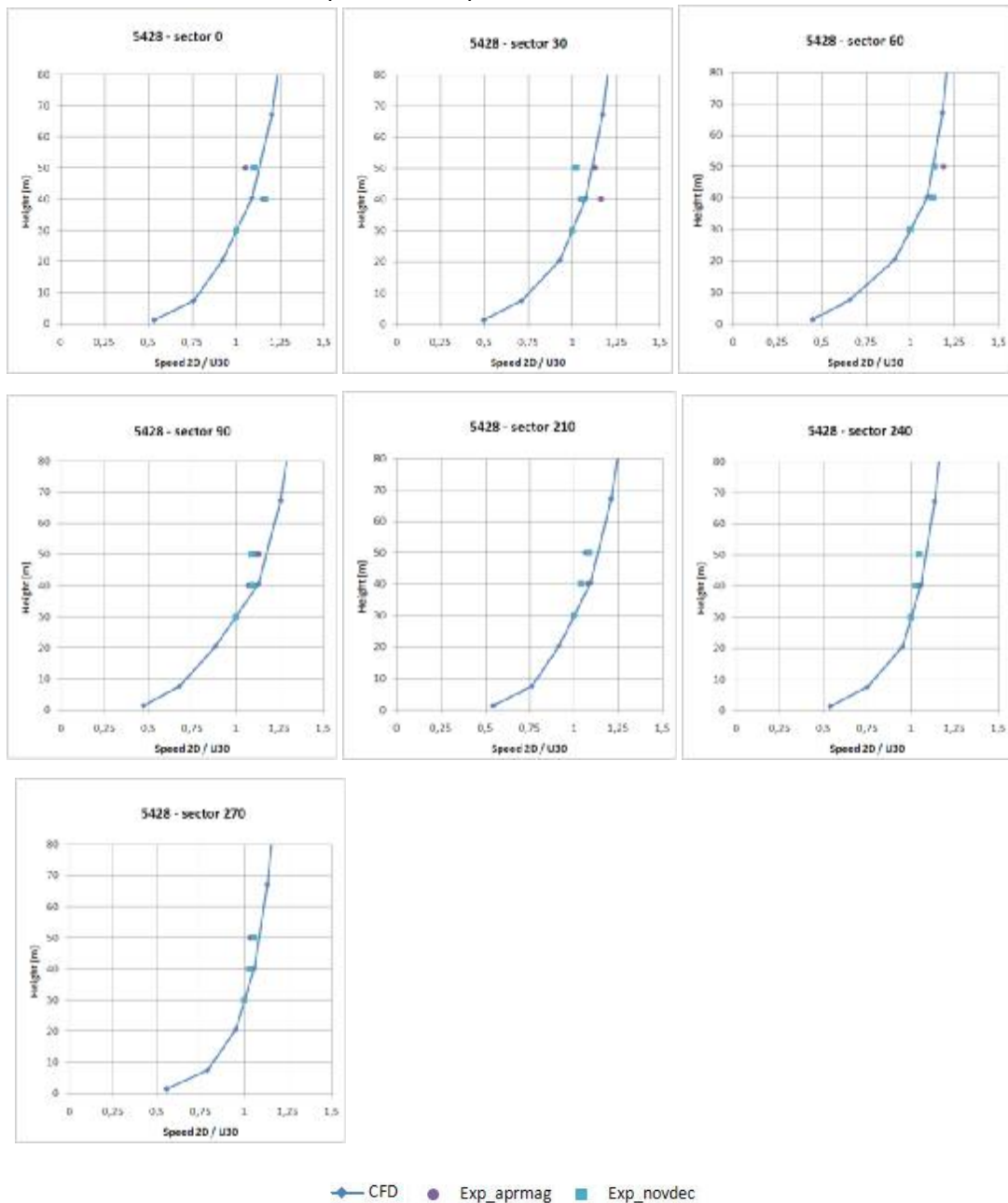


Figure 6.3 a/b/c/d/e/f/g: Velocity vertical profiles for tower 5428.

From the error analysis, it is found that during the period April-May the best sector is 240° and the correspondent U30 is one of the highest; the worst sector is 0 and the U30 has the lowest value (see table 6.5 and 6.6). An analog behavior is found for data in the period November-December, when the best sector is 270° and the worst is 30°.

Table 6.5: U30 in [m/s].

		U30 - Tower 5428						
		0	30	60	90	210	240	270
AprMay		2,24	2,24	3,26	3,33	2,46	3,79	5,52
NovDec		2,18	2,34	4,13	3,45	3,62	4,2	4,17

Table 6.6: RMS errors.

		RMS errors - Tower 5428						
		0	30	60	90	210	240	270
AprMay		0,0599	0,0529	0,0375	0,0383	0,0378	0,0223	0,0302
NovDec		0,0491	0,0541	0,0219	0,0493	0,0409	0,0265	0,0186

The similarity between RMS errors and wind velocity trends of the tower 5424 (CLIFF) in the sector analyzed, is also confirmed in the tower 5428 (PLATEAU). The threshold value of the wind speed can still be identified in a range 3-4 m/s, which seem to be a correct value to distinguish between low and high speeds.

According to the analysis carried on the profiles of vertical velocity, it can be stated, as a general rule, that CFD can well reproduce experimental data when the thermal effects are reduced, that means when the wind speed at 30 meters in the sector is higher than 3-4 m/s. The choice of a single threshold value was not possible due to the different behavior of the flow in the two measurement masts and in the two different periods.

The characteristics of the relation between wind speed U_{30} and RMS errors on mean wind speed are summarized in figure 6.4. It can be stated that errors calculated for tower 5428 are lower than those calculated for tower 5424. The fact could be related to the height of the three measurement stations of the tower 5428 (30-40-50 m) respect to the stations of tower 5424 (15-30m). Measurements close to the ground are difficult to predict due to the complex orographic effects.

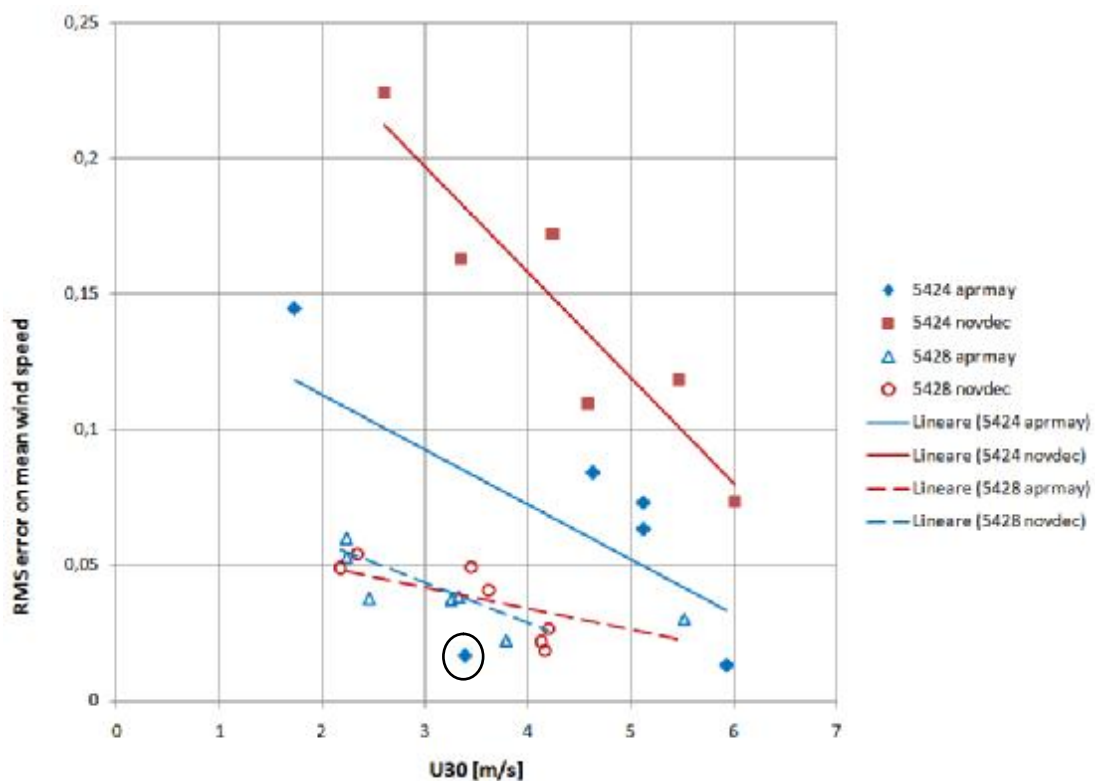


Figure 6.4: Relation between RMS errors and sector wind speed.

Moreover, a general trend for RMS errors can be underlined in the graph, showing that errors become lower when wind speed U_{30} is becoming higher. Again, this fact is related to the importance of thermal effects, that decreases in presence of a high wind speed.

Anyhow, the trend of the linear interpolation is different according to the measurement period. Thermal effects are different along the year, due to solar radiation and thermal inertia of the terrain, inducing different wind breezes in the all area.

Finally, it must be noticed that in the tower 5424 for the period November-December there is the best alignment of the data, while in the period April-May the marker circled in black shows a non correct behavior compared with the other data and should be excluded by the analysis. The point corresponds to sector 0° , which has a low frequency during the considered measurement period and could be influenced by errors in the statistic analysis. Concerning the tower 5428, data are well disposed in a trend in both periods.

6.6 Turbulence Intensity Vertical Profiles

In this paragraph vertical profiles of turbulence intensity (T.I.) are reported and analyzed. The figures 6.5 and 6.6 show the comparison between the measured turbulence intensity values and the results of CFD simulation.

The experimental T.I. values have been calculated by the time-history of wind data as a ratio between the standard deviation of wind velocity and the mean wind velocity evaluated for the instantaneous measurements taken in 10 minutes

$$6.3.3) \quad T.I. = \frac{\sigma}{U}$$

where

σ	is the standard deviation of wind velocity	[m/s]
U	is the mean velocity	[m/s]

By organizing experimental data for wind direction it is possible to group these angles in sectors (in the specific case one sector every 30° for a total of 12 sectors) and to consider the wind velocity and the associated T.I. value. Then, it is possible to create twelve scatter graph, where T.I. is a function of U. In general, these graphs show a cloud of points that can be approximate by an hyperbolic function. Data reaching an infinite value of T.I. for low wind velocity are present while the expected trend goes toward zero when wind velocity grows. The common practice eliminates the high values of turbulent intensity associated to low wind speed, that are probably affected by thermal effects, when an interpolation of the cloud of points is required. According to this practice rule the TI was evaluated rejecting the values connected to low mean wind velocity. The threshold value was chosen following the conclusions of the previous paragraph 6.5, where a threshold range was indicated in $U_{30} = 3\div 4 \text{ m/s}$. In these particular analysis the value of $U=4 \text{ m/s}$ has been used to filter experimental data.

The CFD simulations did not take into account thermal effects, thereafter under estimating the turbulence intensity data. A procedure of rejecting these data is then needed to reduce the thermal effects in the experimental data used for comparison.

Concerning the CFD simulations, values of T.I. have been evaluated using the equation 1.12 mentioned in chapter 1.

The T.I. graphs in the figures 6.5 and 6.6 are organized for each measurement period, measurement mast and wind sector. RMS errors are also calculated, in order to understand the accuracy of the results.

Tower 5424 (CLIFF) experimental data filtered at $U=4\text{ m/s}$

According to the analysis by temporal frequency conducted in chapter 4, the sectors considered for the turbulence intensity study are the same used for the velocity profiles analysis. Sector which frequency was below 1,5 % have been excluded and the considered ones are $0^\circ, 30^\circ, 60^\circ, 180^\circ, 210^\circ, 240^\circ$.

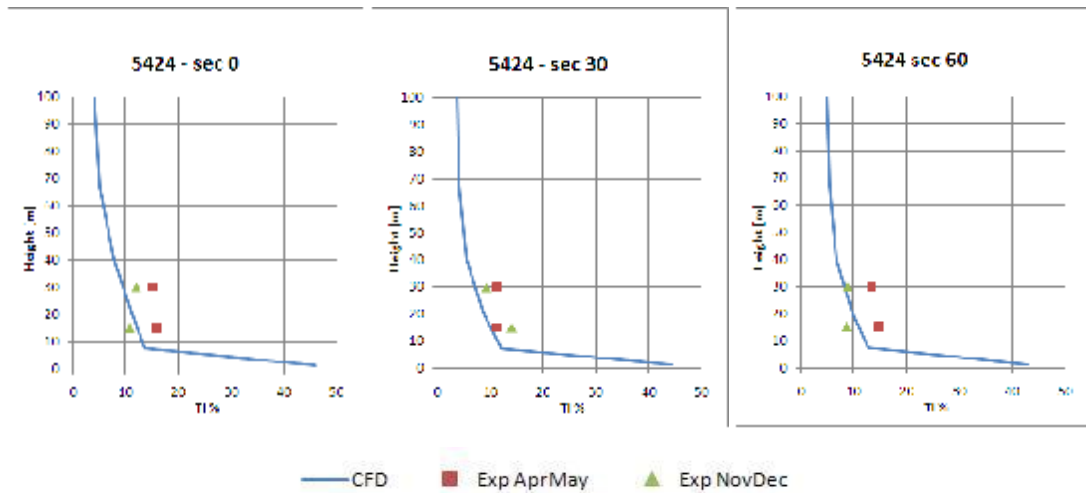


Figure 6.5 a/b/c: Turbulence Intensity vertical profiles for tower 5424.

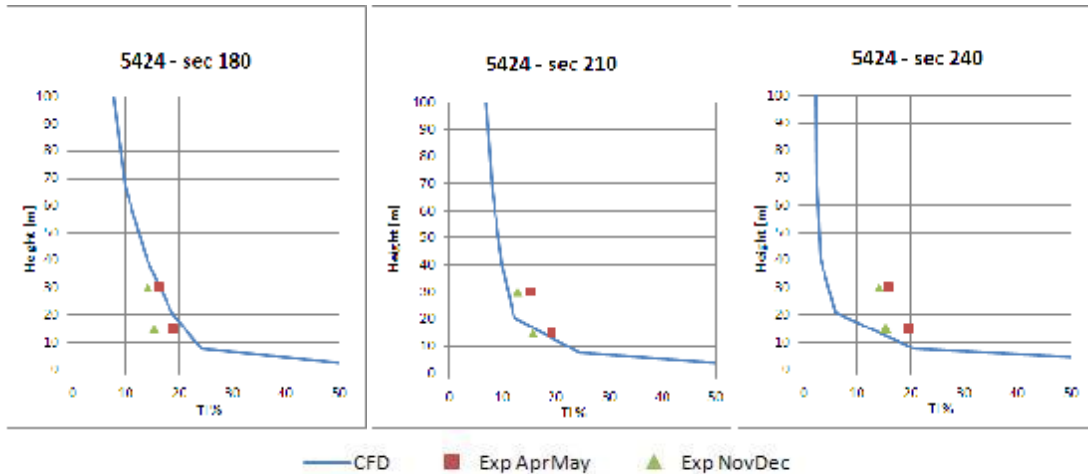


Figure 6.5 d/e/f: Turbulence Intensity vertical profiles for tower 5424.

By general overview of the turbulence intensity vertical profiles it seems that CFD well approximates the experimental data. In tables 6.7 and 6.8, RMS errors related to vertical profiles have been calculated and presented together with the U30 wind velocity.

Table 6.7: RMS errors for tower 5424.

T.I. RMS error - Tower 5424						
	0	30	60	180	210	240
Apr-May	4,87	3,07	4,64	1,56	3,27	9,56
Nov-Dec	2,09	3,19	1,69	4,31	1,73	7,18

Table 6.8: Values of U30 for tower 5424.

U30 - Tower 5424						
	0	30	60	180	210	240
AprMay	3,39	5,12	5,12	1,73	4,63	5,93
NovDec	3,35	5,47	6,01	2,61	4,59	4,24

Analyzing RMS error table for tower 5424, it is found that the sectors where turbulence intensity level is higher is 240° , during both periods April-May and November-December. Sectors where turbulence intensity is lower are sector 180° and 60° , respectively for periods April-May and November-December. By considering the RMS table and the U30 table, different behaviors for the two periods of reference can be found. In fact, during April-May the error seems to increase with the wind speed, while during November –December the error seems to decrease when U30 is growing. This behavior is well represented by the trend lines shown in figure 6.7.

In general, it can be stated that the values of the T.I. seem to be strongly linked to the seasonal thermal effects. As was already said before, the solar radiation and the thermal inertia of the terrain can generate different wind breezes in the measurement periods, also influencing the turbulence level. A stronger influence on the turbulence intensity level can be due also to the orographic effects. The tower 5424 is placed in a cliff, where vertical component of velocity are present and can be strongly variable in intensity during the whole year because of the different solar radiation. Particularly, strong solar radiation near a cliff can generate a intense vertical wind, where the fluctuations are stronger. This phenomenon could explain the trend relative to April-May period in figure 6.7, where in general the errors increase with the U30.

Finally, it must be said that turbulence is a parameter very difficult to measure, especially in atmospheric flows, and that a only a longer measurement campaign can guarantee more accurate results. Due to all this consideration, the agreement between experimental data and CFD results for mast 5424 was fairly good.

Tower 5428 (PLATEAU) experimental data filtered at $u=4$ m/s

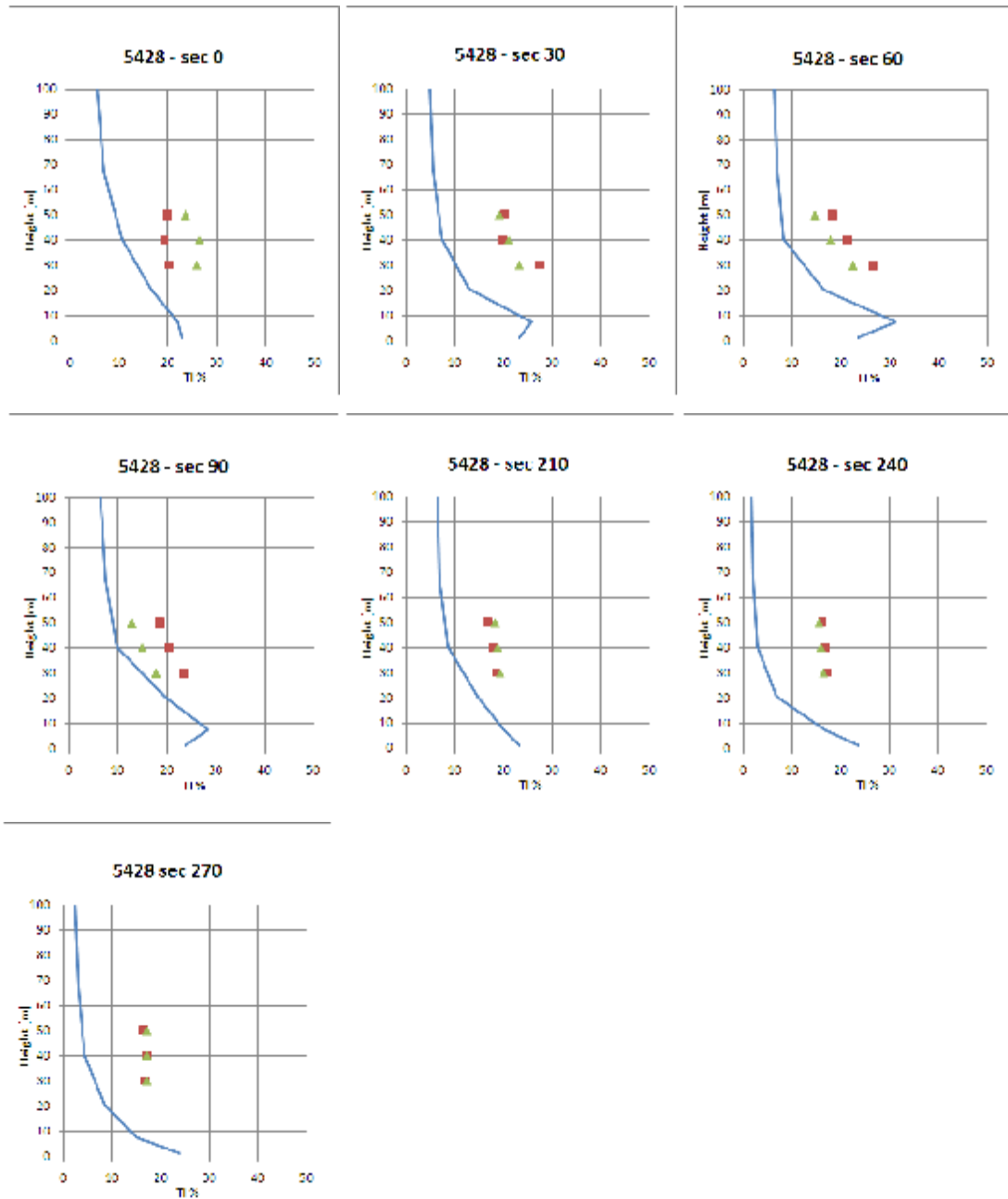


Figure 6.6 a/b/c/d/e/f/g: Turbulence Intensity vertical profiles for tower 5428.

— CFD ■ Exp AprMay ▲ Exp NovDec

A general overview of the figure 6.6 shows that CFD under estimated the experimental turbulent intensity level in the tower 5428. This fact could be related to a different influence of thermal effects on the experimental data respect to the mast 5424.

Table 6.9: RMS errors for tower 5428.

T.I. RMS error - Tower 5428							
	0	30	60	90	210	240	270
Apr-May	8,71	14,65	12,55	9,64	8,19	13,35	11,99
Nov-Dec	14,06	13,10	8,81	3,92	9,31	12,46	12,29

Table 6.10: Values of U30 for tower 5428.

U30 - Tower 5428							
	0	30	60	90	210	240	270
AprMay	2,24	2,24	3,26	3,33	2,46	3,79	5,52
NovDec	2,18	2,34	4,13	3,45	3,62	4,2	4,17

The analysis of the relation between RMS errors and U30 wind speed has been done also for mast 5428 and summarized in figure 6.7. It is confirmed the general trend by which during April-May the errors increase with the wind speed, while during November-December the error decreases while U30 wind speed is increasing. As for mast 5424, the dependence of the experimental data by seasonal thermal effects seems to be strong. Moreover, in this mast the errors are higher than the errors calculated for tower 5424 and it is evident the under estimation of the turbulence intensity level.

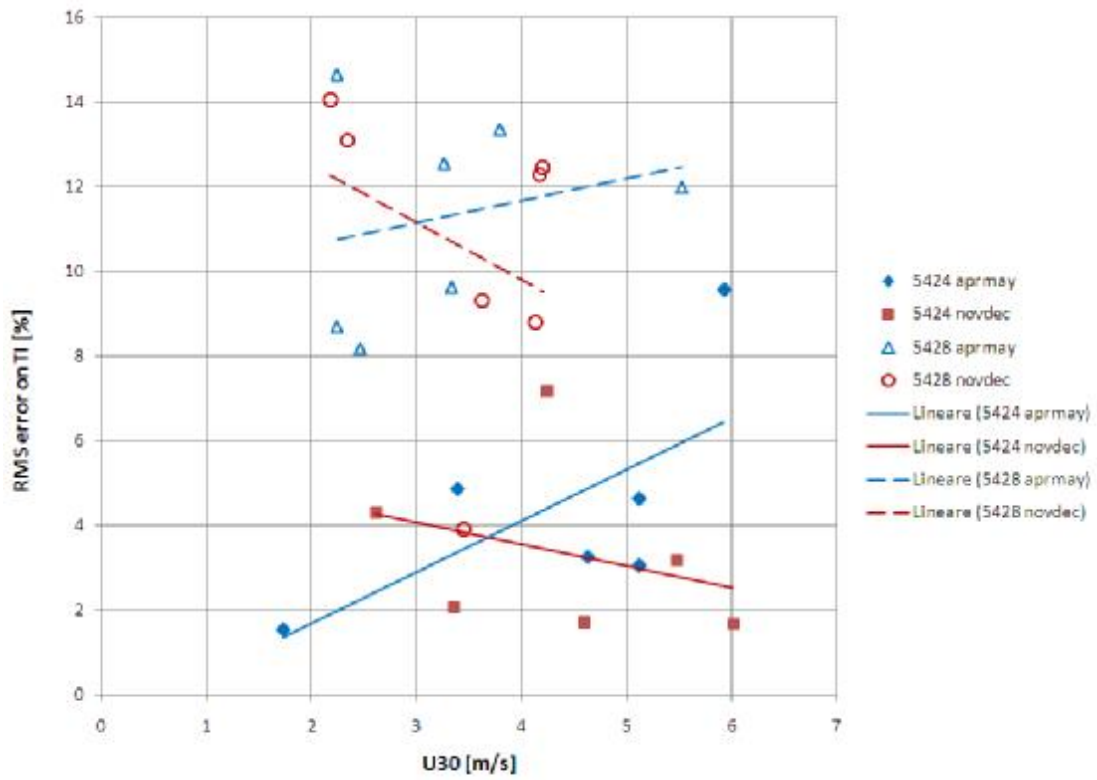


Figure 6.7: Relation between RMS errors and U30.

6.7 Cross Checking of wind data

In the following paragraphs the cross checking of the experimental data, as explained in chapter 1, is presented. The WindSim feature, called transferred climatology object, was used. As previously described when multiple measurement masts are available in one site, it is possible to predict by CFD calculation the wind resources in one measurement station using data available in the other measurement station. In other words, it's possible to transfer wind data from the original position to the other taking into account the different orographic and roughness effects

The cross checking has been studied using two different procedures, using as scaling variable the *mean wind velocity* per sector and the *speed-up* values, respectively.

In the first case the experimental mean velocity values have been extracted by the time history, while all the CFD values have been calculated using the sector interpolation function [4] of WindSim.

In the second case, relative to the speed-up calculation, two methods have been used, one considering the sector interpolation and the other one without this option. Concerning the experimental speed-up, this value was calculated directly from experimental dataset, by a ratio between the mean velocity in the reference mast and the mean velocity in the new position.

Sector interpolation

The sector interpolation function is a correction function of WindSim, needed when the CFD results must be weighted against one experimental data-set, in order to determine the horizontal wind speed map.

When the wind direction boundary condition is given, it is applied at the boundaries of the model, by describing a geostrophic wind. Nevertheless the wind evaluated in the internal zone can be deviated by orographic effects, especially close to the ground. For instance, if two simulations are run for sectors 0° and 30° as boundary conditions, in the center of the domain the wind direction could be, for instance, 2° and 35° respectively. In this case there will be a mismatch between the directions of

the wind rose and the directions of the CFD results. To solve this problems, the sector interpolation function is helpful, in order to modify the CFD results.

The wind directions found with the CFD must be referred at one direction of the wind rose and this is done by a linear interpolation procedure. For instance, the correct CFD velocity for the direction 30° is obtained weighting the two velocities calculated (2° and 35°) in which the 30° direction is comprised. The weight is proportional to the distance of the actual CFD velocity from the wind rose direction. Once this weighted average of sector 0 and 30 is done, the obtained dataset is used to scale the experimental climatology data for sector 30, allowing the preparation of the wind resources map. This procedure is repeated for all the directional sectors at the climatology.

6.7.1 Mean wind velocity

The concept of the accuracy of the CFD calculation ha also been applied in the cross checking procedure, and all the sector with a temporal frequency below 1,5 % have been excluded. Moreover, the conclusions valid for velocity vertical profiles have been also applied to the cross checking and the sectors in which the thermal effects are strongly reduced have been highlighted with a red circle, i.e. the sectors for which U30 is included in or greater than the range 3 ÷ 4 m/s.

In figures 6.8 the results of the cross-checking procedure by mean wind velocity are presented. For instance in figure 6.8a, for each direction sector, the velocity on mast 5424 in the period April-May, forecasted by data on mast 5428 at different height, are reported.

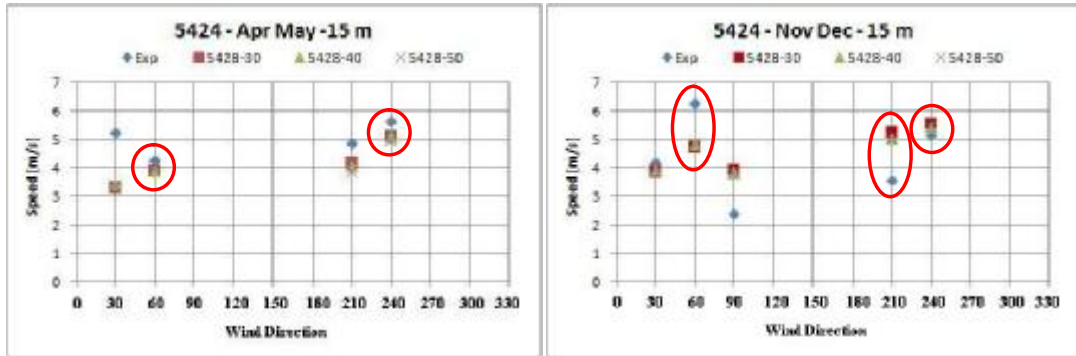


Figure 6.8 a/b

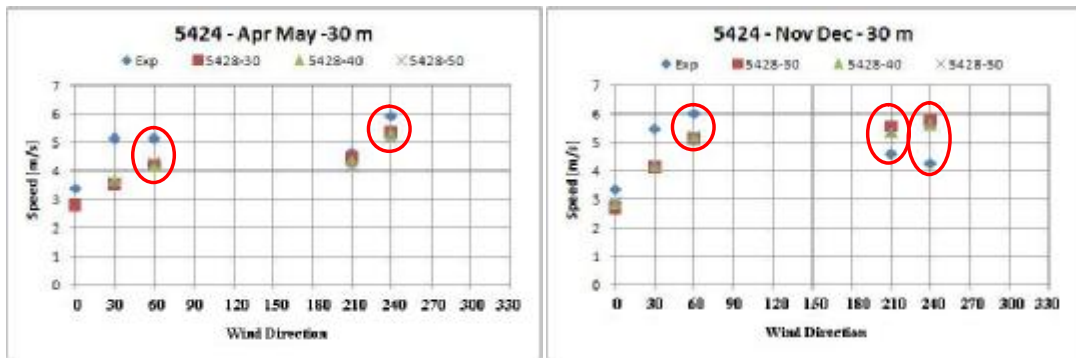


Figure 6.8 c/d

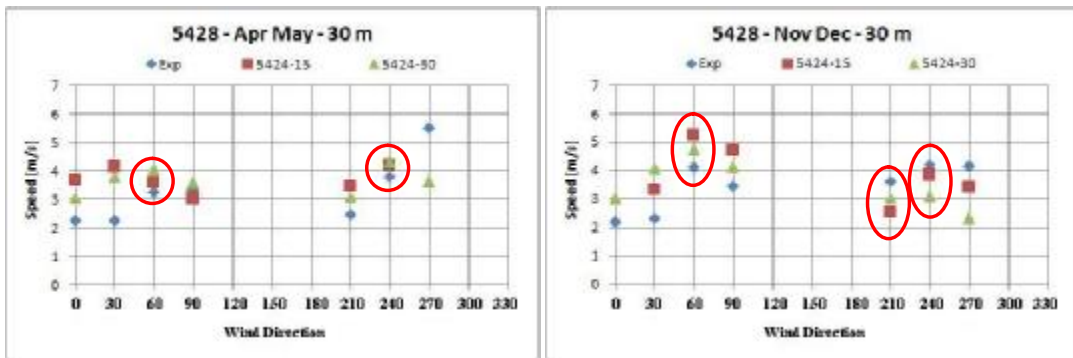


Figure 6.8 e/f

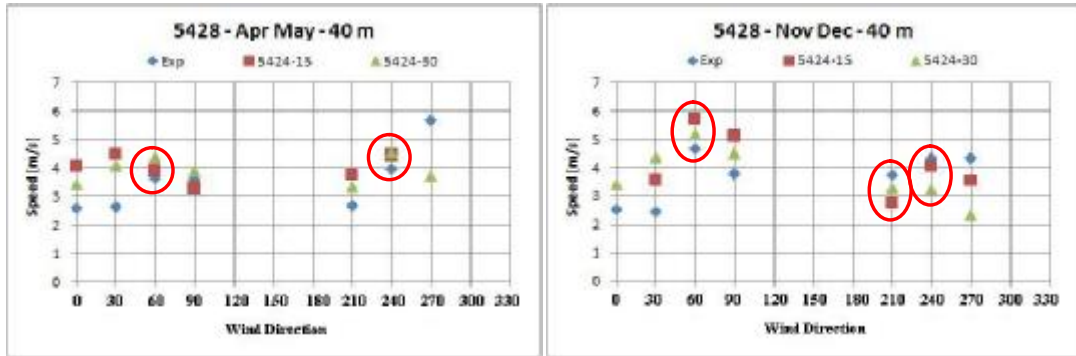


Figure 6.8 g/h

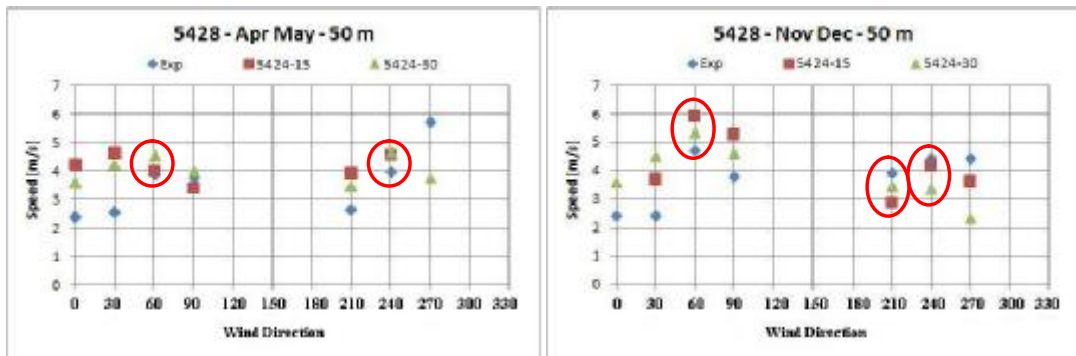


Figure 6.8 i/l

The graphs show a fair agreement between experiments and CFD in the sector where seasonal thermal effects are reduced, i.e. in sectors marked with red circles in the figures 6.8. Some disagreement is found in the period Nov-Dec at 15 m. In the other sectors the errors generally increase, showing again the influence of the mean velocity on the errors. The RMS errors, shown in table 6.8, have been calculated taking into account only the sectors in which the thermal independence is obtained, in order to establish which set of data can reproduce better the experimental data. The table should be read in this way: on the left vertical column, the mast and height for which experimental data are used as reference values in the cross checking are listed. On the horizontal row, the mast and height where the wind data are forecasted is arranged. Obviously, from 5428 mast only data on the 5424 mast can be deduced, and vice versa.

Table 6.11: RMS error for cross checking by mean wind velocity.

Period April-May							Period November-December						
NEW POSITION							NEW POSITION						
ORIGINAL POSITION	MEAN SPEED					Total	ORIGINAL POSITION	MEAN SPEED					Total
	5424 30	5424 15	5428 50	5428 40	5428 30			5424 30	5424 15	5428 50	5428 40	5428 30	
5424-30			0,690	0,676	0,661	0,676	5424-30			0,751	0,722	0,600	0,759
5424-15			0,446	0,396	0,389	0,410	5424-15			0,938	0,862	0,921	0,907
5428 50	0,811	0,547				0,671	5428 50	1,027	1,280				1,179
5428-10	0,798	0,469				0,631	5428-10	1,018	1,196				1,107
5428-30	0,795	0,455				0,625	5428-30	1,169	1,139				1,254

According to the table 6.11 the best match is found in the period April-May, starting from data of the station 5424 at 15 meters height and predicting the wind characteristics in station 5428 at 30 meters height, situation that corresponds to the graph 6.8 e.

Table 6.11 allows also to understand which dataset is to be used for a prediction of the wind in the other tower. The best dataset corresponds to station 5424 at 15 meters high, as shown by the total RMS error. Considering that the 5424 tower is positioned near a cliff this result could seem somehow peculiar. An explanation of this phenomenon can be related to the fact that is relatively straightforward to predict the flow at higher quotes and for non complex orography (mast 5428), even starting from data at low quotes and placed in critical positions as a cliff (tower 5424).

In general, from table 6.11 it can be seen that errors are lower in the period April-May and this could be explained by assuming that seasonal thermal effects are more influent when the terrain is cooling (Nov-Dec) instead of warming up (Apr-May).

6.7.2 Speed-up

In the figures 6.9 and 6.10 the results of the cross-checking by speed-up are reported; the label “*transf clima*” refers to the simulation with the sector interpolation option enabled, while the label “*wecs*” refers to the case with this option switched off. The green triangular marker is related to the experimental speed-up, calculated as the ratio between the two specified measurement stations.

For instance in figure 6.9a, for each direction sector, the velocity on mast 5424 at 15 meters in is forecasted by data on mast 5428 at 30 meters. The CFD calculations with sector interpolation and without this options are compared with the experimental speed-up.

The RMS errors are calculated only for sectors where thermal effects are reduced, marked with the red circles.

Period April-May 2007

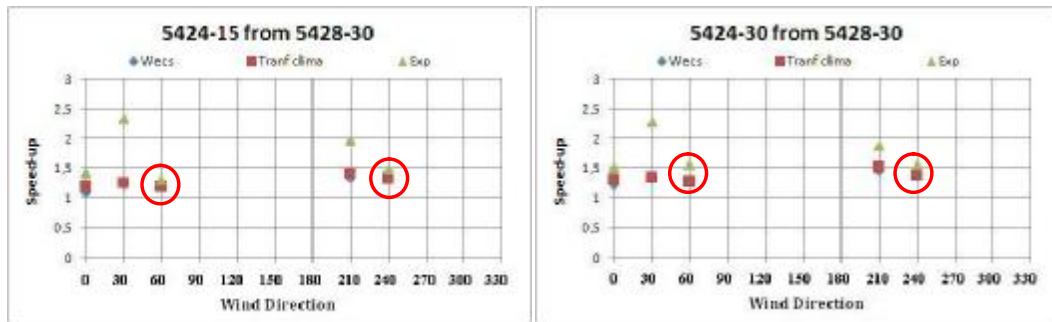


Figure 6.9 a/b

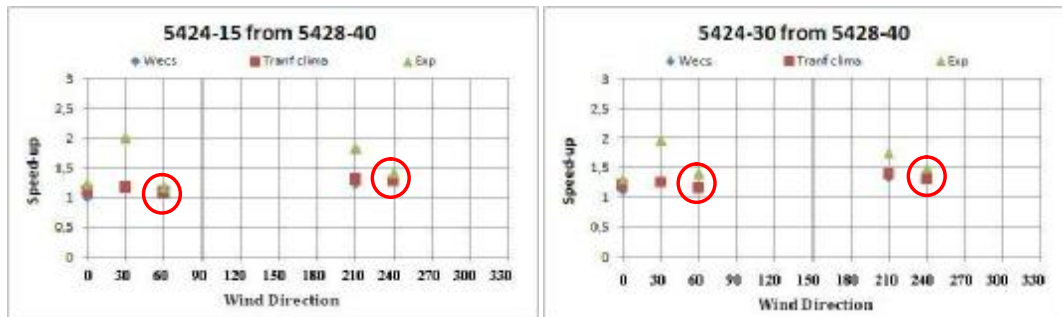


Figure 6.9 c/d

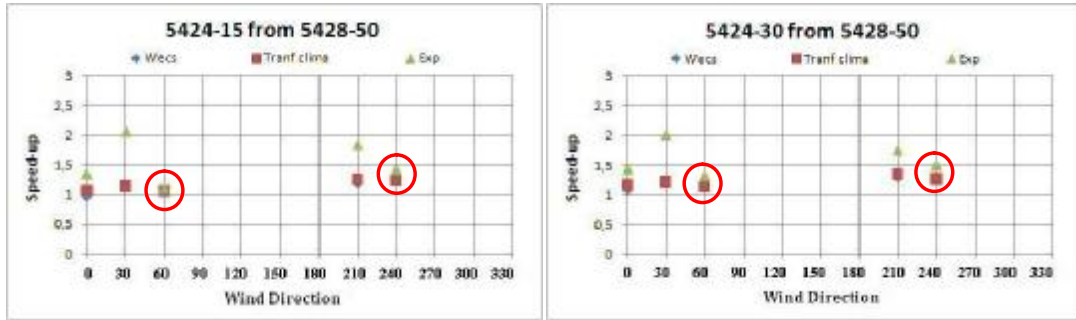


Figure 6.9 e/f

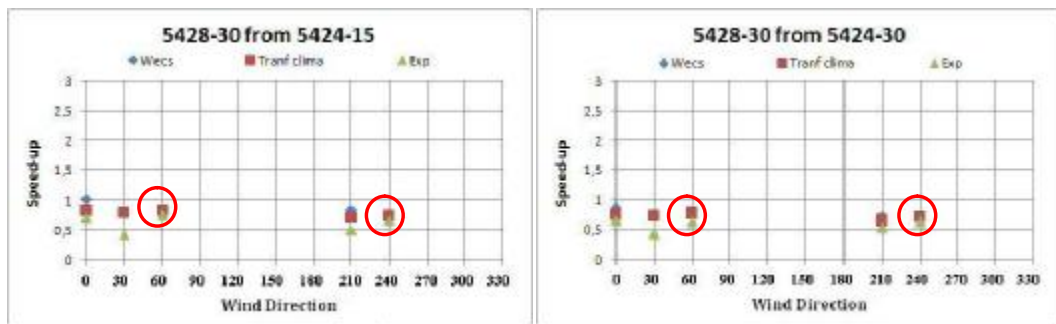


Figure 6.9 g/h

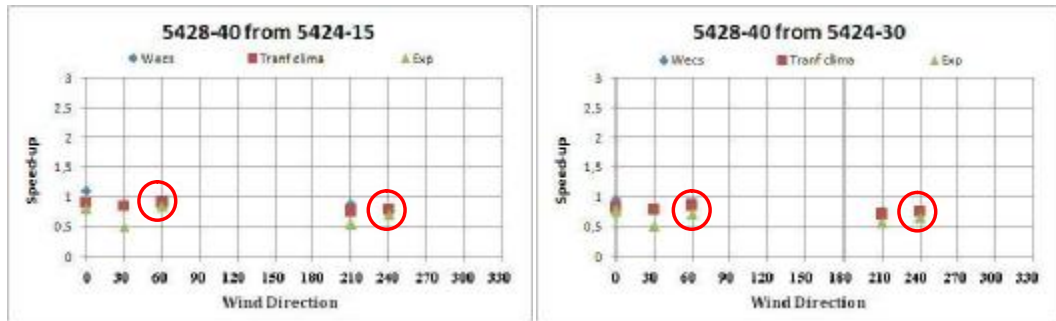


Figure 6.9 i/l

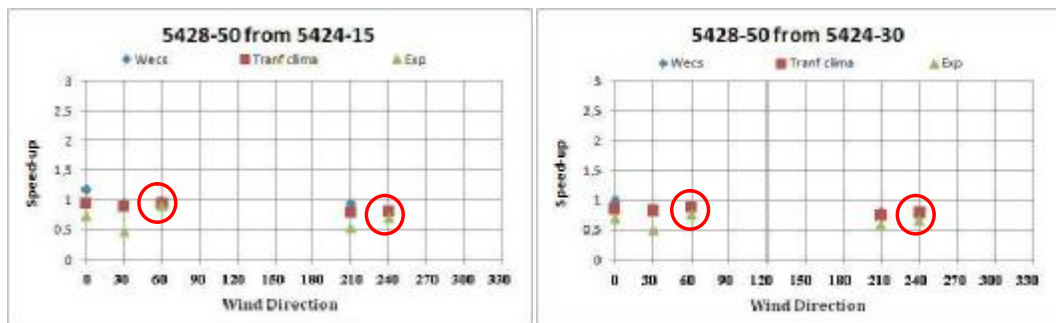


Figure 6.9 m/n

Period November December 2006

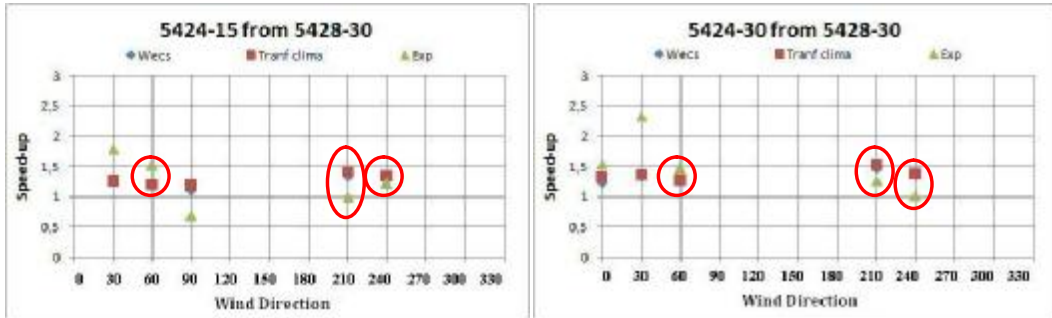


Figure 6.10 a/b

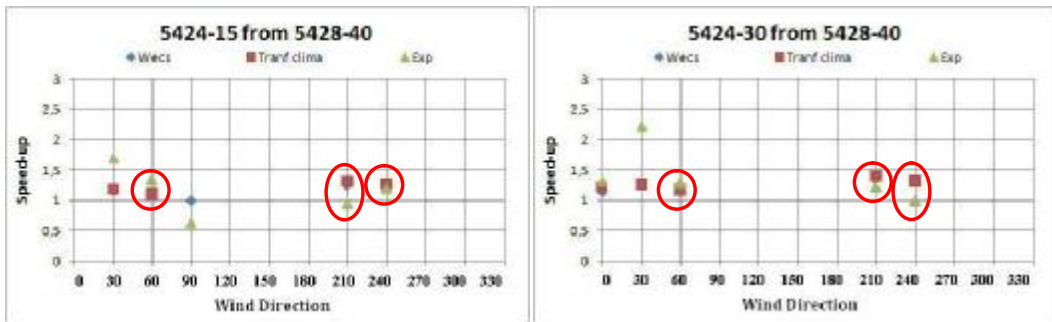


Figure 6.10 c/d

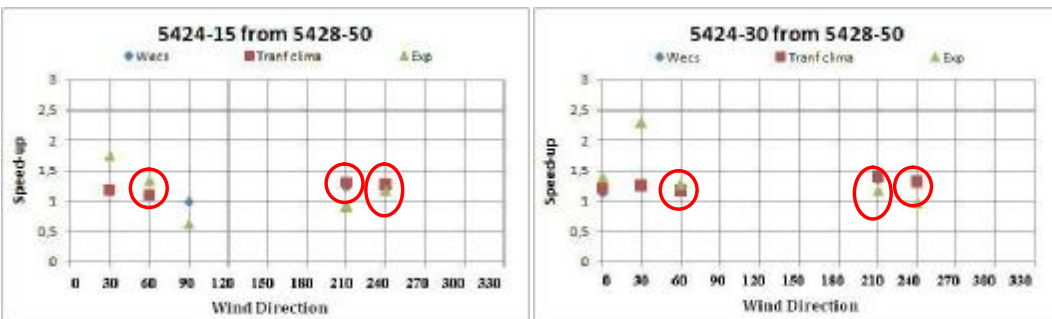


Figure 6.10 e/f

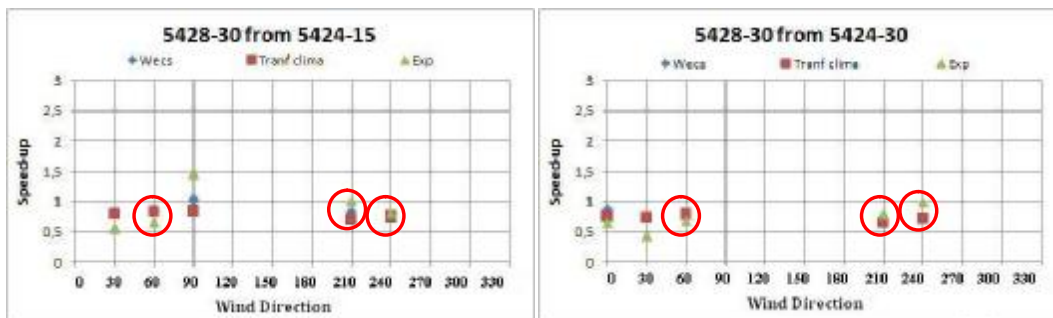


Figure 6.10 g/h

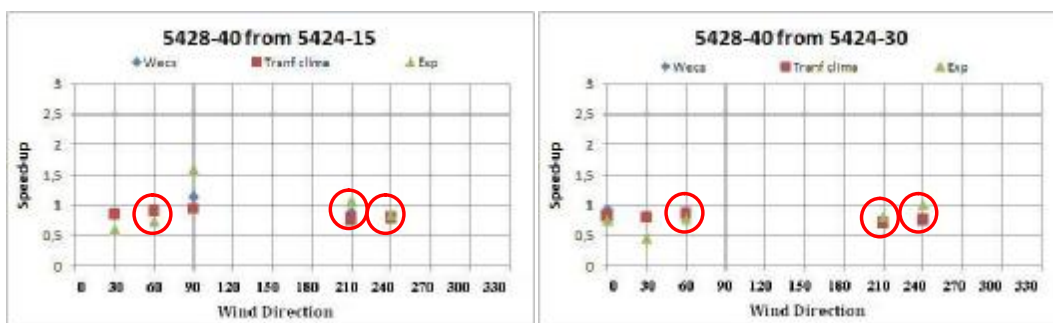


Figure 6.10 i/l

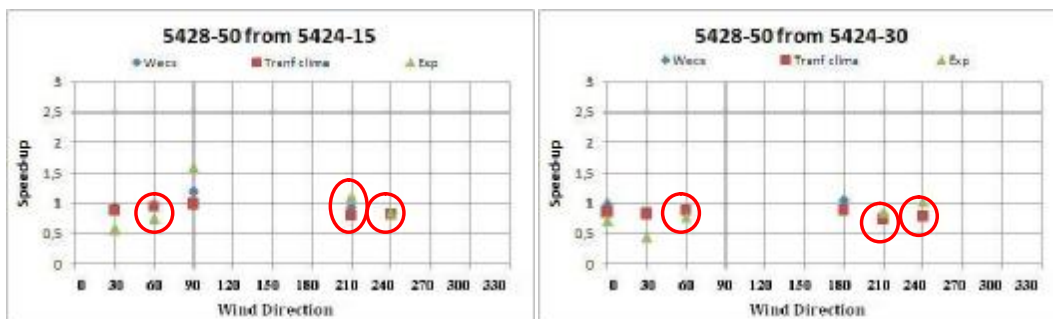


Figure 6.10 m/n

Tables 6.12 and 6.13 summarize the calculation of the RMS errors relative to the cross-checking procedure conducted by speed-up and reported in the figures 6.9 and 6.10.

Table 6.12: RMS errors for period April-May.

		Period April-May						
		NEW POSITION						
ORIGINAL POSITION	SPEED UPP	5424-30	5424-15	5428-50	5428-10	5428-30	Total	
		5424-30			0,122	0,123	0,121	0,122
		5424-15			0,080	0,074	0,071	0,076
		5428-50	0,208	0,135				0,172
		5428-10	0,215	0,122				0,158
		5428-30	0,247	0,133				0,190

		Period April-May						
		NEW POSITION						
ORIGINAL POSITION	SPEED UPP (no sector interpolation)	5424-30	5424-15	5428-50	5428-10	5428-30	Total	
		5424-30			0,138	0,138	0,124	0,133
		5424-15			0,100	0,093	0,082	0,092
		5428-50	0,211	0,134				0,172
		5428-10	0,219	0,122				0,171
		5428-30	0,243	0,133				0,191

Table 6.13: RMS errors for period November-December.

		Period November-December						
		NEW POSITION						
ORIGINAL POSITION	SPEED UPP	5424-30	5424-15	5428-50	5428-10	5428-30	Total	
		5424-30			0,169	0,165	0,180	0,172
		5424-15			0,206	0,195	0,210	0,204
		5428-50	0,435	0,255				0,345
		5428-10	0,227	0,254				0,241
		5428-30	0,280	0,318				0,299

		Period November-December						
		NEW POSITION						
ORIGINAL POSITION	SPEED UPP (no sector interpolation)	5424-30	5424-15	5428-50	5428-10	5428-30	Total	
		5424-30			0,155	0,151	0,161	0,156
		5424-15			0,156	0,145	0,153	0,152
		5428-50	0,209	0,234				0,221
		5428-10	0,215	0,228				0,222
		5428-30	0,264	0,287				0,275

The results of the cross checking by speed-up instead of mean wind velocity confirmed the results of the previous analysis. The period April-May was still better reproduced by CFD calculations and the best dataset was the one at station 5424 at 15 meters height. Moreover, the sector interpolation option was found to be helpful

to have a better agreement during the period April-May, while in the period November-December CFD results show worst match against experimental data. It must be noticed also that the sector interpolation function does not influence in a strong way the total error (shown in tables 6.12 and 6.13), which keeps the same order of magnitude both in cases April-May and November-December.

6.8 Remarks

A CFD simulation of the windfield of a site in central Italy was carried out. The chosen parameters to build the geometrical grid and to set the solver settings were chosen to reach a sufficient accuracy level without any divergence problem.

The velocity vertical profiles were studied and the agreement between CFD and experimental data was correct only in the dominant sectors where thermal effects were reduced.

Turbulence intensity was also studied, showing that the WindSim code underestimates the experimental data and that turbulence intensity level is strongly influenced by seasonal thermal effects.

The cross checking of the experimental data collected by the two masts measurement was finally performed. The best wind dataset in predicting wind experimental data was found in station 5424 at 15 meters height,

The “*sector interpolation*” option [4] was tested, demonstrating that it can be helpful in some case to find a better agreement between experimental and CFD data.

CHAPTER 7. Proposal for a plant layout.

7.1 Introduction

In the present chapter a proposal for a wind farm layout has been presented.

To do this, a wind resources map has been calculated by weighting the CFD results against the chosen experimental climatology, precisely station 5428 at 50 meters height.

According to literature [9], it is better to measure data at a vertical high distance from the ground in a zone not influenced by orographic effects. In this case the chosen station has all the required characteristics.

A wind park with 16 turbines was studied and the Annual Energy Production (AEP) yield was finally calculated

7.2 Dataset April-May 5428-50

The wind rose of the climatology 5428-50 during the period April-May is reported in figure 7.1, with the main parameters of the Weibull analysis.

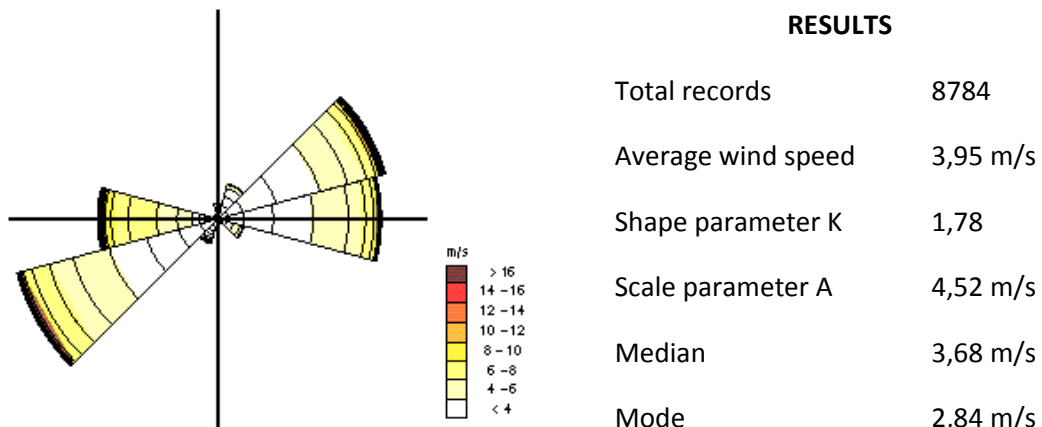


Figure 7.1: Climatology 5428-50 Apr-May

In the graph reported in figure 7.2 the Weibull fitting of the experimental data is presented, showing good agreement.

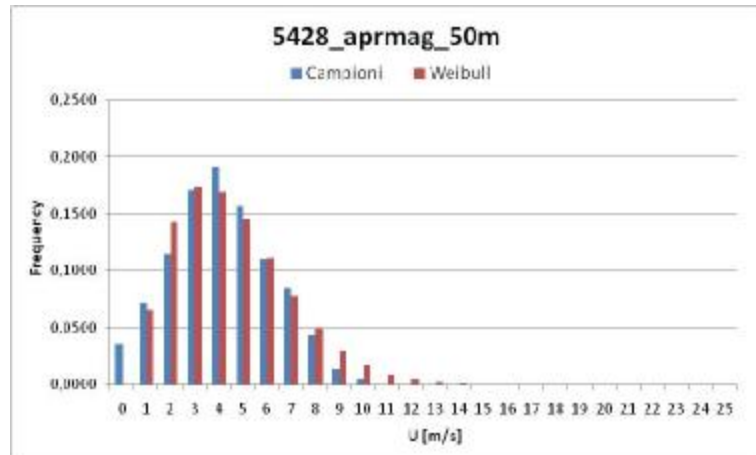


Figure 7.2: Weibull fitting of the climatology 5428-50 Apr-May.

7.3 Wind Resources

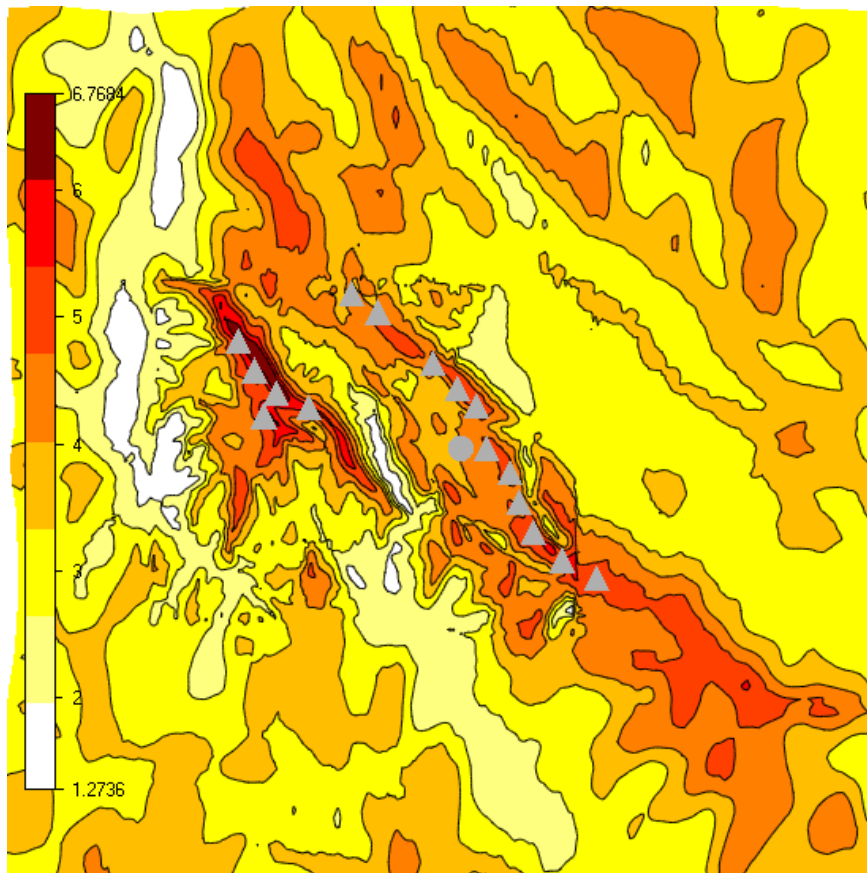


Figure 7.3: Horizontal wind velocity of the site at 60 meters height.

The figure 7.3 shows the horizontal wind velocity extracted in the site at 6 meters height.

As it can be seen the higher speed values are located in the zone of the mountains, surrounding the plateau where the stations were placed.

According to this consideration and the conclusions of the chapter 4, finding sectors 60 and 240 as dominant wind directions, 16 turbines have been placed in two rows perpendicular to the dominant wind directions.

The position of the turbines can be seen by the figure 7.4, while a perspective view of the wind park has been also presented in the figure 7.5.

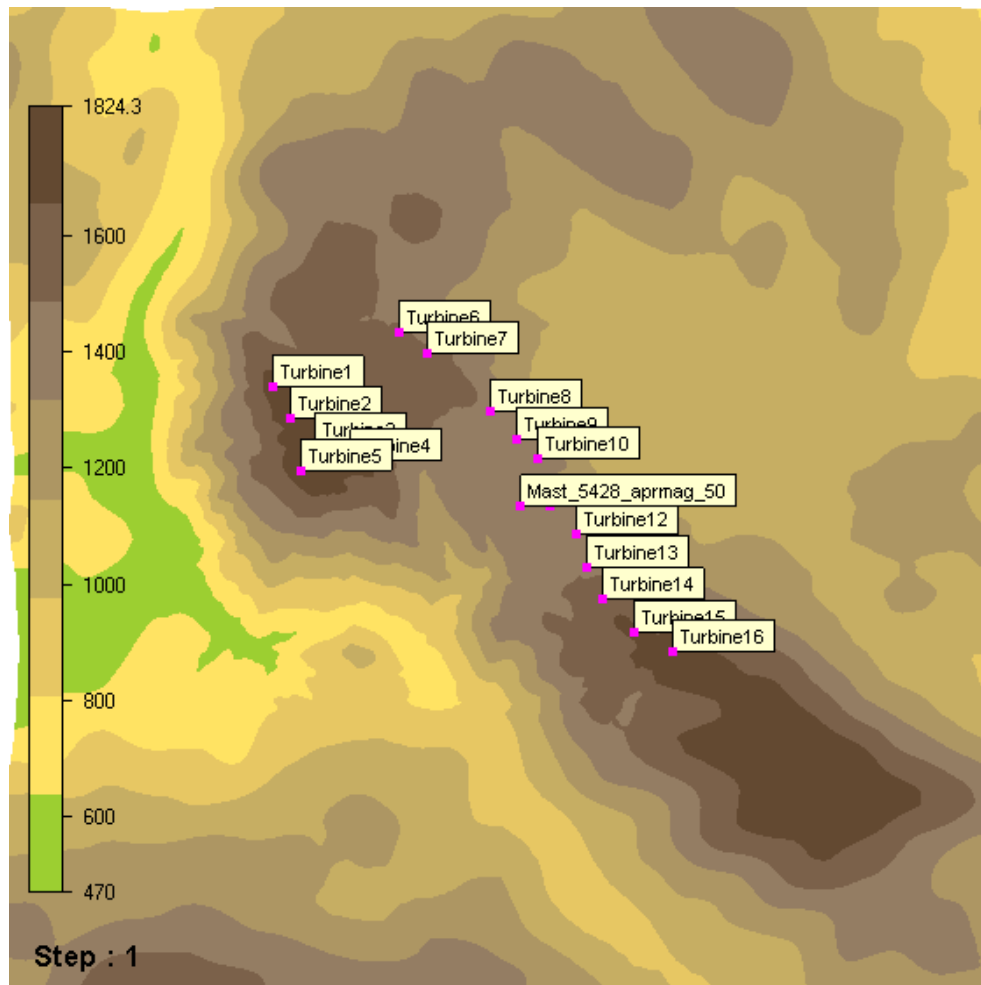


Figure 7.4: Orography map with wind turbines position.

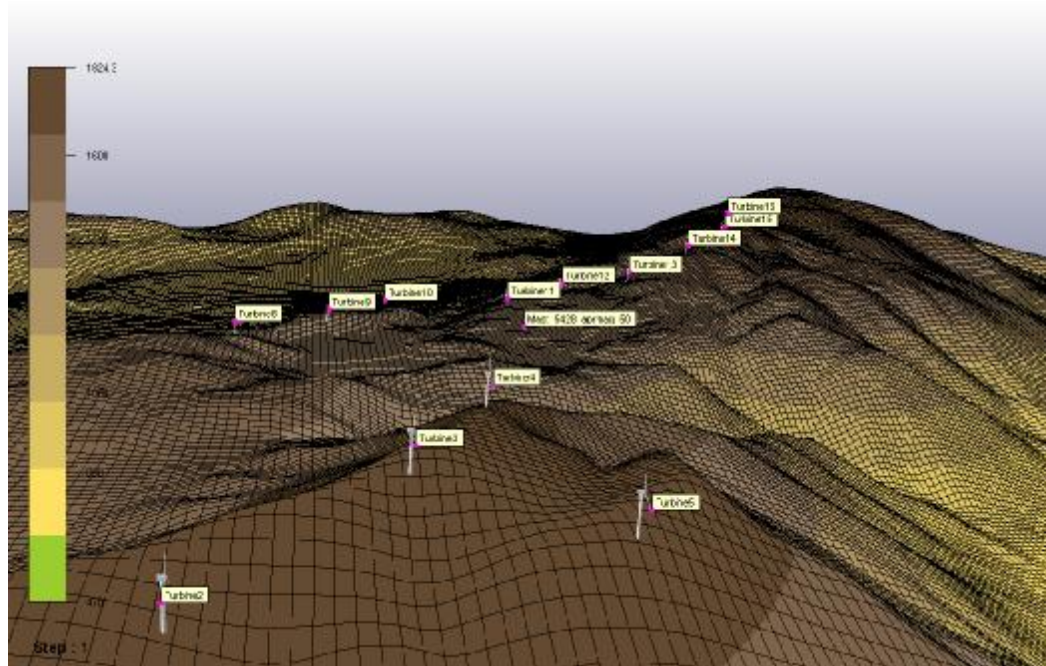


Figure 7.5: Perspective view of the wind park from West direction.

7.4 Annual Energy Production AEP

The turbine chosen to calculate the AEP is the *Gamesa G80* with 2000 kW of nominal power output. Its characteristics are summarized in table 7.1.

Table 7.1: Characteristics of the turbine Gamesa G80

Model	G80
Power	2000 kW
Rotor diameter	80 m
Number of blades	3
Lenght of the blades	39 m
Weight of each blade	6500 kg
Material of construction for blades	Glass fiber reinforced with epoxy resin
Cut in velocity	4 m/s
Cut off velocity	25 m/s
Hub height	60 m
Control system	The generator is a doubly fed machine (DFM), whose speed and power is controlled through IGBT converters and PWM control

A picture of the turbine is presented in figure 7.6.



Figure 7.6: Turbine Gamesa G80 – 2000 kW

Other technical data of the turbine are shown in figure 7.7, reporting the power curve.

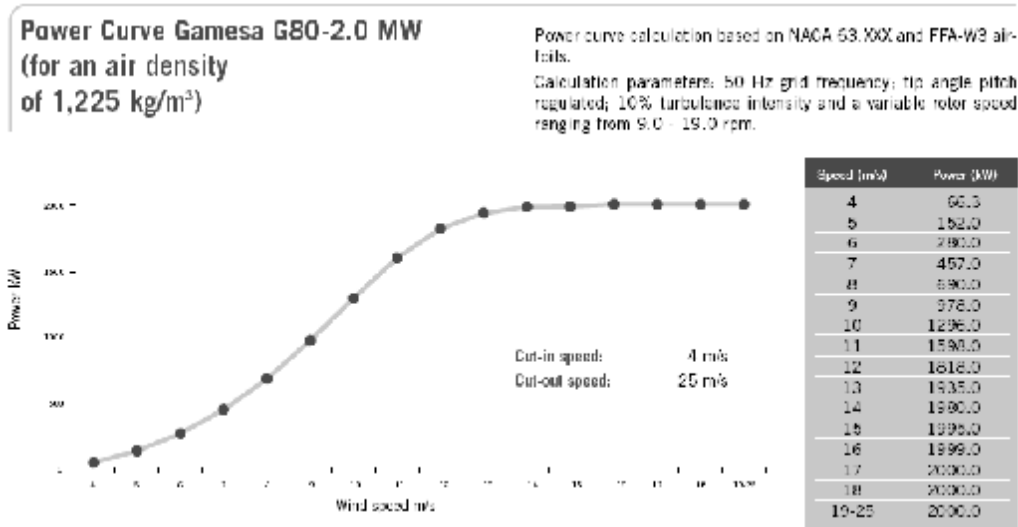


Figure 7.7: Power curve of the turbine Gamesa G80 -2000 kW.

The calculated AEP has been reported in the table 7.2: this table shows the energy production of every single turbine with its mean wind speed at hub height and the total amount of the wind park.

Table 7.2: Annual Energy Production of the wind park.

Turbine	Manufacturer	Model	Nominal power (kW)	Hub height (m)	Wind speed (m/s)	Energy (MWh/y)
Turbine1	Gamesa	G80	2000	60	6,50	4791,7
Turbine2	Gamesa	G80	2000	60	5,81	3776,7
Turbine3	Gamesa	G80	2000	60	6,34	4564,9
Turbine4	Gamesa	G80	2000	60	5,45	3149,7
Turbine5	Gamesa	G80	2000	60	6,06	4179,6
Turbine6	Gamesa	G80	2000	60	4,52	1996,1
Turbine7	Gamesa	G80	2000	60	4,52	1957,6
Turbine8	Gamesa	G80	2000	60	4,61	2018,7
Turbine9	Gamesa	G80	2000	60	4,89	2403,3
Turbine10	Gamesa	G80	2000	60	5,00	2539,6
Turbine11	Gamesa	G80	2000	60	4,68	2148,2
Turbine12	Gamesa	G80	2000	60	4,84	2360,4
Turbine13	Gamesa	G80	2000	60	4,49	1880,6
Turbine14	Gamesa	G80	2000	60	5,24	2877,7
Turbine15	Gamesa	G80	2000	60	5,30	2915,6
Turbine16	Gamesa	G80	2000	60	5,06	2600,5
All park			32000			46160,9

By dividing the AEP for the total nominal power output of the wind park , it is possible to calculate the equivalent hours of work, another parameter that indicates the production of the park, that is equal to 1443 hours per year.

7.5 Remarks

The AEP of a wind farm constituted by 16 turbines has been calculated. The position of the turbines was chosen considering the main wind direction sectors and the wind resource map, obtained by weighting the CFD results against the climatology 5428 (plateau) at 50 meters height in the period Apr-May.

The total energy production of the park was 46,2 GWh per year, corresponding to 1443 hours of work at nominal power output.

For further analysis it could be interesting to have a complete one year wind data set and to have wind data filtered with cases of atmospheric stability, in order to avoid strong thermal effects which WindSim was not able to reproduce.

Conclusions

In this thesis the evaluation of wind energy resources over a site in central Italy was conducted by means of the CFD code WindSim.

A series of 2D simulations were performed in order to validate the code. The code validation process brought to find the correct discretization parameter capable to guarantee grid-independent results. The “wall with no friction” boundary condition, to be applied at the top of the geometrical model was chosen, in order to describe a neutral stratified ABL. The comparison of the CFD results with numerical and experimental results found in literature showed a fair agreement.

The second part of the study concerned the evaluation of the wind resources o a site in Central Italy. The experimental wind data of two measurement masts (with multiple measurement stations) were reduced and statistically analyzed.

Taking advantage of the validation phase results, 3D simulations were carried out for a geometrical model reproducing the complex orography of the site. Firstly, a correct value of the free stream velocity of the geostrophic wind was found, to avoid convergence problems. A “nesting technique” in a coarse grid model was used to reduce the computational resources.

Finally the simulations over a 14×14 km extended model with fine grid resolution have been performed and the wind field over the whole site was calculated.

The results of the CFD calculations have been processed, in order to establish a comparison between experimental data and calculated values. The Vertical profiles of velocity and of turbulence intensity allowed the evaluation of RMS errors against the experimental data. RMS errors of velocity profiles exhibit a direct relation with the mean velocity of the directional sector, hereby showing a strong dependence on thermal effects. Indeed, these effects are not taken into account by the CFD simulations, whose numerical model is addressed to reproduce a neutral ABL. The analysis of results on the wind rose sectors allowed to find a range of mean velocity (3÷4 m/s) that discriminate whether or not data are influenced by thermal effects. For high mean velocity a fair agreement was evident.

A cross-checking procedure of wind data was also performed. Using each mast as a reference, the best agreement was found again for sectors where mean velocity is high and thermal effects where negligible. The discrepancy between CFD results and experimental data is acceptable because accurate ABL measurement are difficult to obtain. Furthermore, this kind of analysis is normally addressed to evaluate the correct wind resources distribution, instead of finding the true velocity field over the site.

For the site under study, a wind resources map at 60 meters height has been predicted and a proposal for a plant layout has been advanced, taking into account the dominant direction sectors and the higher horizontal speed areas.

As a conclusion, the methodology presented in this thesis is relatively straightforward and suitable for wind resources analysis. Nevertheless, the correct setting of the CFD model and long term wind measurement are needed to guarantee the correct accuracy. Fairly accurate results with neutral ABL simulations are obtained only for directional sectors where negligible thermal effects are found. This work demonstrates that the procedure is suitable for sites with complex orography, in this showing advantages over linear methods normally for siting evaluation.

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